

Degradation of groundwater quality in expanding cities in West Africa. A case study of the unregulated shallow aquifer in Cotonou

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1 Degradation of groundwater quality in expanding cities in West Africa. A

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- 3
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24 Abstract

25 In Cotonou, as in many expanding West African cities, major population growth and 26 infrastructural development has not kept up with informal settlement development onto floodable plains and marshes. The population of the slum, which makes up about 60 % of the city's inhabitants, 27 is highly disadvantaged and vulnerable to rising sea levels, flooding, sanitation and waste management 28 issues. However, the risks associated with the use of contaminated shallow groundwater for domestic 29 purposes are less recognised. Our study demonstrates that, in many instances, the cheaper option of 30 the use of shallow groundwater from the coastal Quaternary aquifer for domestic purposes represents 31 a significant risk for the poorer residents of Cotonou through the voluntary (drinking) or non-32 33 volontary (dish washing, cooking) consumption of this unmonitored and untreated water resource. In 34 the 3 neighbourhoods surveyed, environmental tracers (major ions, Cl/Br molar ratios and stables 35 isotopes) showed that this shallow aquifer is degraded by seawater intrusion as well as septic and 36 sewerage contamination. In particular, the higher NOx concentrations correspond to ranges associated 37 with sewerage and septic tank effluent pollution and the major ion concentrations and $\delta^2 H - \delta^{18} O$ 38 signatures showed that high salinity values are where groundwater mixes with saline Lake Nokoue 39 water. The population using this resource from local wells should be made aware of seasonal changes 40 in groundwater contamination and potential health risks associated with sewerage and septic tank contamination. 41

42 Key words: Groundwater, sewerage and septic tank contamination, salinization, urbanization,

43 recharge and discharge.

44

45 **1. Introduction**

With cities in developing countries expanding at unprecedented rates, ensuring clean water supplies for all inhabitants is becoming increasingly more challenging. This is particularly the case in more impoverished urban areas, where infrastructure development often lags behind population growth (Lapworth et al., 2017). In addition, there often exist two types of water supplies: firstly, there is the official water supply that is monitored, treated and comes at a cost for consumers; and then there is the second unregulated water supply that is often sourced from wells accessing shallow groundwater that is unmonitored, untreated but free.

53 The health risks in using the unregulated groundwater resource for domestic purposes, including drinking water supplies, is high due to water quality issues in many urban areas (e.g. 54 Ouedraogo et al., 2016; Hassane, 2010). The same factors driving the demand for water supplies, 55 accelerated urban growth and the expansion of informal settlements, are also significant drivers of 56 57 groundwater pollution (UNESCO, 2017). One of the many threats to shallow groundwater quality in cities is from sanitary wastewater and solid waste disposal (Lu, 2015). Uncontrolled seepage of 58 wastewater from septic tanks and human activities as well as infiltration of urban stormwater lead to 59 groundwater contamination (Dhanasekarapandian et al., 2016). Many parts of the world are now 60 61 reporting groundwater and surface water nitrate pollution issues (Roy et al. 2007; Stuart et al. 2007; Zhang et al., 2014; Ogrinc et al., 2019). For example, in the Coimbatore city, India, population 62 growth, pit latrines and septic tanks, industrial effluents, and irrigation water return flows are the main 63 64 sources of groundwater contamination (Selvakumar et al. 2017). In Florida, the proximity of wells to 65 septic tanks contributes to increasing fecal coliform, nitrate and phosphate concentrations during wet season compared with the dry season (Arnade, 1999). Many studies have identified sewerage and 66 latrine contamination as a significant public health issue due to the resultant presence of faecal 67 bacteria in waters such as Escherichia coli, faecal Streptococci, Salmonella and Shigella (Boukari, 68 69 1998; Odoulami et al. 2013; Yadouléton, 2015). During this study, for the first time, the seasonal 70 variation in the contamination of waters by sewerage and latrine contaminantion is investigated using 71 environmental tracers.

In addition, shallow groundwater resources in coastal cities are particularly vulnerable to salinity problems due to mixing with saline surface waters and seawater intrusion (Barker et al., 1998; Cary et al., 2015; Petelet-Giraud et al., 2016; Najib et al., 2017; Liu et al., 2017). For example, in Ho Chi Minh city, Vietnam, the groundwater resources are under threat due to saltwater intrusion in the shallow aquifer (Ngo et al., 2015).

77 In this study, we use the example from Cotonou, Benin, to identify processes resulting in nitrate and salinity contamination of a city's unmonitored and untreated shallow groundwater 78 resource. The complexity of contamination processes in evolving urban environments makes it 79 difficult to study the risks of shallow groundwater quality. This can be due to both natural seasonal 80 81 shifts from climatic and environmental influences, and significant anthropogenic influences on the 82 hydrogeological system. In this study, we analyse the distribution and concentration of major ions and stable isotopes (δ^2 H and δ^{18} O) to investigate the temporal variation in groundwater quality across 3 83 84 sites in Cotonou. These 3 neighbourhoods were selected to represent contrasting hydrogeological 85 environments: (i) a neighbourhood (St Jean) where there is no surface inundation; (ii) a 86 neighbourhood (Ladji) bordering Lake Nokoué which overflows towards the end of the wet season; 87 and (iii) a neighbourhood (Agla) located in a swamp low land and is subject to inundation during both the small and large wet seasons. In doing so, we identify hydrological conditions and seasons when 88 groundwater is at its greatest vulnerability in terms of salinization and latrine contamination. At each 89 90 site, there are wells where local inhabitants have access to unmonitored and untreated shallow 91 groundwater for domestic consumption.

92

93 2. Study Area

94

95 2.1 Location and climate

The city of Cotonou is bordered by Lake Nokoué to the north, and the Atlantic Ocean to the south. The location of Cotonou and the average monthly interannual rainfall at Cotonou station from 1971 to 2016 are shown in Figures 1 and 2. The topography of the city is relatively flat with an altitude varying between 0 m and 6 m (Boukari, 1998). The average annual rainfall for Cotonou is 100 1,300 mm (Yabi et al. 2006). The city's climate is characterised by 4 seasons; main dry season
101 (average rainfall is 25 mm from mid-November to mid-March), main wet season (average rainfall is
102 152 mm from mid-March to mid-July), small dry season (average rainfall is 55 mm in mid-July to
103 mid-September), and small wet season (average rainfall is 75 mm from mid-September to mid104 Novembre).

105

106 **2.2 Aquifer geology**

Cotonou is located in the coastal sedimentary basin comprised of Quaternary (Holocene) 107 sediments, which include facies of the littoral plain (sands) and alluvial deposits, underlain by 108 sediments from the Mio-Pliocene (Continental Tenninal), Paleocene and Upper Cretaceous (Maliki, 109 110 1993). Previous studies by Oyédé (1991), Maliki (1993), Alidou (1994) and Boukari (1998) have described in detail the different sedimentary units of this coastal sedimentary basin and the 111 distribution are presented in Figure 1. The shallow Quaternary sediments consist predominantly of 112 fine to medium sands (85.5%), silts (5.5%) and clays (9%) (Oyédé, 1991). The low clay content in 113 these sediments results in the high permeability of the shallow sandy soils (Maliki, 1993), and 114 therefore a high vulnerability of the aquifers to the transfer of surficial pollutants. The clays in 115 Quaternary sediments are constituted by smectite and a kaolinite content, which increases with depth. 116 The traditional domestic wells were sampled in this study so that we could analyse the water 117 consumed and used by local residents. However, due to the intense pumping of these wells, the 118 fluctuations in hydraulic head are heavily influenced by pumping, therefore we were unable to use the 119 hydraulic head data to infer subsurface flow directions. There has however been previous work that 120 has addressed this question, the work of Maliki (1993) and Boukari (1998) focused on the piezometry 121 of the Cotonou water table during the flood and low water periods, and the interactions with the Lake. 122 It appears from these studies that during the flood period, there is a flow direction of groundwater 123 from the center of Cotonou (piezometric dome) to outlets such as Lake Nokoué, the lagoon, the 124 Atlantic Ocean and swamps. In comparison, during the low-water period it was observed that water 125 from the Lake Nokoué flows into the groundwater aquifer 126

127

128 2.3 Sewerage and waste

129 The city of Cotonou has about 166,433 households and the population evaluated at 679 012 130 inhabitants in 2013 (INSAE, 2015) and is subject to increasing urbanization. The poorer dwellings are located in the neighborhoods along the coastline, along the edges of the Cotonou Channel and Lake 131 Nokoué, and in the swamp areas. In these neighbourhoods, solid and liquid wastes are released into 132 the immediate environment without treatment as illustrated in the Figure 1(c). In addition, they are 133 also scattered in various places of the city where garbage heaps have formed. 78.5% of household 134 135 wastewater and 33 % of solid wastes are ejected in gardens, streets, gutters, unused wells and empty 136 blocks (INSAE, 2016). These poorer neighborhoods are also without adequate sanitation systems. In 137 Cotonou 64.9% of households use latrines that are reportedly leak-proof, whereas 13.5% adopt unsafe 138 and non-hygienic practices such as still latrines and open defecation. Only 20.8% of the population use septic tanks, which are generally evacuated by drainage structures (INSAE, 2016). Since there are 139 140 few market gardens in this heavily urbanised city, the potential sources of nitrate in the study area are predominantly from wastewater, solid waste and septic tanks. Therefore, the risk of pollutant transfers 141 142 to the shallow aquifers is high in the poorer neighbourhoods because they are exposed to both (i) frequent episodes of inundation and (ii) high levels of pollutants from poorly constructed landfills and 143 144 latrines.

145 **2.4 Groundwater quality and use**

Several studies have evaluated the impact of human activities on the quality of the shallow 146 groundwater in the Cotonou region and have reported relatively high levels of nitrates (e.g. up to 100 147 mg/L; Maliki, 1993; Boukari, 1998; Odoulami et al. 2013; Totin et al. 2013). In addition, the 148 vulnerability of this groundwater resource is also related to its proximity to the Atlantic Ocean and the 149 150 saltwater or brackish lakes (Boukari, 1998; Totin et al. 2013). Data from previous studies have shown an increase in salinity levels in groundwater near Lake Nokoué. The salinity of the Lake Nokoué also 151 exhibits a marked seasonal cycle (0 psu during the small wet season in October and 25-30 psu at the 152 end of the dry season) (Stephen et al. 2010; McInnis et al. 2013; Totin et al. 2013). The coastal 153 groundwater system in Cotonou has relatively high chloride (23.6 to 160 mg/L) and sulphate (6.4 to 154 155 25.7 mg/L) concentrations (Alassane et al. 2015; Nlend et al. 2018). Lake Nokoué and the Atlantic

Ocean are therefore a major concern in terms of the source of salinization of groundwater resources inthe coastal zone (Alassane et al. 2015).

Due to the groundwater pollution problems, shallow groundwater is excluded from the official supply of drinking water to the city of Cotonou. However, the low rate of services to the public water supply network in the neighborhoods of peri-urban areas (51%) leads the population to use the groundwater from the shallow Quaternary aquifer for their various domestic uses (INSAE, 2016). This is despite the fact that the Quaternary aquifer is not part of the official monitoring services of the National Water Company of Benin, the only legal water distribution structure.

164 Groundwater from the shallow Quaternary aquifer, is accessed by the residents of Cotonou via 165 large-diameter wells installed on private household property. These large-diameter wells are shallow 166 and are designed to pump groundwater from the water table surface so as not to pump deeper saline 167 groundwater (Maliki, 1993). The Quaternary aquifer sands have a high permeability (in the order of 168 10^{-2} to 10^{-4} m/s) and locally contains fresh to brackish groundwater whose exploitation is related to the 169 position of the lake water and seawater instrusions, and to the replenishment of freshwater lenses.

In Cotonou, 81% of the neighbourhoods have wells, and whilst 9% have not any water supply
(no wells nor stand pipe for drinking water supply) for domestic purposes, they sometimes buy water
for drinking purposes from neighbours (Hounkpè et al., 2014). Furthermore, the groundwater from the
large-diameter private wells is used in households for drinking, cooking, laundry, bathing and washing
dishes (Yadouléton, 2015; EAA, 2018).

175 In this study, we expand on previous work to elucidate the seasonality of hydrogeological 176 processes resulting in the contamination of groundwater from latrine waste (notably nitrate and nitrite 177 concentrations, herein referred to as NOx) and from salinization.

178

179 **3. Methods**

180 Three neighbourhoods in Cotonou were selected for sampling based on differences in their 181 hydrogeological environments: (i) St Jean, where there is no long-term surface inundation and the 182 heavy rainfall infiltrates the soils or leaves the site via overland flow; (ii) Ladji, which is located at the 183 shore of Lake Nokoué and is inundated towards the small wet season; and (iii) Agla, which is located 184 in a flood plain and is inundated early during both the small and large wet seasons. The locations of these three neighbourhoods and sampling sites are presented in Figure 1. In each neighbourhood, 185 groundwater wells were selected for sampling and, where/when possible, surface water samples 186 187 (permanent and temporary pools, as well as Lake Nokoué) were also included. The depths to the water table and the electrical conductivity (EC) of groundwater in the Quaternary aquifer were monitored 188 monthly at 9 wells and 3 piezometers from each of the 3 neighbourhoods from June 2017 to June 189 2018. The wells are separated by an average distance of ~ 200 m within the same neighbourhood 190 probe. 191

192 The hydrochemistry of waters were analysed 6 times between November 2016 to June 2018 193 during the months of February (main dry season), June (main wet season) and November (small wet season), resulting in a total of 127 groundwater, and 60 surface water samples. In addition, for the 194 period from June 2017 to June 2018, 13 rainfall samples were analysed to determine water stable 195 196 isotopes (δ^2 H and δ^{18} O). Lastly, 2 water samples from atlantic ocean (AO) were collected of the 197 periods from January 2019 (main dry season). Groundwater was sampled from the Quaternary aquifer via the large-diameter wells. Groundwater and surface water temperature, EC, and dissolved oxygen 198 (DO) were measured *in-situ* using a WTW 3430i portable digital multiparameter. 199

Water samples were collected for the analyses of cations (filtered at 0.45 µm and acidified 200 with HNO₃), anions and stable isotopes, that were conducted at the Laboratory of Hydrogeology of the 201 University of Avignon, France. Major ions were analysed using ion chromatography (Dionex; 202 203 ICS1100 and autosampler AS-AP). Ion analysis uncertainty is in the order of 3 %, and all ionic 204 balances were ≤ 5 %. The alkalinity was measured using a HACH digital titrator, and stable isotopes 205 were analysed using a Picarro Analyser L 2130-I. For the stable isotopes of waters, the error is $\pm 0.1\%$ for δ^{18} O and $\pm 1\%$ δ^{2} H. The results are presented in Tables 1 and 2. The mixing ratio between lake and 206 207 shallow groundwater is calculated by using stable isotopes values according to Paran et al. (2015):

208 $\delta^{18}O_m = f \cdot \delta^{18}O_L + (1-f) \cdot \delta^{18}O_{GW}$.

The same equation is used with $\delta^2 H$ to compute the $\delta^2 H_m$ mixing ratio. O_m is water sample isotope composition where groundwater (GW) and Lake (L) mixing is supposed. O_L is Lake isotope composition. O_{GW} is groundwater isotope composition and *F* is mixing fraction.

One questionnaire in each sampling site was administered, 10 per study area (Ladji, Agla and St. John). A total of thirty (30) questionnaires were administered through an interview with households in each study area. Main information such as the the source of drinking water supply, groundwater use, waste and wastewater management are mentioned in the questionnaire (supplementary data).

217

218 **4. Results**

219 4.1 Seasonal variations of the water table and groundwater EC

The monthly monitoring of the depth to water table and EC over a one-year period highlights the variations in seasonal fluctuations of the aquifers between the sites. Typical of unconfined shallow aquifers, the Quaternary aquifer shows marked spatial and seasonal variations and is in phase with the monthly rainfall. The seasonal variations of the water table and groundwater EC in each of the neighbourhoods are illustrated in Figures 3 and 5.

225 Between the neighbourhoods, St Jean has a maximum depth to water table (1.5 m at site J5) corresponding with the dry season (February) that is relatively deep compared with the depths of 1.4 226 m and 1.0 m observed in Ladji (site L7) and Agla (site A4), respectively (Fig. 5). In addition, at 2 of 227 228 the 4 sites in Agla, the negative depth values indicate that the water table is above the land surface 229 during the wet season (thus contributing to the floods), whereas groundwater at St Jean and Ladji remains below the surface. The seasonal fluctuation amplitudes also vary. In St Jean, the maximum 230 change between wet and dry season water table depth is of 1.2 m. Here the water table fluctuations 231 also vary between wells. Well J5 strongly reacts to the increased rainfall in October (small wet season) 232 233 whereas the deepening of the water table in wells J1 and J2 remain relatively weak (Fig. 5). These relatively low water table values can be attributed to the use of wells by households (Boukari, 1998; 234 Kadjangaba et al., 2018). Similar to St Jean, in Ladji, the seasonal fluctuations in the water table 235

between the wet and dry seasons vary by 1.3 m (Fig. 5). In comparison, in Agla, the seasonal amplitude of the water table variation is lower (0.5 m). Since Ladji is located on the shore of Lake Nokoué, the lake fluctuations likely influence the variation of the water table depths. An increase of about 90 cm in the water table is observed at Ladji between the wet and dry seasons is consistent with the observed elevation of the lake level during the same period (IRD/IRHOB, unpublished data). The rise in water level starts in January and increases gradually to April/May, which is a 2-month time lag from Agla and St Jean.

243 With seasonal variations in the water table depth, the groundwater EC also shows varying trends (Fig. 3). In St Jean, the groundwater EC values (297-1,285 µS/cm) are relatively weak 244 245 compared to Agla and Ladji, and show variable seasonal changes (Fig. 3). Wells J1 and J2 recorded 246 higher EC values in the dry season (511-1,285 µS/cm) when groundwater levels were low compared with the wet season (297-799 µS/cm). In comparison, the well J5 and the piezometer had higher EC 247 values (444-1,169 µS/cm) in the rainy season (rise of the water table) compared with the dry season 248 (497-685 µS/cm). In Agla, the groundwater EC ranges between 353-1,468 µS/cm, which is low 249 250 compared with local surface waters (including temporary ponds and swamps; 2,800-2,900 µS/cm). EC 251 values are similar during the rainy season and the dry season (Fig. 3). In Ladji, the groundwater EC values are higher than in the other neighbourhoods (633-5,340 µS/cm) and similar to the local 252 temporary ponds (480-4,090 μ S/cm), but strongly lower than the Lake Nokoué (up to 49,700 μ S/cm) 253 254 during dry season). With the exception of well L2, no difference was observed in the variation of the EC during the dry season and the rainy season $(1,478-5,340 \,\mu\text{S/cm})$. 255

256

257 4.2 Seasonal variations of major ions

The composition of major ion concentrations of waters is different in each of the neighbourhoods as highlighted in Figure 4. In St Jean, the groundwater is of Na-Ca-HCO₃-Cl type. In comparison, the groundwater in Ladji is of Na-HCO₃-Cl type, and Agla groundwater is more of the Na-Ca-HCO₃-SO₄-HCO₃ type (Fig. 4). The ponds sampled in both Ladji and Agla show some similarities in major ion compositions with local groundwater, thus indicating a potential connection between surface and subsurface waters. However, this connection is spatially and/or temporally variable, in particular in Agla, since some of the pond samples have a greater concentration of HCO_3 relative to Cl and SO₄, and increased Ca relative to Na and Mg compared with groundwater. In addition, in Ladji, the Lake Nokoué has a greater ratio of Cl to SO₄ and HCO₃ compared with groundwater.

268 The temporal evolution of these major ions at each site is presented in Figure 5 and, similar to EC values, the seasonal changes vary both between sites and between neighbourhoods. Greatest 269 seasonal variations are observed in Ladji, and particularly for groundwater sampled at L6 and L7 270 where dry season conditions results in increases in Na (by 244-374 mg/L), Ca (by 76-81 mg/L), Cl (by 271 301-447 mg/L) and HCO₃ (by 530-622 mg/L). These sites are located between 100 to 150 m from the 272 273 Lake shore. The increase of the ions' concentration during dry season is associated with a deepening 274 of the water table and may therefore indicate influences of lake infiltration. The Lake Nokoué has 275 greater concentrations of Na (3,168-12,219 mg/L), Ca (146-419 mg/L), Cl (6,166-1,9310mg/L), and 276 HCO₃ (73-110 mg/L) compared with all groundwater samples (Tables 1 and 2).

277 In Agla, the seasonal variations in major ions remain relatively stable except for HCO₃ and SO₄ concentrations (Fig. 5). However, between sites, there are large differences in seasonal trends. 278 Four of the eight groundwater wells (A4, A5, A9 and A10) show an increase in Na (by 47-128 mg/L), 279 280 Ca (by 20-88 mg/L), HCO₃ (by 155-750 mg/L) and Cl (by 61-155 mg/L) during the month of February (dry season). During the wet seasons (June and October/November), the remaining four 281 wells (A2, A6, A8 and A7) show an increase in Na (by 93-158 mg/L), Ca (by 34-78 mg/L), HCO₃ (by 282 99-436 mg/L), Cl (by 110-176 mg/L), and SO₄ (by 75-276 mg/L) concentrations. These trends 283 284 correspond to changes in the water table depths, where the increasing wet season major ion 285 concentrations occurs in wells where the groundwater is (sub)artesian during the wet season (A2 and A8), compared to wells where groundwater remains well below the surface and has increased dry 286 287 season major ion concentrations (A4). These trends may highlight the influence of groundwater 288 mixing with pond water in the discharge areas, compared with infiltrating rainfall diluting groundwater in other areas during the wet season. The temporary and permanent pond waters have 289 similar concentrations of Na (56-166 mg/L), Ca (34-147 mg/L), Cl (43-240 mg/L), and HCO₃ (97-738 290 291 mg/L) compared with Agla groundwater samples (Tables 1 and 2).

292 In St Jean, groundwater at seven of the ten wells sampled (J1, J2, J10, J3, J4, J6 and J9) show an increase in Na (by 36-115 mg/L), Ca (by 49-114 mg/L), HCO₃ (by 119-439 mg/L) and Cl (by 42-293 294 111 mg/L) concentrations during the dry season (February and March, Fig. 5). This occurs at the time 295 when the water table is at its deepest level. In addition, small wet season peaks in major ion 296 concentrations are also observed in groundwater from wells J5, J7 and J8. These increases include Na (by 63-65 mg/L), Ca (by 63-80 mg/L), HCO₃ (by 128-272 mg/L) and Cl (by 61-79 mg/L) 297 concentrations. Since St Jean does not have any surface water bodies, the wet and dry season major 298 ion variations are probably due to variations in the composition of infiltrating water, mixing between 299 subsurface waters, reactions during water table fluctuations, and evaporation. 300

301

302 4.3 Stables isotopes

The δ^{18} O and δ^{2} H values of rainfall in Cotonou (obtained from 95 samples from the IITA and 303 304 IAEA Global Network of Isotopes in Precipitation (GNIP) from 2005-2016) has a weighted average of -2.8 ± 1.6 % for δ^{18} O and -11.2 ± 11.7 % for δ^{2} H. The local meteorological water line (LMWL) of 305 rainfall obtained from these data has a slope of ~7 and is presented together with the global meteoric 306 307 water line (GMWL; slope of ~8, Craig, 1961) in Figure 6. These data were also used to calculate the mean weighted average values of δ^{18} O and δ^{2} H for different ranges in monthly rainfall. The results 308 show that when monthly rainfall is < 50 mm, the isotope values are greater compared with monthly 309 rainfall ranging from 50 mm to 500 mm (Fig. 6a). The isotope values of rainfall collected in Cotonou 310 show an overall mass effect; a depletion in heavy isotopes correlates with an increase in rainfall 311 amounts. In Cotonou, the local rainfall has stable isotope values covering a large range; from -5.4 to 312 0.8 % for δ^{18} O, and from -30.9 to 14.5 % for δ^{2} H (Fig. 6a, black crosses). The groundwater values 313 from the three sites are also presented in Figure 6a, which also highlights a large range in values (from 314 -4.17 to 0.35 % for δ^{18} O, and from -21.9 to 4.3 % for δ^{2} H). 315

In St Jean, the stable isotope values for groundwater vary between -4.17 and -2.37 ‰ for δ^{18} O and between -21.87 and -8.98 ‰ for δ^{2} H. These values show that the groundwater in St Jean is depleted at δ^{18} O and δ^{2} H. The linear regression between δ^{18} O and δ^{2} H values (slope of 7.5) lie close to the LMWL (Fig. 6a), suggesting that groundwater likely originates from local rainfall.

In Ladji, groundwater has enriched in δ^{18} O and δ^{2} H compared with St Jean (-3.36 to 0.35 and 320 -12.96 to 4.17 ‰, respectively; Fig. 6b). This discrepancy may be due to three distinct phenomena: (i) 321 322 the recharge of relatively low volume rainfall events (< 50 mm/month) in Ladji compared with St 323 Jean; (ii) evaporation effects (some values lie to the right of the LMWL indicating a slope of 5.2), or (iii) groundwater mixing with lake water. During the dry season, groundwater isotopic values are 324 close to the lake Nakoue values (Fig. 6b), which suggests a strong mixing between the lake and 325 groundwater during this season. There are also a number of groundwater samples whose δ^{18} O and δ^{2} H 326 values remain close to those of the temporary and permanent ponds. δ^{18} O and δ^{2} H mixing ratios 327 indicate ranges from 72 to 74 % of shallow groundwater mixing with lake water during the dry season 328 in Ladji. 329

Like in Ladji, groundwater in Agla is also enriched in δ^{18} O and δ^{2} H (-3.57 to -0.03, and -18.22 to 2.78 ‰, respectively) compared with St Jean. Some samples from Agla (mostly during the dry season) lie to the right of the MWL with a slope of 5.4 (Fig. 6c). Therefore, these groundwaters may also be subject to evaporation and mixing with enriched surface waters such as the lake water (via the lowlands) and the pond water. In addition, groundwaters (e.g., wells A9 and A10) that are close to lake water are located in discharge areas. δ^{18} O and δ^{2} H mixing ratios indicate ranges from 54% to 62% of shallow groundwater mixing with lake water in Agla.

337

338 4.4 Nitrogen

The Beninese quality standard for drinking water for nitrogen is 45 mg/L (0.70 mmol/L) for NO₃, 3.2 mg/L (0.07 mmol/L) for NO₂, and 250 mg/L (7.0 mmol/L) for Cl (Decree N ° 2001-094 of February 20th). In this study, 26 % (values of 0.73-5.06 mmol/L) of the groundwater sampled exceeds these limits in terms of NO₃, and 7.0 % (values of 0.12-2.04 mmol/L) for NO₂ (Table 1).

Highest values of NOx (\sum NO₂+NO₃) are observed in groundwater of St Jean, where the concentrations reach values of 5.06 mmol/L (average of 0.90 mmol/L). For NH₄, the groundwater concentrations remain low (up to 0.80 mmol/L, average of 0.04 mmol/L). The highest concentrations of NOx are observed during the dry season (> 3 mmol/L) compared with wet season samples (NOx < 347 3.37 mmol/L). Two sites (J6 and J8) have elevated NOx concentrations in both the wet and dry
348 seasons (1.15-5.06 and 0.99-2.58 mmol/L, respectively). NOx in St Jean groundwater may originate
349 from natural fixation of nitrogen in the soil and human pollution via leaky latrines and human
350 defacation in the streets.

In Ladji, the groundwater NOx concentrations ranges up to 2.51 mmol/L (average of 0.54 351 mmol/L), and NH₄ is up to 0.43 mmol/L (average of 0.11 mmo/L). During the wet season, the 352 groundwater NOx concentrations are highest (1.06-2.51 mmol/L) compared to dry season 353 concentrations (<0.13 mmol/L). The wet season increase in NOx of groundwater in Ladji may either 354 be due to infiltration of anthropogenic pollution, or nitrogen fixed by vegetation in the soils. The NOx 355 356 concentrations are low in the permanent ponds and the lake (<0.01 mmol/L). However, the wet season 357 temporary ponds have high concentrations of NH₄ (up to 1.55 mmol/L), which may undergo 358 nitrification during infiltration and may also result in the NOx contaminantion of groundwater.

The highest value of NH₄ in groundwater is observed in Agla, but only at one site (A10: 4.91 mmol/L). All other sites have NH₄ concentrations in groundwater ranging from 0.00 to 0.52 mmol/L (average of 0.17 mmol/L). Compared with Ladji and St Jean, groundwater NOx concentrations in Agla are low, with a maximum of 0.86 mmol/L (average of 0.15 mmol/L). In contrast, permanent ponds in the Agla lowlands have high NOx levels only at one site (A13: 17.50 mmol/L), particularly in the dry season.

365 As described by Katz et al. (2011), the Cl/Br ratio can be a valuable first assessment of septic tank contamination of shallow groundwater. This is based on the assumption that sewerage waters and 366 septic tank effluent exhibit distinct ranges and higher values of Cl/Br molar ratios and Cl 367 concentrations compared with rainwater (Davis et al. 1998), as observed in many samples form this 368 369 study (Fig. 7a). Elevated Cl/Br ratios may also result from the dissolution of halite. However, halites 370 have not been reported in the local Quaternary aquifer. In addition, many of the groundwater values 371 correspond to ranges reported for sewerage and septic tank effluent (Davis et al. 1998; Vengosh and Pankrativ, 1998; Katz et al. 2011) compared with waters with halite dissolution (Davis et al. 1998; 372 Pastén-Zapata et al. 2014; Panno et al. 2006), rainwater from coastal areas, and seawater (Alcalá and 373 374 Custodio 2008; Davis et al. 1998). According to Figure 7b, groundwater values with similar Cl/Br

375 ratios and Cl concentrations than sewerage wastewater also show higher concentrations of NOx. Three 376 samples (wells J2, L9 and A6) plot above the Cl/Br ratios for sewage or septic tank effluent (Cl/Br molar ratios 676-1350; Cl <1000 ml/L), likely indicating animal manure. As evidenced in Figure 7b, 377 values that fall into the animal manure/animal urine (Cl/Br molar ratios, 2,810-3,730; Cl <1000 mg/L) 378 379 also have high NOx concentrations (0.76-0.81 mmol/L). Other evidence for a septic tank and animal manure influence are the corresponding elevated dissolved organic carbon (DOC) concentrations 380 (Table 1). However, not all the groundwater samples have high DOC concentrations. According to 381 Broun et al. (2009), rapid oxidation of DOC into carbon dioxyde may account for the low DOC 382 concentration in groundwater. 383

384 Alternatively, samples with low NO_x concentrations and relatively low Cl/Br ratios, such as 385 groundwater from Agla (Fig.7b), can indicate areas of organic biodegradation (e.g. McArthur et al., 2012). The relationship between HCO₃, NO_x and DO for all groundwater sampled at Agla is 386 387 shown in Figure 8. The samples where dissolved oxygen (DO) and NOx are low compared with HCO₃ correspond to discharge areas. This is especially the case for groundwater sampled during the dry 388 season at wells A10, A9 and piezometers A11, J11 where HCO3 values increases (4.80-12.31 389 mmol/L) and NOx decreases (0.0-0.40 mmol/L) with low DO values (0.06-4.22 mg/L). However, this 390 391 may also indicate areas of recent rainfall infiltration resulting in a dilution of NOx concentrations (e.g. 392 for well A4).

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394 5. Discussion
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395 5.1 Variations in the drivers of groundwater degradation

The vulnerability of groundwater degradation from either salinization or sewage leakage in Cotonou is temporally and spatially variable. The Quarternary aquifer is exposed to large fluctuations in the water table in response to rainfall changes; all sites show seasonal water table fluctuations higher than 0.5 m. However, both the timing of the water table fluctuations and the seasonal changes in groundwater quality varies between urban areas. 401 In Agla, a strong driver of groundwater quality is the proximity to lowlands. Agla is scattered 402 by these low elevation zones, which have either temporary or permanent ponds where groundwater seasonally discharges. As highlighted by the major ion results, wet season increases in the water table 403 404 can either result in increased groundwater salinity (EC up to $1,468 \mu$ S/cm) in these lowland discharge 405 areas due to mixing with pond water, or can result in lower groundwater salinity (EC < 842μ S/cm) due to the dilution effect from infiltrating rainfall. During the dry season, the groundwater stable 406 isotope values also highlight increased mixing with pond water and saltier Lake Nokoué water. 407 Elevated NOx concentrations in groundwater at Agla is due to sewerage contamination and was 408 recordered during both dry and wet seasons (NOx concentrations up to 0.86 mmol/L). Groundwater 409 contamination from sewerage may infiltrate directly from leaky latrines or from mixing with the 410 permanent ponds, which have accumulated NOx concentrations up to 17.50 mmol/L. According to 411 412 Starr et al. (1993), denitrification tends to occur in aquifers with very shallow groundwater compared to aquifers with deeper groundwater (more than 2 m). Low dissolved oxygen levels combined with 413 low NOx levels in shallow wells (even less than 1.0 m) in the study area (Fig. 8) where higher levels 414 415 can be expected, may be related to biodegradation or denitrification processes (Postma et al. 1991, 416 Jorgensen et al. 2004, Hassane et al. 2016 and Kadjangaba et al. 2018). Either by reducing NO_3 to 417 HCO_3 (1) or N_2 (2). According to Postma et al. (1991) and Anornu et al. (2017), the absence or low 418 concentrations of NO₂ and NH₄ in groundwater is probably related to the reduction of NO₃ to N₂. The 419 overall denitrification process can be described as (Berner, 1980):

420
$$5CH_2O + 4NO_3^{-} \longrightarrow 2N_2 + 4HCO_3^{-} + CO_2 + 3H_2O$$
 (1)

421 $2NO_3^- + 12H^+ + 10e^- \longrightarrow N_2 + 6H_2O$ (2)

However, denitrification may not be the only process that contributes to the increase of HCO₃
in shallow groundwater, which probably also results from the decomposition of organic matter and
mineral dissolution reactions (e.g. Alassane et al. 2015).

In Ladji, the main influence on groundwater salinization (EC up to 5,340 μ S/cm) is the mixing with the Lake Nokoué waters. This was obvious during the dry season where both stable isotopes values and major ion concentrations clearly highlight mixing, and the water level also shows an early dry season rise in values due to lake infiltration. In comparison, the sewerage contamination 429 of groundwater is mostly evidenced during the wet season (groundwater NOx up to 2.5 mmol/L).
430 During the wet season, increases in NOx from sewerage sources may result from both (i) rises in the
431 water table mobilising NOx in the unsaturated zone, and/or (ii) infiltrating rainfall and temporary
432 ponds recharging the shallow groundwater (as suggested by stable isotope results) and therefore
433 transporting NOx to the saturated zone.

434 St Jean is the only studied neighbourhood that does not have any surface water bodies and no groundwater discharge sites. Therefore, the wet and dry season variations in groundwater degradation 435 are potentially due to variations in the composition of infiltrating water, mixing between subsurface 436 437 waters, and reactions during water table fluctuations. Although St Jean has low groundwater salinity 438 levels (EC up to 1,285 uS/cm), this area has the highest NOx concentrations recorded in this study 439 (NOx up to 5.06 mmol/L). Since the local rainfall is the only origin of the shallow groundwater (as 440 seen from stable isotope values), it is expected that either recharging rainfall or rising water tables 441 transfers NOx from the sewerage sources to the groundwater system during the wet season. However, at St Jean, groundwater NOx concentrations are greater during the dry season (up to 5.06 mmol/L) 442 443 compared with the wet season (up to 2.58 mmol/L). So, instead, the contamination may be constant leakage from latrines throughout the year and wet season rainfall may act to dilute this contamination. 444 Almost all the pit latrines and septic tanks in Cotonou have depths between 1.50 and 2.50 m (Hounkpè 445 et al. 2014; Yadouléton, 2015). In St Jean, the maximum depth to water table is observed at 1.53 m in 446 447 the dry season, which means that dry season saturated zones remain close to leaking sewerage sources.

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449 **5.2** Periods of increased risk due to groundwater degradation

Generally speaking, groundwater samples with nitrate levels that exceed the Beninese quality standards for drinking water originated from wells J1, J3, J6, J7, J8, J9 and J11 (piezometer) during the dry season (February and March) in St Jean, from wells L7 and L11 during the rainy season (June and October) in Ladji, and from wells A2, A6 in the dry season (February and November) in Agla. Thus, risk for nitrate pollution in shallow groundwater shows high seasonal variation between sites. Similar results were obtained by Boukari (1998) and Totin et al. (2013) at different sites in Cotonou with higher NO₃ levels (up to 1,61 mmol/L) observed during recharge of the wet season. The measured EC values show a large amplitude of spatial and temporal variation. The EC values in Agla and St Jean are in the same order of variations (200-1,800 μ S/cm) whereas the groundwater in Ladji records EC values in the order of 750-5,340 μ S/cm. In St Jean, the waters are more saline during the dry season. Unlike St Jean, the wells in Agla and Ladji are saline during both the dry and the rainy seasons due to mixing processes with lake water.

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464 5.3 Shallow groundwater, an unregulated water resource in expanding urban environments

In urban areas of sub-saharan Africa, groundwater from shallow wells is commonly used to 465 466 partially or fully supply drinking water resources (Okotto at al. 2015). In major cities, such official groundwater resources are monitored and treated. For example, in the coastal city of Douala 467 (Cameroon), shallow groundwater is the main source for domestic and drinking purposes (Takem et 468 al. 2015). Likewise, in Bamako (Mali), around 55 % of the population uses water from aquifer 469 resources (British Geological Survey, 2002). Where shallow groundwater is not the official resource, 470 the shallow groundwater usually free and therefore commonly used in the impoverished areas of urban 471 472 sprawls, including for domestic uses. Normally this resource is meant for washing only, but it 473 commonly ends up for drinking, dish washing, shower and cooking water supplies. For example, the 474 interviews conducted during this study showed that 10 on 10 of the households in St Jean, 9 on 10 in Agla, and 4 on 10 in Ladji reported to use well water only for dish washing. In St Jean, 1 on 10 475 476 households declared to use well water for drinking water supply compared with Agla and Ladji where no households have reported such use. The results of the interviews have therefore showed that there 477 478 are differences in groundwater use between St Jean, Ladji and Agla. The underprivileged areas of the 479 city are subject to greater groundwater quality issues, and this also correspond to the parts of the city where the residents are more dependent on the groundwater as a domestic water resource. This is the 480 case in Kinshasa (Democratic Republic of Congo) where in peri-urban and rural inhabitants widely 481 482 use the unregulated and untreated shallow groundwater resources for drinking water supply (Ndembo, 2009). 483

484 Unfortunately, as seen in this case study of Cotonou, the shallow groundwater in urban 485 environments is often contaminated. This was also observed in Blantyre (Malawi) where drinking 486 water from shallow groundwater was heavily polluted by a lack of sanitation facilities and 487 indiscriminate waste disposal (Mkandawire, 2008). Likewise, high NO₃ concentrations in urban 488 shallow groundwater resources have been observed in Dakar, Senegal (NO₃ > 100 mg/L; Ndeye et al. 489 2017); Douala, Cameroon (NO₃ up to 241 mg/l; Ketchemen-Tandia et al. 2017) and in different 490 regions of Ghana (NO₃ up to 507 mg/L; Rossiter et al. 2010). 491 In many urban cases, the nitrate contamination of shallow groundwatear is from 492 anthropogenic sources (Martínez-Santos et al. 2017). This was also observed in this study and previous 493 work in Cotonou (Boukari, 1998; Maliki, 1993; Boukari et al., 1996; Totin et al. 2013), and as also 494 noted in other major sub-saharan cities such as N'djaména, Tchad, (Kadjangaba et al. 2018) and 495 Djibouti (Ahmed et al. 2017). As highlighted in the study of Ouagadougou, Burkina Faso (Yaméogo, 496 2008), such nitrate contamination is linked to the high population density that relies on archaic or nonexistent sanitation systems. This is a particularly significant challenge in informal settlements, like in 497 the cities of Douala (Takem et al. 2010; Ketchemen-Tandia et al. 2017) and Kampala, Ouganda 498 (Nyenje et al. 2013) where the increase in nitrates and chlorides in shallow groundwater are related to 499 faeces from pit latrines, sewages, landfills, surface discharges and droppings from domestic animals. 500

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- 502

503 6. Conclusion

Although pollution sources are identical for each of the three neighbourhoods studied, the 504 resultant transfers and reactions controlling concentrations are distinct. In the neighbourhood where 505 there is no surface inundation and acts as a local recharge area, the groundwater salinity values remain 506 507 low (EC \leq 1285 μ S/cm), however NOx concentrations are the highest recorded in this study (up to 5.1 mmol/L). In the neighbourhood bordering a lake where there is seasonal inundation and groundwater 508 509 discharge, the dry season interaction with lake water results in groundwater with highest observed EC values (up to 5340 μ S/cm). Stable isotope (δ^2 H and δ^{18} O) mixing ratios indicate mixing of up to 74 % 510 511 with lake water; In the neighbourhood located in a swamp that is subject to inundation during both the 512 small and large wet seasons also shows mixing with lake water (up to 62 % using $\delta^2 H$ and $\delta^{18}O$ values) resulting in groundwater EC values up to 1468 µS/cm. This discharge site notably has lower 513 514 NOx concentrations (up to 0.86 mmol/L), however this is not indicative of less sewerage 515 contamination only greater degradation processes. Understanding the seasonal changes in processes controlling groundwater quality between each site is key to identifying risks to the residents who use 516 this unregulated shallow groundwater resource for domestic purposes, including drinking water 517 supplies. Seasonal variations highlight heightened risks from sewerage and septic tank leakage during 518

the wet season in neighbourhoods located in discharge areas compared with increased risk during the dry season in the recharge area. In addition, there are increased risks of shallow groundwater salinization during both the small wet and dry seasons in discharge areas.

522 Stable water isotopes showed a direct relationship between local rainfall water and 523 groundwater at St Jean compared with Ladji and Agla where groundwaters are also be subjected to 524 mixing with enriched surface waters such as the lake water (via the lowlands) and the pond water, 525 particularly during the dry and small wet seasons.

526 Groundwaters chemistry in each of the neighbourhoods are different. In St Jean, the 527 composition of major ion concentrations of waters is dominated by Na-Ca-HCO₃-Cl groundwater 528 type, while Ladji is of the Na-HCO₃-Cl type and Agla of the Na-Ca-HCO₃-SO₄-HCO₃ one. The ponds 529 sampled in both Ladji and Agla showed some similarities in major ion compositions with local 530 groundwater, thus indicating a potential connection. The Lake Nokoué has greater concentrations of 531 Na and Cl compared with all groundwater samples.

532 Groundwater samples indicated that 26 % for NO₃ and 7.0 % for NO₂ do not comply with the 533 Beninese quality standard for drinking water. Based on Cl/Br molar ratios, sources of NOx in 534 groundwater appear to be dominated by infiltration of sewerage and septic tanks in dry season in St 535 Jean, while in Agla and Ladji, contamination was obvious in wet season following infiltrating rainfall 536 and ponds recharging. Low NOx for some of the groundwater samples may indicate effects from 537 biodegradation in discharge area or dilution from rainfall.

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Figure 1. (a) Location of Cotonou and sampling sites in the three neighbourhoods; (b) Sedimentary unit cross section (A-B) of marginal-littoral area of Benin (modified from Boukari, 1998); and (c) Environmental situation at Agla (A9 site)



Figure 2: Average monthly interannual rainfall at Cotonou station from 1971 to 2016 (ASECNA station)



Figure 3. Relationship between water-table variation and electrical conductivity at each site



Figure 4. Piper diagram for groundwaters, ponds and lake at each site



Figure 5. Temporal variation between, depth water table, electrical conductivity, major ion and isotopic stable at each site.



Figure 6. Relationship between ¹⁸O and ²H with Global Meteoric Water Line (GMWL) and Local Meteoric Water Line (GMWL).



Figure 7. (a) Groundwater, pond and lake water Cl/Br molar ratios with changes in Cl concentrations highlights different sources of Cl in the waters. (b) Relationship between Cl/Br ratios and Cl concentrations showing higher concentrations in NOx (circle sizes are relative to NOx concentrations).



Figure 8. Relationship between HCO3, NOx and DO for all groundwater sampled at Agla



Figure 9. Schematic of conceptual model for contaminant sources and transfers at (a) Agla, and (b) Ladji and Saint-Jean.

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Table 1. Chemical	narameters and	isotopic y	values in	shallow	groundwater
rable r. Chennear	parameters and	isotopic v	anues m	shanow	Stoundwater

ID	Sampli ng month	Origin	Dept h	T℃	pН	DO	EC	NOx	HCO3	CI	NO ₂ ⁻	Br	NO ₃	PO4	SO4	Na⁺	NH_4^+	K⁺	Mg ⁺⁺	Ca**	DOC	d ¹⁸ O	d²H
Units			m			ma/l	uS/cm	meq/							ma/l								200
J1	Nov-16	well		29.50	6.97	ilig/i	590	0.67	221.49	27.59	DBL	DBL	37.68	8.69	3.90	47.20	DBL	19.57	5.42	58.13		-3.07	-14.04
	Mar-17	well	0.80	29.70 28.20	7.24	2.59	998 1002	1.84	244.00 318.12	74.90 77.18	0.80 DBI	0.97 DBI	66.12 DBI	8.26 6.99	85.49 115.78	86.41 94.37	0.25 DBI	29.43	9.19 9.99	93.52 89.42	10.85	-2.85 -2 94	-12.48
	Oct-17	well	0.90	27.90	7.57	3.93	799	0.44	314.76	28.52	7.13	DBL	17.67	5.68	4.87	44.68	DBL	22.27	9.69	92.45	5.89	-3.89	-21.22
	Feb-18	well	1.23	29.10	7.16	2.35	1285	1.18	384.30 152.50	85.78 57.74	DBL	DBL	73.45	8.34	8.62 54.63	14.51	13.65	35.65	8.76	79.28	6.80	-3.36	-16.10
J2	Nov-16	well	0.00	29.00	6.92	0.24	502	0.76	99.56	34.52	0.12	0.28	46.95	6.93	36.23	45.47	0.28	12.67	2.80	44.77		-2.81	-11.64
	Mar-17 Jun-17	well	0.59	29.90 28.60	6.93 7.04	2.42	552 347	0.70	113.46 128.10	42.75 25.67	0.17 DBL	0.65 DBL	37.70 DBL	7.13 5.27	31.75 17.90	41.98 31.66	0.30 DBL	18.60 7.32	3.24 1.94	49.58 33.24	8.30	-3.35 -3.65	-15.40
	Oct-17	well	0.77	28.60	7.36	5.13	297		121.60	15.94	DBL	DBL	DBL	6.52	13.00	22.97	DBL	7.66	1.56	28.22	6.31	-4.17	-21.87
	Feb-18 Jun-18	well	1.04 0.75	29.70 29.03	6.88 7.06	2.56 3.39	586 341	0.50	119.56 86.93	52.64 2.19	0.43 DBL	DBL	3.49 12.43	6.40 3.77	49.37 26.87	37.97 22.87	0.62	13.28 8.45	3.14 1.79	46.95 31.72	5.55	-3.39 -3.54	-15.90
J3	Nov-16	well		29.70	7.02		658	0.74	125.58	46.97	0.53	0.61	45.39	9.82	54.88	62.97	1.64	17.99	3.46	53.96		-3.01	-13.16
	Jun-17	well		29.20 27.60	7.15	4.61 7.85	891	0.63	191.54 254.68	81.73 75.52	0.31 DBL	0.13	98.13	6.28 5.25	53.17 52.61	86.30 84.25	2.25 DBL	21.71 22.82	4.96	68.47 63.45	1.17	-3.02 -3.19	-12.99
	Oct-17	well		29.00	5.91	3.55	630	1 10	29.84	56.65	DBL	DBL	DBL 72.08	5.19	4.84	51.28	DBL	13.35	4.54	6.62	4.53	-3.56	-16.94
	Jun-18	well		29.40	6.90	3.90	792	1.12	158.60	7.88	DBL	DBL	69.65	3.83	41.29	66.47	0.38	14.90	5.79	63.11	5.57	-3.38	-15.19
J4	Nov-16	well		29.80	7.11		500	0.22	18.45	27.63	0.77	0.45	13.48	2.47	36.46	37.84	0.31	17.42	3.77	51.35		-2.37	-8.98
	Mar-17 Jun-17	well well		29.90 27.90	7.15 7.63	4.95 3.98	572 368	0.31 0.33	195.20 132.37	33.68 2.34	0.11 0.15	0.79 DBL	18.69 DBL	2.36 1.99	37.86 23.22	41.44 29.65	0.12 DBL	13.45 15.53	6.22 2.62	57.91 31.66	14.96	-2.58 -3.23	-9.93 -14.08
	Oct-17 Feb-18	well well		28.70 29.20	8.06 7.15	5.35 3.33	342 434	0.22	81.74 96.38	29.18 47.54	DBL DBL	DBL DBL	DBL 13.85	2.45 2.27	24.79 27.26	25.31 28.87	DBL 0.81	7.95 3.52	2.97 3.64	29.79 36.94	4.24 3.79	-3.07	-13.61
	Jun-18	well		28.80	7.41	4.08	379	0.33	122.00	25.24	DBL	DBL	18.82	2.37	28.23	24.75	0.48	16.77	3.00	36.36		-3.52	-15.23
J5	Nov-16	well		29.70	6.83		642	0.45	28.68	38.17	DBL	0.82	0.28	0.82	3.38	46.38	DBL	43.42	6.31	56.38		-3.22	-14.19
	Mar-17	well	1.05	29.70	6.89	3.25	715	0.69	17.80	57.97	0.43	0.96	42.66	0.21	59.64	72.32	0.34	37.83	3.69	49.65	11 57	-3.06	-13.09
	Oct-17	well	0.43	27.10	7.03	1.31	630	0.40	272.60	37.12	DBL	0.19	DBL	DBL	26.37	43.66	DBL	19.78	5.82	65.35	4.93	-3.61	-17.56
	Feb-18	well	1.53	28.60	6.96	2.34	558	0.46	164.70	38.37	DBL	DBL	2.89	DBL	44.12	34.22	0.50	16.46	3.98	46.12	8.59	-2.89	-12.01
.16	Jun-18 Nov-16	well	0.66	27.80 29.10	6.97 6.76	2.03	639 548	0.25	265.35 66.49	35.12 41 99	DBL	DBL 0.49	1.53 71.59	DBL 5.27	31.40 51.16	44.85 47.46	0.62	19.72 21.98	5.49 3.87	62.37 4.27		-3.46 -3.24	-15.23
00	Mar-17	well		29.40	6.99	3.60	1326	5.67	85.40	111.13	1.95	0.18	311.33	2.68	91.61	115.00	0.73	37.66	13.4 6	17.85		-2.85	-11.68
	Jun-17	well		29.20	7.06	5.12	1129	3.80	154.79	99.84	54.7 9	DBL	157.26	2.57	87.74	99.79	0.39	4.98	1.37	87.46	16.65	-2.60	-10.19
	Oct-17	well		28.50	6.99	4.40	637	1.94	54.90	44.28	2.28	DBL	117.19	2.33	53.18	46.00	DBL	21.69	4.32	48.44	8.31	-3.28	-15.11
	Feb-18	well		29.00	6.97	4.35	961	3.37	439.20	51.14	0.17	DBL	28.87	2.62	47.94	93.79	0.16	29.53	3	114.74	8.63	-3.32	-15.88
J7	Jun-18 Nov-16	well		29.00 30.10	7.17 6.87	5.96	346 431	0.43 0.74	54.90 93.25	29.35 26.42	DBL 0.19	DBL 0.40	24.95 45.76	4.54 3.19	2.49 27.42	26.22 34.79	0.27	11.87 13.95	2.27 2.89	24.73 36.26		-3.23 -2.85	-13.82
	Mar-17	well		29.90	7.05	5.12	695	0.87	184.22	44.50	0.22	0.11	53.78	0.79	44.46	53.47	0.18	2.82	4.77	55.50	0.00	-3.08	-12.98
	Jun-17	well		30.20	7.00	3.78	742	0.56	22.36	61.44	0.26	DBL	DBL	1.14	68.21	63.46	DBL	22.43	5.45	63.31	6.80	-2.61	-9.98
	Oct-17 Feb-18	well		28.90 29.00	7.14	4.20 4.19	652 647	1.29	16.14 113.46	38.98	5 0 4 9	DBL	57.27 95.59	0.73	59.96 59.66	4.94 38.17	DBL 0.23	15.58 24.95	5.28 4.69	62.60 48.66	4.59 7.08	-3.13 -3.04	-14.66
	1 00 10			20.00						0.00							0.20	21.00		10.00	7.00	0.01	
J8	Jun-18 Nov-16	well		29.40 29.40	7.08 6.99	4.19	616 701	0.66	14.30 126.41	37.00 4.12	0.58	0.66	4.89 136.47	1.87 4.61	56.59 32.89	47.58 38.52	0.24	31.95 18.17	4.20 5.39	43.35 78.67		-3.38 -3.35	-14.29
	Mar-17	well		29.70	6.78	3.62	885	2.58	115.90	72.18	0.48	0.16	159.26	5.87	36.16	74.78	2.78	29.87	4.94	61.28		-2.77	-11.16
	Jun-17	well		29.20	7.27	4.67	894	2.56	128.86	79.23	93.7	DBL	32.54	4.43	54.90	8.34	0.71	25.57	5.65	72.48	5.82	-2.92	-11.84
	Oct-17	well		27 80	7 4 9	4 75	564	0.99	128 10	22 92	45.5 7	DBI	DBI	5.67	3 47	36 54	DBI	19 17	3 18	53 58	5 54	-3.64	-17 99
	Feb-18	well		28.50	6.91	3.30	840	2.60	118.34	56.71	0.46	DBL	16.63	6.20	6.15	59.92	0.32	25.37	4.63	59.75	4.09	-3.12	-13.91
J9	Jun-18 Nov-16	well		28.40 29.50	7.65 7.19	4.56	377 448	0.62	128.10 213.58	14.14 25.43	0.23	0.47	38.59 0.18	4.18 DBL	18.97 16.48	19.89 18.58	0.46 4.60	1.39	2.17 3.11	38.72 49.79		-3.55 -2.96	-13.45
	Mar-17	well		30.10	6.98	4.82	461	0.88	82.96	42.48	0.69	0.19	54.42	DBL	21.57	36.85	0.33	11.14	2.94	41.27		-2.64	-9.55
	Jun-17 Oct-17	well well		29.10 28.10	7.20 7.21	4.06 4.04	437 323	0.49 0.46	126.58 97.60	24.64 27.93	8.35 1.56	DBL DBL	19.20 0.42	DBL DBL	26.35 11.57	29.73 17.19	0.26 0.58	1.53 5.12	3.20 2.22	46.72 36.84	8.89 0.41	-3.06 -3.44	-12.08 -15.80
	Feb-18	well		28.60	7.00	4.63	242	0.13	56.12	28.38	DBL	DBL	8.23	DBL	11.53	14.34	0.60	7.64	1.82	26.23	4.58	-3.74	-18.62
.110	Jun-18 Nov-16	well		28.60 29.20	7.04	3.67	251 572	0.15	122.00 13.63	14.44 41.52	DBL 0.32	DBL 0.49	9.45 43.32	2 64	15.59 38.84	11.91 56.63	0.43	5.34 26.72	1.62	38.59 42.15		-3.43 -3.44	-13.35
0.0	Mar-17	well		28.50	6.89	2.37	615	0.39	173.24	42.31	0.86	0.96	23.27	DBL	48.34	42.65	1.74	37.39	4.64	52.68		-3.17	-14.39
	Jun-17 Oct-17	well		29.50 29.30	7.13 8.03	3.43 4.94	770 481	1.12	163.18 98.82	65.67 39.22	DBL DBL	DBL	69.32 DBL	DBL 2.51	61.25 47.65	72.53 28.13	1.46 DBL	33.84 13.14	4.70 3.82	57.83 47.88	11.67 3.81	-3.17 -3.17	-13.29
	Feb-18	well		29.60	6.90	2.77	943	0.99	229.36	8.94	0.56	DBL	6.44	0.52	47.17	6.59	14.36	28.62	5.34	56.68	7.19	-3.14	-13.91
	Jun-18	well		30.50	6.75	0.73	878	0.90	25.10	65.43	DBL	DBL	55.54	1.52	47.69	66.77	5.41	27.62	6.57	65.57		-3.28	-14.91
J11	Feb-18	Piezo meter	1.40	31.30	7.21		497	0.89	92.72	57.13	DBL	0.60	55.30	4.48	27.92	42.63	2.29	11.97	4.67	46.88	13.50	-3.09	-15.03
	lup 10	Piezo	0.99	20.40	7 99	6 70	444	0.65	141 50	19.61	DBI	DBI	4.96	6 47	22.82	12.44	0.15	10.46	4 01	40.02		0 70	17.66
Δ2	5011-10	meter	0.00	23.40	7.55	0.75	444	0.05	141.52	13.01	DDL	DDL	4.20	0.47	22.02	13.44	0.15	12.40	4.01	43.32		-5.75	-17.00
712	Nov-16 Mar-17	well well		28.60 29.90	6.60 6.28	3.32	1164 990	0.72 0.87	432.95 13.70	118.77 123.64	DBL 0.39	0.38 0.22	0.44 53.39	DBL DBL	33.32 136.56	15.85 125.35	0.25 0.23	35.77 28.66	1.85 7.45 17.4	8.73 49.37		-1.81 -2.72	-5.39 -10.94
	Jun-17	well	-0.01	27.30	7.21	2.42	1273		435.85	16.53	DBL	0.35	DBL	DBL	74.85	147.93	DBL	35.17	6	78.28	22.27	-2.32	-9.17
	Oct-17	well	0.04	26.50	7.02	2.33	1150		441.64	14.00	DBL	0.53	DBL	DBL	1.45	111.34	DBL	29.87	7	81.80	25.62	-2.02	-7.51
	Feb-18	well	0.69	28.30	6.96	2.83	733	0.83	22.52	97.21	DBL	0.40	0.51	DBL	32.23	86.88	0.32	15.55	14.7 5	41.54	14.65	-2.90	-12.08
	Jun-18	well	0.04	28.90	7.02	1.61	959		292.80	111.26	DBL	DBL	DBL	DBL	3.56	93.69	DBL	24.14	1.24	68.46		-2.75	-11.56
A4	Nov-16 Mar-17	well well		29.60 30.90	6.02 5.85	2.95	326 269	0.27 0.39	31.95 27.00	17.78 27.70	0.35 0.57	0.12 0.17	1.63 2.32	DBL DBL	78.71 44.29	29.22 29.25	0.16 0.38	6.91 6.53	3.73 3.67	16.65 13.12		-3.22 -3.12	-14.01 -14.05
	Jun-17	well	0.47	30.70	6,01	5,34	353	0.43	37.82	56.98	DBL	0.23	2.52	DBL	38.15	32.98	0.54	8.24	5,30	2,99	9,86	-2.96	-12.84
	Oct-17	well	0.29	28.00	5.95	4.53	497	0.18	54.90	69.62	DBL	0.35	1.95	DBL	61.86	46.93	DBL	1.46	6.77	27.46	8.79	-2.68	-10.44

A5	Feb-18 Jun-18 Nov-16 Oct-17	well well well well	0.99 0.25	29.50 29.00 30.70 27.20	5.84 5.76 6.00 7.01	3.97 3.32 2.75	548 666 1037 500	0.42 0.66 0.27	36.60 36.60 13.62 15.60	86.61 97.38 158.50 5.92	DBL DBL DBL DBL	0.46 DBL 0.25 0.13	25.86 37.26 16.97 DBL	DBL DBL DBL DBL	5.46 59.26 135.17 19.18	51.32 6.43 131.15 51.31	0.76 0.10 0.90 DBL	1.49 12.96 17.93 9.32	7.16 1.98 6.68 6.26	2.14 32.71 37.45 33.27	10.38 10.06	-2.89 -2.81 -2.77 -3.52	-12.28 -11.54 -11.89 -17.61
A6	Feb-18 Jun-18 Nov-16 Jun-17	well well well well		29.00 29.20 30.30 29.00	6.55 6.74 6.16 7.12	2.39 4.33 8.45	652 652 983 1338	0.39 0.17 0.86 0.90	154.94 155.55 155.55 245.53	6.87 48.84 1.56 175.75	0.20 DBL 1.67 8.61	0.12 DBL 0.17 0.17	23.85 1.53 51.84 38.63	DBL DBL 1.38 DBL	58.67 52.19 125.27 216.32	67.19 52.72 113.75 158.45	2.58 0.12 DBL DBL	13.73 11.19 22.17 38.76	9.77 8.15 5.98 1.17	47.38 4.83 45.76 96.61	8.85 14.60	-3.18 -3.46 -2.99 -2.02	-14.14 -16.07 -12.98 -8.14
A7	Oct-17 Feb-18 Jun-18 Nov-16	well well well well		27.60 30.30 29.60 34.00	6.50 6.22 6.72 7.51	5.40 4.25 8.31	1095 967 935 1302	0.37 0.77 0.59 0.36	162.26 87.84 144.88 125.58	128.13 137.65 91.65 44.84	0 DBL DBL DBL	0.26 0.33 DBL 0.12	0.14 47.68 31.57 0.23	DBL DBL DBL DBL	172.62 165.24 19.87 43.42	13.59 128.49 13.93 58.32	DBL 2.42 0.24 1.63	28.97 26.57 25.35 13.37	6.95 7.92 6.24 3.69	62.43 47.19 51.88 21.27	7.43 7.08	-2.49 -2.62 -2.66 -3.57	-10.75 -10.75 -11.51 -18.22
	Mar-17 Jun-17 Oct-17	well well well		30.20 30.70 29.10	6.21 6.31 6.28	2.01 2.52 2.12	631 763 687	0.75 0.42 0.35	97.60 95.75 112.24	84.84 11.86 13.59	DBL 0.52 DBL	0.26 0.28 0.29	0.46 0.19 0.19	DBL DBL DBL	72.31 131.16 74.26	76.93 93.31 84.18	0.91 0.68 0.36	21.42 19.39 16.73	8.34 1.26 6.42	25.66 34.31 28.45	6.80 6.19	-2.96 -2.92 -2.89	-13.00 -12.16 -13.15
A8	Feb-18 Jun-18 Nov-16	well well well		29.30 28.30 29.50	6.20 6.38 6.20	2.12 2.83	689 816 551	0.56 0.63 0.25	79.30 155.55 75.41	12.42 18.42 7.68	DBL DBL DBL	0.38 DBL 0.13	0.35 0.39 0.16	DBL DBL DBL	74.59 67.64 59.96	8.83 89.20 66.77	1.77 0.89 0.12	21.11 27.33 6.86	8.15 9.20 1.90	29.82 32.96 26.92	7.79	-2.99 -3.02 -2.58	-13.11 -13.62 -10.57
	Mar-17 Jun-17 Oct-17	well well well	0.02 -0.12	29.10 29.70 27.80	6.03 6.37 6.03	2.13 3.98 4.51	560 842 645	0.28 0.27	78.80 98.67 64.23	73.69 11.17 70.00	0.19 0.19 DBL	0.24 0.24 0.26	1.26 1.39 DBL	DBL DBL DBL	69.90 171.18 149.72	84.16 16.87 92.14	1.28 2.49 0.22	7.44 13.32 3.37	2.28 6.34 4.96	22.49 44.49 2.22	9.95 7.31	-2.90 -2.87 -2.67	-12.29 -12.41 -10.51
A9	Feb-18 Jun-18 Nov-16 Mar-17	well well well well	0.48 -0.20	28.90 29.10 29.00 29.70	6.29 6.15 7.04 6.85	2.93 7.49 2.34	642 778 666 765	0.18	73.20 143.35 246.67 174.46	77.37 78.94 67.77 82.83	DBL DBL DBL 0.53	0.30 DBL 0.23 0.26	1.14 DBL DBL 0.45	DBL DBL 1.46 0.39	116.63 139.61 8.92 93.83	82.51 95.79 67.94 91.12	0.88 0.91 DBL 0.34	5.91 7.69 18.82 1.21	4.88 9.16 9.32 9.13	2.54 33.59 34.30 52.77	8.89	-2.69 -2.64 -0.33 -2.52	-10.74 -10.29 0.51 -11.02
	Oct-17	well		28.90	7.78	8.22	931		374.83	77.28	DBL	0.29	DBL	4.26	2.17	7.46	2.98	71.94	15.2 4	49.57	20.79	-2.29	-10.29
	Feb-18	well		28.40	7.84	4.22	1475	0.37	475.80	155.00	DBL	0.66	0.23	5.15	93.18	127.81	5.78	55.96	19.5 7	87.88	26.91	-0.03	2.78
	Jun-18	well		27.40	7.29	0.56	1200		488.00	94.55	DBL	DBL	DBL	3.19	66.50	86.96	1.48	53.93	2.64	92.00		-2.27	-9.04
A10	Nov-16 Mar-17	well well		28.50 31.10	6.56 7.55	2.48	700 1023	0.96 0.96	35.84 381.86	34.77 96.39	0.62 0.14	0.74 0.24	5.14 0.50	3.69 DBL	25.72 5.69	39.59 92.29	9.44 2.52	16.33 31.14	5.00 1.53	64.39 8.56		-3.29 -1.11	-14.98 -1.57
	Jun-17 Oct-17	well well		30.20 29.20	7.60 7.84	6.36 6.43	837 835	0.17	193.68 245.22	84.73 119.79	DBL DBL	DBL 0.20	DBL 0.66	DBL DBL	136.69 4.98	85.33 11.90	DBL DBL	22.51 26.97	7.26 8.75	75.40 44.20	14.76 22.53	-1.85 -0.29	-6.87 -1.64
	Feb-18	well		28.70	7.77	0.12	1870	0.12	751.52	112.19	DBL	0.76	0.77	DBL	37.48	121.80	88.40	44.42	9.35	45.65	26.47	-0.26	1.72
	Jun-18	well		28.80	7.33	3.71	654		262.30	6.83	DBL	DBL	DBL	DBL	4.68	67.89	DBL	14.47	6.75	47.42		-2.03	-6.96
A11	Feb-18	Piezo meter Biozo	0.40	32.20	6.61	1.19	1212	0.14	346.48	12.41	0.56	0.69	0.94	DBL	84.86	85.96	7.45	43.27	9.63	8.93	23.4	-2.86	-13.02
	Jun-18	meter	-0.26	28.20	7.23	0.06	1468	0.42	491.66	148.36	DBL	DBL	24.94	DBL	54.75	114.51	5.11	64.13	6	91.52		-2.46	-10.21
L2	Mar-17	well		29.30	6.80	1.69	2090	0.17	247.66	385.88	1.26	0.14	0.62	DBL	17.50	317.38	2.52	53.61	4 29.4	54.18	10.00	-2.21	-8.43
	CCL-17	well	0.41	20.20	6.02	3.10	2180	1.07	323.30	428.37	1.03	0.20	2.07	DBL	204.40	315.83	2.56	46.02	2 26.6	42.72	15.09	-2.25	-8.75
	Jun-18	well	0.64	28.00	6.92	2.52	2030	0.63	366.61	431.66	1.35	0.15	8 59	DBL	294.16	327.83	0.23	62 13	5 31.2	64.83	13.10	-2.38	-8.85
L6	Nov-16	well	0.01	28.50	6.87	2.02	1608	0.00	41.99	255.80	0.65	0.41	2.57	DBL	78.11	213.45	2.75	43.28	0 23.1	71.64		-1.98	-6.20
	Mar-17	well		29.50	7.12	2.37	1660		512.40	31.17	0.89	0.35	0.19	2.17	1.91	244.35	1.97	45.50	23.7	76.22		-0.33	1.60
	Jun-17	well	0.19	28.70	7.73	2.29	1543		47.50	261.68	0.85	0.96	0.59	DBL	43.27	22.77	4.32	38.96	21.1 4	76.80	24.63	-2.30	-9.74
	Oct-17	well	0.09	26.30	7.69	1.61	1352		386.74	219.13	0.63	DBL	DBL	2.54	7.87	172.98	DBL	38.28	17.5 2	56.54	15.12	-1.92	-8.77
	Feb-18	well		28.50	7.84	1.63	1669		529.48	232.88	0.91	0.18	1.59	1.84	4.19	192.75	5.76	48.46	17.6 6	64.46	18.10	0.35	4.27
	Jun-18	well	0.36	28.60	7.39	0.56	750		266.57	85.77	0.24	DBL	DBL	1.83	1.46	8.56	0.15	26.70	9.42	39.00		-1.34	-4.37
L7	Nov-16	well		31.90	7.13		1623	0.13	323.22	21.94	0.55	1.66	12.69	4.45	121.78	197.25	0.65	48.74	24.8 6	81.14		-2.66	-10.42
	Mar-17	well		30.70	7.13	1.70	2350	0.17	585.60	414.29	1.29	0.23	14.25	2.38	134.38	365.33	1.26	39.34	42.1 5	81.31		-1.99	-6.96
	Jun-17	well	0.38	29.40	7.69	7.32	1357	0.39	279.69	18.85	0.56	1.44	89.49	4.15	135.36	177.48	0.19	44.73	18.8 3	58.85	7.03	-2.32	-8.25
	Oct-17	well	0.31	27.00	7.39	4.32	1378	5.46	263.52	16.51	0.41	0.90	48.44	3.49	119.67	152.12	DBL	47.78	19.4 0	66.91	5.93	-2.70	-12.96
	Feb-18	well	1.25	29.10	7.17	1.98	2600		622.20	446.98	1.68	DBL	DBL	0.83	151.54	374.26	2.51	31.75	5 15.6	66.38	13.96	-2.11	-7.32
	Jun-18	well	0.54	29.70	7.68	4.87	1210		248.27	13.48	DBL	1.21	75.72	4.45	117.46	14.17	0.33	44.64	0	48.68		-3.36	-14.50
L11	Feb-18	well		30.00	7.00	3.60	1945	0.15	35.00	253.78	0.62	0.94	57.98	2.16	273.55	236.26	2.69	59.67	8	78.17	11.85	-2.38	-9.83
1.10	Jun-18	well piezom	0.00	29.70	6.89	3.32	2004	2.51	266.57	292.14 1561.3	DBL	1.72	13.25	4.33	248.81	229.91	0.76	/.81	8 63.4	97.83	0.15	-2.41	-9.59
L18	Mar-18	eter	0.39	33.40	6.84	0.10	5340		4.16	5	5.28	0.87	0.54	DBL	136.77	832.90	7.33	13.91	8	13.00	8.45	-1.86	-6.72

Table 2: Chemical	parameters a	nd isotopia	values in	surface water
rable 2. Chemiear	parameters a	ina isotopi	varues m	surface water

	Samplin					= 0															180	.2
ID	g month	Origin	I℃	рН	DO	EC	NO _x	HCO ₃	Cl	NO ₂	Br	NO ₃	PO ₄	SO_4	Na⁺	NH₄⁺	K*	Mg**	Ca++	DOC	d™O	d⁺H
Units					mg/L	μ3/c m	meq/L							mg/L							%	50
A2	Nov-16	pond	32.40	7.02		818		287.54	82.19	BDL	0.26	BDL	1.98	19.41	76.54	0.26	14.21	7.84	62.59		-1.51	-3.18
	Jun-17	pond	25.20	8.72	7.77	113		41.02	5.19	BDL	BDL	BDL	0.60	4.27	6.27	0.41	3.47	0.74	1.79	10.91	-2.13	-12.69
	Oct-17	pond	26.20	7.25	4.65	970	0.00	358.68	108.79	BDL	0.57	BDL	4.49	16.65	96.12	BDL	22.87	1.32	84.99	25.43	-2.65	-12.20
A3	Jun-18 Nov-16	pond	28.80	0.90 7.51	1.51	1242	0.02	380.03 459 10	95.54 149.59	0.74	0.43	0.99	2.20	24.29	79.57 128.17	1.22	52.62 58.34	13.90	59.34 59.61		-2.89	-12.18
7.0	Jun-17	pond	29.10	8.24	9.53	1619	0.02	408.70	249.90	BDL	0.39	BDL	1.53	154.25	26.27	BDL	67.54	16.88	89.66	36.99	-2.26	-8.64
	Oct-17	pond	29.80	8.43	11.61	346		137.86	24.35	BDL	0.18	BDL	3.41	0.69	29.56	BDL	14.22	4.19	26.47	18.45	-2.41	-14.50
46	Jun-18	pond	30.10	7.05	1.07	998		256.20	62.27	BDL	BDL	BDL	0.57	47.74	64.51	0.18	22.14	7.55	57.43	0 10	-3.74	-17.80
A0 A7	Nov-16	pond	28.40	6.20	7.40	504		498.68	4.50	BDL	0.37	BDL	4.53	17.86	174.65	BDL	8.33	9.77	37.36	0.12	3.88	24.74
	Jun-17	pond	28.00	8.04	6.66	250		103.85	13.47	0.37	BDL	BDL	BDL	15.79	17.78	BDL	6.42	1.99	27.58	7.86	-2.74	-15.01
	Oct-17	pond	35.70	9.05	15.70	1070		470.31	92.01	BDL	0.24	BDL	4.50	42.50	117.64	BDL	83.62	17.53	41.35	39.75	-0.01	-6.12
48	Jun-18 Nov-16	pond	30.70	8.46	4.33	534		1/3.85	38.68	BDL	BDL	BDL 0.12	5.77	18.98	49.25	BDL 6.42	35.93	7.92	21.97		-4.27	-23.07
70	Jun-17	pond	35.90	9.64	12.38	1278	0.05	97.60	116.75	0.67	0.16	1.89	BDL	395.69	113.71	0.40	39.75	9.80	131.91	26.26	-0.33	-1.29
	Oct-17	pond	29.70	6.38	2.34	659		147.62	62.96	BDL	0.45	BDL	BDL	8.25	73.47	BDL	13.73	6.78	38.46	15.41	-2.63	-11.50
	Jun-18	pond	26.90	6.80	0.08	695	0.08	253.15	43.38	3.46	BDL	BDL	1.85	52.53	56.39	0.39	24.74	8.47	47.66		-3.45	-16.37
A9	Nov-16	pond	31.10	6.88	0.01	764		305.53	74.35	BDL	0.24	BDL	2.13	4.88	71.43	BDL	19.62	12.90	44.65	05.00	0.35	5.38
	Oct-17	pond	29.40	8.41	2.40	1643		581.33	240.73	BDL	0.17	BDL	5.25	11.54	75.29 166.41	3.32 BDL	145.23	26.22	38.85	25.28	-2.09	-6.20
	Jun-18	pond	31.20	8.58	10.72	711		260.47	69.58	BDL	BDL	BDL	0.40	16.46	7.35	BDL	35.68	12.37	34.24		-3.29	-16.12
A10	Nov-16	pond	30.10	6.69 7.48	2.07	1310	0.02	544.58 314.15	127.78	BDL	0.44	1.51	0.50	2.21	12.24 92.28	0.18 BDI	4.32	14.42	73.28	43.28	-0.93	-1.43
	Oct-17	pond	28.70	6.96	1.82	621		250.71	69.50	BDL	0.19	BDL	0.85	6.44	62.84	BDL	21.67	7.69	41.23	27.14	-3.30	-16.92
Δ12	Jun-18 Feb-18	pond	28.70	7.17	3.29	758 768		271.45 42.70	75.47	BDL	BDL 0.99	0.25 BDI	BDL	15.37 254.62	72.84	0.37	23.60	8.97	54.34 27 38	18.06	-2.90	-11.70
A12	Jun-18	pond	28.10	6.69	0.17	605		176.90	53.70	BDL	BDL	BDL	1.47	19.75	63.75	0.25	15.75	6.33	35.76	10.00	-3.98	-19.48
A13	Feb-18	pond	31.90	4.12	6.77	2900	17.55	0.00	256.87	0.59	BDL	188.	12.8	24.59	219.67	28.2	95.68	68.63	221.89	55.14	-2.44	-9.10
	Jun-18	pond	29.30	6.72	0.03	537		247.05	43.53	BDL	BDL	BDL	BDL	6.82	46.36	BDL	12.94	9.75	39.46		-3.27	-14.28
A14	Feb-18	pond	29.60	7.28	0.01	1671		456.28	189.73	BDL	0.43	BDL	5.98	291.84	146.16	3.67	92.25	25.54	139.40	31.90	-1.51	-0.95
	Jun-18	pond	27.90	6.90	0.30	685	0.19	271.45	55.38	0.32	BDL	11.1 5	4.33	7.70	54.90	0.42	5.35	9.86	42.82		-3.08	-13.31
A15	Jun-18	pond	27.80	6.93	0.12	829	0.14	298.90	65.99	BDL	BDL	8.95	4.87	18.13	69.75	0.25	3.56	1.98	66.66		-3.35	-14.37
A16 A17	Jun-18	pond	28.40	7.32	0.86	712		282.43	57.09	BDL	0.14	BDL	2.75	6.71 12.48	55.33 83.16	BDL	37.75	1.63	53.39 67 19		-2.87	-11.27
A18	Jun-18	pond	31.90	7.44	7.33	563		250.10	40.80	BDL	0.92	BDL	1.45	16.12	45.58	BDL	26.43	6.77	42.95		-3.60	-16.65
A19	Jun-18 Oct-17	pond	27.30 24.80	7.15	0.19	796 732		284.26	74.60 84.00	BDL	0.38	BDL	1.83	14.76 37 13	78.96 83.44	0.28	17.67	6.92 5.85	62.75 28.74	15.26	-2.05	-7.31
1.2	Nov 16	loko	29.00	7.00	2.10	19630	0.01	79.49	6166 20	BDL	10.45	0.02	7.50	011.04	3168.6	0.02 PDI	144.00	375.64	146 70	07.10	2 20	-11.13
L3	NOV-10	lake	32.90	7.18	C 15	18630	0.01	73.43	50.04	BDL	0.10	0.82	BUL	011.04	4	BDL 0.77	144.99	3/5.64	146.79	27.19	-3.29	-14.44
	Jun-17 Jun-17	pond	29.10	8.05	7.19	480		401.08	946.12	BDL	2.21	BDL	3.45 BDL	34.35 339.79	53.23 655.14	BDL	158.33	4.37 48.63	23.57 85.54	27.60	-1.85	-4.57
	Oct-17	pond	28.80	7.43	2.60	1371		235.46	314.06	BDL	1.34	BDL	3.41	31.84	180.00	BDL	27.68	22.49	35.97	22.60	-1.93	-9.03
L6	Jun-17	pond	29.90	8.17	0.02	2050		555.56	255.63	BDL	0.55	BDL	14.9	98.67	275.76	27.9 7	137.28	12.36	28.24	28.67	-0.98	0.71
1.8	Nov-16	lake	29.60	7.01	/	23200		93.10	8273.96	BDL	25.72	0.34	BDL	134.38	4495.2	BDL	158.90	54.36	171.41		-0.71	-0.85
														1867.2	4 8478.3						••••	
	Mar-17	lake	31.00	7.65	5.10	47100	0.01	109.80	15655.71	0.33	49.97	BDL	BDL	8	9	8.64	251.58	153.00	268.12		1.13	7.60
	Jun-17	lake	28.90	7.55	3.75	38800		103.70	14687.98	BDL	5.48	BDL	BDL	2232.5	848.96	1.94	286.80	966.60	298.53	2.30	1.09	8.46
	Oct-17	lake	25.70	7.29	2.00	886		102.48	229.88	BDL	0.94	BDL	BDL	19.31	123.87	BDL	11.32	14.36	16.98	13.01	-2.20	-10.87
	Feb-18	lake	28.70	7.86	3.32	47300		109.80	18119.73	BDL	62.32	BDL	BDL	2473.3	9449.1	0.71	36.92	928.45	34.41	0.98	0.97	7.46
	1 10		00 70		0.54	00400		100 71	10111.00		00.40	001	001	2293.2	, 8415.5	0.47	00.50	454.00	070.00		4.00	07.00
1.0	Jun-18	аке	29.70	7.41	2.54	39100		128.71	16114.88	BDL	62.43	BDL	BDL	9	2	0.47	36.52	151.93	276.92	45.40	-4.99	-27.32
L9 L10	Jun-17	pond	31.10	9.85	0.02	2000		156.62 491.97	254.45	BDL	0.23	BDL	2.25	89.27	195.47 295.17	0.10	64.58 96.45	2.42	11.95 64.45	45.49 27.02	-0.57	-16.01
	Oct-17	pond	28.70	9.61	3.28	2440		500.20	509.98	BDL	0.82	BDL	8.83	158.87	384.79	2.74	12.76	12.65	4.88	39.97	-1.13	-5.51
L14	Feb-18	lake	29.60	7.63	4.40	47600		109.80	19310.02	BDL	74.44	BDL	BDL	2667.7 1	12219. 87	21.2 4	381.49	1256.9 1	419.23	0.71	0.75	6.26
	Jun-18	lake	30.10	7.60	3.82	39400		131.15	16379.80	BDL	57.53	BDL	BDL	236.46	8487.5	0.51	323.79	158.49	263.38		-3.36	-8.85
115	Feb-18	lake	30.20	7 90	7 29	49700		113.46	18031.65	BDI	69 73	BDI	BDI	249 49	9579.6	25.3	32.24	955.30	323 52	1.02	0.93	7 34
210	10010	laito	00.20	1.00	7.20	107.00			10001.00	001	00.70	002	002	2165.4	3 7884 5	3	02.21	000.00	020.02	1.02	0.00	7.01
	Jun-18	lake	37.00	7.67	5.85	38100		57.95	15173.94	BDL	52.19	BDL	BDL	7	8	0.34	296.81	994.89	254.81		-1.34	-14.50
L16	Mar-18	pond	29.20	8.65	1.16	3830		1115.0 8	824.56	BDL	3.41	BDL	7.25	2.78	698.15	BDL	34.17	73.68	11.74	40.37	0.56	5.05
	Jun-18	pond	29.10	7.37	0.01	1024		289.75	120.41	BDL	0.17	BDL	6.98	41.34	121.28	0.83	39.94	12.58	39.45		-1.34	-4.37
L17	Mar-18	pond	28.90	7.20	0.01	2000		274.50	356.31	BDL	5.95	BDL	BDL	224.64	314.95	1.96	12.20	19.63	43.30	33.59	-2.02	-6.86
1.10	Jun-18	pond	30.90	6.86	0.35	831		259.25	46.16	BDL	BDL	BDL	1.66	1.43	47.63	2.57	1.98	8.63	44.32		0.50	5.33
LIS	Jun-10	ponu Occara	30.30	0.00	0.04	533		201.32	53./3	BUL	BUL 7.00	BUL	3.00	8.52 2317.0	100.00	0.24	20.00	5.28 1219.0	32./1		0.50	0.15
AUT	Jan-19	Ocean				51000		132.00	19091.00	BUL	7.00	BUL	BUL	0	1170.00	0.70	384.00	0	391.00		0.52	4.00
AO2	Jan-19	Ocean				51300		135.00	19172.00	BDL	64.00	BDL	BDL	2329.U 0	0	0.40	382.00	1216.U 0	386.00		0.54	4.69

