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1 **RESEARCH ARTICLE**

2

3 **Title**

4 Distribution of soil properties along forest-grassland interfaces: influence of permanent environmental
5 factors or land-use after-effects?

6

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19

20 **Abstract**

21 Soil properties vary spatially according to land use; both because land users have selected specific soil
22 properties for specific land uses, and land uses modify the soil properties. However, permanent
23 environment factors and land-use effects are unlikely to display the exact same spatial patterns. Study
24 of the spatial and historical patterns of distribution of soil properties could help to separate between
25 these two causes. In this aim, we studied 22 forest-grassland interfaces with controlled historical
26 configurations in northeast France. In each land use (forest and grassland), three distances to the edge
27 (edge, periphery and core) and two land-use histories (ancient and recent) were studied.

28 Along forest-grassland interfaces, forests were usually located slightly upslope of grasslands, and
29 mainly because this non-random topographic position the topsoil texture was significantly more silty
30 in forests, and clayey in grasslands. After statistically controlling for the effects of topography and soil
31 texture, we observed two main gradients of variation in soil properties according to the distance-to-
32 edge (acidity in forest and nutrient content in grassland). In forest, pH and Ca dropped from the edges
33 to the peripheries (15 m distance), while in grassland, C, N, P and Na sharply increased from the edges
34 to the cores (25 m distance). These results demonstrate, through the edge effect, the strong influence
35 of the land use on a part of soil properties. Furthermore, less than two centuries after grassland
36 afforestation or deforestation, we observed that soil properties in recent forests and recent grasslands
37 were respectively closer to their current land use than to their former land use. These results
38 demonstrate a rapid change in soil properties after land-use change. However, recent forests and recent
39 grasslands kept a legacy of soil texture from their former land use, respectively. Recent grasslands also
40 kept a lower soil density, N and Na content compared to ancient grasslands.

41 Hence, this study of forest-grassland interfaces show strong and short-scale relationships between land
42 use and soil properties and suggest that they express both original choices of land users for specific
43 soil properties and land-use after-effects. The non-random topographic position of the forest-grassland
44 interfaces indicates a conscious choice of this positioning by the land users, for agronomic reasons.
45 Beyond that, land use, through vegetation composition and management practices, also has a strong
46 impact on soil properties. The fact that land-use changes affect most soil properties after only a few
47 decades confirms the existence of land-use effects over time.

48

49 **Keywords**

50 soil; forest; grassland; land use; edge effect; land-use change

51

52 **1. Introduction**

53 Five soil-forming factors, *i.e.* climate, geology, topography, land use and time have been identified as
54 the source of the heterogeneity of soils in current landscapes (Jenny, 1941). Macro-climate, geology
55 and time mainly affect soil and humus properties at intermediate and large scales (McKenzie and
56 Ryan, 1999; Ponge *et al.*, 2011). At the local scale of a hillslope or a small catchment, geology,
57 topography and land use (including micro-climate) are generally the dominant factors affecting soil
58 properties (McBratney *et al.*, 2003). As man has increasingly changed his environment during the
59 Holocene, the effect of current and past land uses is a central question of ecology.

60 In most temperate regions, landscapes are currently marked by mosaics of open (*e.g.* grassland,
61 cropland) and closed (*e.g.* forest) land uses. Although relatively small compared to forest-cropland
62 contrast, differences in soil properties have been reported between forest and grassland: forests
63 generally have lower pH, thicker humus, higher C/N and lower P than grasslands (Joimel *et al.*, 2016).
64 However, the distribution of soil properties along forest-grassland interfaces remains poorly known.

65 Differences in vegetation and management practices, *i.e.* land-occupation and land-use influence, may
66 be intrinsically responsible for most of these differences in soil properties (Jenny, 1941). Differential
67 erosion, leaching, and bioturbation may modify the transport and redistribution of soil particles and
68 elements (Wang *et al.*, 2001). Grassland harvesting or grazing may decrease soil fertility, whereas
69 mineral fertilization and manuring may increase it, by increasing nutrient (N, P, K) contents relative to
70 forests (Chantigny, 2003). By spillover, fertilizer applications can also induce edge effects on soil
71 properties within the forest edges adjacent to the grasslands (Piessens *et al.*, 2006). In addition,
72 between forest and grassland, differences in tree, shrub and herb cover, as well as in taxonomic and
73 functional composition may be the source of different plant-soil feedbacks (Ehrenfeld *et al.*, 2005;
74 Harrison and Bardgett, 2010). Consequently, the species turn-over observed in plant communities
75 along forest-grassland interfaces (Burst *et al.*, 2017) can both induce and reflect an edge effect on soil
76 properties.

77 To properly study land-use effects on soil properties, all other soil-forming factors (climate, geology,
78 topography and time) should be equal (Motzkin *et al.*, 1996), but in most studies this condition is

79 rarely met. Time effects or slight topographical variations are often neglected. Yet, following the
80 catena concept (Hook and Burke, 2000), topography finely controls for erosion, mobilization and
81 redistribution of water, particles and nutrients along hillslopes (Seibert *et al.*, 2007). Indeed, in valley
82 bottoms, soils are better supplied with water and nutrient-richer than on slopes and uplands (Brubaker
83 *et al.*, 1993; Hook and Burke, 2000). Consequently, strong relationships between topography, soil
84 properties and vegetation composition due to these processes have been observed (Binkley and Fisher,
85 2013; Solon *et al.*, 2007). However, in the absence of topographic homogeneity, it is thus difficult to
86 attribute the variation in soil properties to a present land-use effect (Schimel *et al.*, 1985). Soil texture
87 is a typical soil characteristic which is both strongly linked to topographical position and modified by
88 land use and which affects most of soil physical and chemical properties (Hassink, 1997; Laganière *et*
89 *al.*, 2010). As a result, the non-random spatial distribution of soil properties, and more specifically
90 their agronomic and forest values, may have played a crucial role in the spatial distribution of past and
91 current land use.

92 Furthermore, historical land-use changes (*e.g.* afforestation and deforestation) may blur the influence
93 of current land use by reducing differences in soil properties between forests and grasslands. Indeed,
94 after afforestation of former grasslands, topsoil properties of recent forests should gradually tend
95 towards those of the ancient forests (Li *et al.*, 2012) and display intermediate values between those of
96 ancient grasslands and ancient forests. Similarly, after deforestation, soil properties of recent
97 grasslands should display intermediate values between those of ancient forests and ancient grasslands.
98 Thus, depending on the resiliency of the former soil properties after land-use change, a temporary
99 mismatch between current land use and soil properties can be observed (Peña *et al.*, 2016). With the
100 acceleration of land-use changes in the last decades (Alstad *et al.*, 2016), it is becoming urgent to
101 better understand the effects of these changes on soil properties.

102 To decipher the influence of permanent environment factors and of land-use effects on soil properties,
103 we focused our study on forest-grassland interfaces and sought to answer the following questions:

104 1. How are soil properties distributed along forest-grassland interfaces? To what extent land use and/or
105 distance-to-edge impact soil properties?

106 2. To what extent afforestation and deforestation modify soil properties? Do soil properties of recent
107 land uses (forest or grassland) remain closer to those of their former land use or rapidly became
108 similar to those of their new land use?

109

110 **2. Materials and Methods**

111 2.1. Study area and site selection

112 The study was conducted in two sectors of the Lorraine regional natural park (PNRL) in northeast
113 France (figure 1, 215-277 m a.s.l.). The climate is semi-continental with annual rainfall and
114 temperature averaging 775 mm and 10.4°C, respectively, for the period 1981-2010. The geological
115 substrate is mainly a clayey marl from the Oxford Clay formation in the western sector and the Keuper
116 formation in the eastern sector. All soils studied are vertic cambisols (IUSS Working Group WRB,
117 2015), derived from post-Holocene aerial deposits of silts of various thickness overlaying the clayey
118 marl material (Antoine *et al.*, 2003).

119 Vectorized land-use maps (scale 1:40,000) over the PNRL allowed us to preselect forest-grassland
120 interfaces. Land-use changes were characterized by comparing the land use on French Etat-Major map
121 (established from 1826 to 1831 in the area, see Favre *et al.* 2013 for the vectorization) and the current
122 land-use map (vectorized since 1998 by the PNRL and updated in 2014, with a typology of land-uses
123 more detailed but included in that of the Etat-Major map) using a geographic information system
124 (ArcGIS®, v.10.2, ESRI, France). When the land use had not changed since 1826-1831, a date which
125 is close to the minimum forest area in northeast France, we categorized it as ancient, and when the
126 land use was different, we qualified it as recent. For each preselected forest-grassland interface, series
127 of aerial photographs (spaced by ten years maximum between 1931 and 2014,
128 <http://www.geoportail.gouv.fr>) were photo-interpreted to confirm the origin and the stability of the

129 forest, the grassland and the edge over time. As a result, we could confirm that the land-use changes,
130 and therefore recent habitats, were all over 50 and under 188 years old.

131 To avoid confounding factors other than present land-use effect, edge effect and land-use change
132 effect, the interfaces were then selected according to several environmental criteria, including:
133 homogeneity of topographic positions within and between interfaces (mainly wide and open valley
134 bottoms and footslopes), minimum area of both adjacent land uses (≥ 1 ha) to ensure to have a core
135 area, linearity of the edge to avoid multiple edge effect and absence of disturbances (*e.g.* paths,
136 streams, trenches). Forests were only mature deciduous forests, treated as high forests or former
137 coppices with standards now in conversion to high forests, mainly dominated by oaks, beech and
138 hornbeam (*Carpinion betuli* Issler 1931) without recent logging. Grasslands were mesophilous to
139 meso-hygrophilous mowing grasslands, dominated by tall grasses (*Arrhenatherion elatioris* Luquet
140 1926 and *Bromion racemosi* Tx. in Tx. et Preising ex de Foucault 2009) *i.e.* temporarily wet meadows
141 (Mucina *et al.* 2016), most of which are flooded each year. Unlike forests managed for the most part
142 by the National Forestry Office, grasslands were owned by private farmers who may have varying
143 management practices. In the last decades, most meadows were mowed once a year around June 15th,
144 and some were fertilized by manure application and / or chemical fertilizers (NPK granules). Traces of
145 burrowing activities (mainly due to earthworms and wild boars) were regularly observed in the
146 periphery of grasslands. After field surveys, 22 interfaces conforming exactly to the above criteria
147 were retained for our study (11 in the eastern sector and 11 in the western sector of the PNRL, figure
148 1).

149

150 2.2. Study design

151 We sampled three types of interfaces with different historical configurations (figure 1): (i) stable
152 interfaces (8 sites) corresponding to ancient forest adjacent to ancient grassland, (ii) afforestation
153 interfaces (7 sites) with recent forest (developed on a former grassland) adjacent to ancient grassland,
154 and (iii) deforestation interfaces (7 sites) with ancient forest adjacent to recent grassland (formerly

155 forested). Therefore, we sampled four land-use histories combining two histories (ancient, recent) and
156 two land uses (forest, grassland). Each interface included six plots, three in forest and three in
157 grassland, positioned at three increasing distances from the forest-grassland boundary and referred
158 according to plot position as edge, periphery and core (appendix A). Each plot had an area of 2 m x 50
159 m (100 m²) and was oriented parallel to the forest-grassland boundary. As the distance of influence of
160 the edge effect is recognized as lower in grassland compared to forest (Schmidt *et al.*, 2017), periphery
161 and core plots were placed closer to the boundary in grassland (5 m and 25 m, respectively) than in
162 forest (15 m and 50 m). A total of 132 plots in the 22 forest-grassland interfaces were sampled.

163

164 2.3. Soil sampling and soil analyses

165 In each plot (appendix A), three equally spaced topsoil samples (excluding litter) were collected in
166 September 2014 to a depth of 4.5 cm with a steel cylinder of 250 cm³ volume (8.4 cm in diameter, and
167 4.5 cm in height). The choice of the sampling at depths of 0 to 4.5 cm makes it possible to focus only
168 on the first soil horizon, which is the one most influenced by the land uses and which interacts most
169 with the composition of the vegetation and the management practices. Samples were bulked,
170 homogenized and air-dried. Samples were then sieved to 2 mm to remove roots, pebbles (< 1 %) and
171 other debris.

172 For each of the 132 topsoil samples, the following soil properties were measured: (i) the particle size
173 distribution in three fractions *i.e.* clay (0-2 μm), silt (2-50 μm) and sand (50-2000 μm) using the
174 pipette method without organic matter destruction, (ii) the dry bulk density (hereinafter simply
175 mentioned soil density or density) based on the residual weight of an aliquot of 30 g of fine soil dried
176 in an oven (105°C, 48h), then extrapolated to the total 750 cm³ soil sample, (iii) the total organic
177 carbon (C) and the total nitrogen (N) using dry combustion, (iv) the available phosphorus (P), (v) the
178 cation exchange capacity (CEC) and the exchangeable cations (Ca, Mg, Na, K, Fe, Mn and Al) after
179 cobaltihexamine extraction, and (vi) the soil pH in water. By adding the calculated C/N ratio, a total of

180 17 soil properties were used in subsequent analyses (see appendix B for the list of soil properties and
181 references for the methods used).

182

183 2.4. Humus and soil description

184 In each plot, in addition to soil sampling to analyses of soil properties, we also described the humus,
185 dug a soil pit down to 30 cm and probed the soil with an auger down to 1 m to describe the different
186 horizons by visual estimation and to the touch for texture, color, structure, compaction, root density,
187 effervescence and redoximorphic features. The humus types were mainly mulls (from dysmull to
188 eumull), except two plots characterized as eumoder type (Jabiol *et al.*, 2009). For all the plots, soil
189 texture was clayey-silty to clayey with less than 10 % of sand on average. As soils derived from the
190 overlay of silty deposits over a clayey material, they all showed a more or less abrupt textural change
191 in the soil sampling, at various depth (Brêthes, 1976). Upslope, the depth to which this transition
192 occurred was 30 cm on the average (10 to 60 cm), while downslope, this depth was about 10 cm (0 to
193 60 cm). Close to the valley bottoms, pseudogley features were observed from an average depth of 10
194 cm (0 to 50 cm).

195

196 2.5. Topographic data collection

197 Within each site, two complementary topographical observations were collected for each plot: (i) the
198 topographic position was noted as an ordered factor (4 levels: bottom = 1, footslope = 2, mid-slope = 3
199 and upper slope = 4), and (ii) the slope percentage from the forest-grassland boundary measured with a
200 Suunto® clinometer (accuracy < 0.5 % of slope). Slopes along the 75 m long interfaces were low, 2.3
201 % on average and 5.7 % at the maximum (see also appendix C). It is interesting to note that in all the
202 sampled cases, the ancient forests were located upslope from the ancient grasslands.

203

204 2.6. Data analysis

205 To identify the main soil gradients along the forest-grassland interfaces and take into account
206 collinearities between soil variables, a principal component analysis (PCA) was performed on all the
207 17 soil properties. Due to the major influence of topography and soil texture on most physical and
208 chemical soil properties, a partial PCA was then performed in order to partial out the effects of these
209 two variables. The topography *i.e.* the topographic position was included as an ordered factor with
210 four levels (bottom, footslope, mid-slope and upper slope) and the soil texture as a continuous
211 variable, the soil clay content.

212 To examine how land use, distance to edge and land-use change influenced soil properties along the
213 forest-grassland interfaces, three linear mixed models (LMMs) were fitted:

214 LMM1: $Y \sim \text{Position} + \text{History}:\text{Position} + \text{Topography} + (\sim 1|\text{Sector}/\text{Site})$

215 LMM2: $Y \sim \text{Position} + \text{History}:\text{Position} + \text{Topography} + \text{Clay content} (\sim 1|\text{Sector}/\text{Site})$

216 LMM3: $Y \sim \text{Position} + \text{History}:\text{Position} + (\sim 1|\text{Sector}/\text{Site})$

217 The first model (LMM1) applied to the three textural properties (clay, silt and sand contents) included
218 as fixed effects: the position along the interfaces (6 levels: forest core, forest periphery, forest edge,
219 grassland edge, grassland periphery and grassland core), the interaction of the history (2 levels: ancient
220 and recent) with the position, and topography as a categorical ordered control variable. This first
221 model allowed us to check whether the land use *per se* had an effect on soil texture, once removed the
222 influence of topography on the particle size distribution along the slopes. The second model (LMM2)
223 applied to all the other soil properties included as fixed effects: the position, the interaction of the
224 history with the position and the control variables topography and clay content. The third model
225 (LMM3) was applied to the first two axes of the partial PCA and included as fixed effects: the position
226 and the interaction of the history with the position. The last two models allowed us to evaluate the
227 influence of land use on the soil properties by controlling for the effects associated with the non-
228 random distribution of land uses along the slopes, which depends on topographic position and soil
229 texture. Topography and clay content were not included in LMM3, because the effects of these control
230 variables had already been removed by partialling them out of the PCA. In a second step, to check to

231 what extent the control for topography and soil texture within the first two models (LMM1 and
232 LMM2) affected our results, the third model was also applied to each soil property. For the three
233 models, the site (22 levels) nested in the sector (2 levels: east and west of the PNRL) was introduced
234 as a random effect.

235 For each each soil property and the first two partial PCA axes, the goodness-of-fit of LMMs was
236 assessed using two indicators: (i) the marginal r^2 (*i.e.* variance explained by fixed effects only) and (ii)
237 the conditional r^2 (*i.e.* variance explained by both fixed and random effects) (Nakagawa and
238 Schielzeth, 2013). The part of random (site) effects relative to the total of random and fixed effects
239 was calculated as: $1 - \text{marginal } r^2 / \text{conditional } r^2$, and expressed in percentage. To determine the
240 statistical significance of fixed effects in models, type-III analyses of variance were performed.

241 Contrast analyses were then performed to test for more specific assumptions using the interaction of
242 the history with the position. Firstly, to properly test for the present land-use effect (*i.e.* test for
243 differences between and due to current forests and grasslands) and notably avoid any confounding
244 effects linked to land-use change and distance-to-edge, only plots from cores of ancient land uses (15
245 plots of ancient forest cores *vs.* 15 plots of ancient grassland cores, AFC-AGC) were compared.
246 Secondly, to test for the edge effect in each land use, we compared the ancient forest cores *vs.* ancient
247 forest edges (AFC-AFE, both 15 plots) and the ancient grassland cores *vs.* the ancient grassland edges
248 (AGC-AGE, both 15 plots). Thirdly, to test for the land-use change effect in recent forests and in
249 recent grasslands, cores and peripheries of each land-use having the same history were grouped in
250 order to improve the power of the tests. The edges of the two land uses were excluded from these
251 groups because of their close physical proximity. We tested for the legacy of soil properties after land-
252 use change by contrasting the recent forests against the ancient forests (RFCP-AFCP, 14 *vs.* 30 plots)
253 and the recent grasslands against the ancient grasslands (RGCP-AGCP, 14 *vs.* 30 plots). We then
254 tested for the change of soil properties after land-use change by contrasting the recent forests against
255 the ancient grasslands (RFCP-AGCP, 14 *vs.* 30 plots) and the recent grasslands against the ancient
256 forests (RGCP-AFCP, 14 *vs.* 30 plots). All the p-values associated with contrast analyses were
257 adjusted by controlling for the false discovery rate (Benjamini and Hochberg, 1995).

258 Statistical analyses were carried out with R 3.2.3 (R Foundation for Statistical Computing, Vienna,
259 AT) and the following packages were used: *vegan* for PCA, *nlme* for LMMs, *piecewiseSEM* to
260 calculate the r^2 and assess the quality of models, *car* for analyses of variance and *lsmeans* for the least-
261 squares means estimations and the contrasts tests.

262

263 3. Results

264 3.1. Effects of topography and soil texture on soil properties

265 The topography showed a significant influence ($p\text{-value} \leq 0.05$) on clay, silt and sand (table 1), but
266 only on sodium among the other soil properties. Besides, the soil texture, *i.e.* the soil clay content,
267 showed a significant influence on most of the other soil properties (table 1), except on Al and Fe. The
268 goodness-of-fit of models was high for most soil properties (conditional $r^2 \geq 0.5$) (table 1). Only the
269 goodness-of-fit for Al and Fe contents was lower (conditional $r^2 = 0.39$ and 0.36 , respectively).

270

271 3.2. Main soil gradients along the forest-grassland interfaces

272 The first PCA axis (44.7 % of the total variation) showed a gradient from coarser textured soils (high
273 silt and sand contents) to soils of finer texture (high soil clay content) (appendix D). The second PCA
274 axis (15.4 % of the total variation) showed a gradient from acidic soils (with relatively high Fe and Al
275 contents and high C/N) to more neutrophilic soils (with a higher pH). After controlling for topography
276 and soil texture (figure 2a), the first partial PCA axis (28.2 % of the total variation) showed a gradient
277 from acidic soils to more neutrophilic soils, the same as the previous second PCA axis obtained
278 without topography and soil texture control. The distribution of types of humus along this axis
279 confirmed its interpretation as an acidity axis (figure 2a). The second partial PCA axis (19% of the
280 total variation) showed a K-Na opposition, the soils richer in Na being also higher in Mg, Mn and P,
281 and displaying a higher organic matter content (C and N concentrations). Correlations of the first two
282 partial PCA axes with soil properties are given in appendix E. Both the first two partial PCA axes

283 reflected a forest-grassland gradient (figure 2b). Along the forest-grassland interfaces, both the first
284 (soil acidity) and the second (gradient of nutrient variations) partial PCA axes were significantly
285 influenced by the plot position (appendix F), but only the second one was also significantly influenced
286 by the interaction of history with plot position.

287

288 3.3. Soil texture

289 From ancient forest cores to ancient grassland cores, the clay content increased while the silt content
290 decreased linearly (figure 3, appendix G and H3), with a pronounced gap between forest edges and
291 forest peripheries. Significant land-use and edge effects were found for these two texture properties
292 (table 1). In recent land-uses (both forest and grassland), clay and silt contents were intermediate
293 between those of ancient forest and ancient grassland (figure 3). Thereby, land-use change also had a
294 significant effect, but the recent land uses were closer to their former land uses than to their current
295 ones (table 1 & 2). For clay, silt and sand, the goodness-of-fit of models was very high (conditional r^2
296 > 0.8), but reflected a high or moderate site effect (table 2).

297

298 3.4. Soil acidity

299 Within interfaces, Fe and Al were higher in ancient forest compared to ancient grassland (appendix
300 G), with a remarkable gap between the two edges, although located a few meters apart. In accordance,
301 pH was higher in ancient grassland than in ancient forest and showed a decline from edges to cores in
302 forest (figure 3). These three soil properties showed a significant position effect (table 1), but none for
303 the interaction of history with plot position. However, the land-use and edge effects were not
304 significant (table 2). After afforestation, we observed no significant change or legacy in recent forest,
305 but after deforestation, Fe was significantly lower (appendix G) and pH higher (figure 3) in recent
306 grassland compared to ancient forest (table 2). A relatively strong site effect was found for these three
307 soil properties (table 2).

308

309 3.5. Soil nutrients, organic matter and density

310 From ancient forest cores to ancient grassland cores, density, C/N and K decreased, and conversely C,
311 N, P and Na increased (figure 3 and appendix G). Most of these soil properties showed a significant
312 position effect, except for density and C (table 1). Thereby, C/N, N, P, Na, K, but also C showed a
313 significant land-use effect (table 2). For almost all of these soil properties, a significant edge effect
314 was observed in grassland, but not in forest (table 2), except for K for which the opposite was found.
315 Indeed, for K, in addition to a decreasing gradient (not significant) from ancient grassland edges to
316 ancient grassland cores, a significant decreasing gradient was observed from ancient forest edges to
317 ancient forest cores (table 2 and figure 3).

318 Interacting with plot position, land-use history showed a significant influence on the density, C/N, C,
319 N and Na (table 1). After afforestation, density, C/N, but also K were higher in recent forest cores and
320 peripheries compared to ancient grassland cores and peripheries, while C, N, Na and P were lower
321 (figure 3 and appendix G). Thereby, significant changes for all these soil properties (including K and
322 P) were observed in recent forest compared to ancient grassland (RFGP-AGCP comparisons, table 2),
323 while no significant differences (legacies of previous land use) were found between recent forest and
324 ancient forest (RFGP-AFGP comparisons, table 2). After deforestation, C/N and K were lower in
325 recent grassland compared to ancient forest, while density, N, Na and P were higher (figure 3).
326 However, only the changes for C/N, P and K were significant (RGGP-AFGP comparisons, table 2).
327 Besides, significant differences for density, N and Na were observed between recent grassland and
328 ancient grassland (RGGP-AGCP comparisons, table 2), with a lower N and Na and a higher density in
329 recent grassland (figure 3). It is remarkable to note that none of these soil properties showed both a
330 significant change and a significant legacy, either after afforestation or after deforestation.

331 For all these soil properties (*i.e.* density, C/N, C, N, Na and K), except P, a low site effect was found
332 (table 2).

333

334 3.6. Effectiveness of topography and soil texture control and validity of the results

335 The use of the third model LMM3 on each of the soil properties, *i.e.* without topography and soil
336 texture control, confirmed the results obtained with the two previous models LMM1 and LLM2
337 (appendix H). Indeed, apart from two rare exceptions, one for Fe and one for K, the control of
338 topography and soil texture in the first two models simply led to a general decrease in the significance
339 of most of the results (appendix H). We can also note that in the absence of topography and texture
340 control, the goodness-of-fit of the third model was weaker for most soil properties except Sand, Fe and
341 Al (appendix H).

342

343 4. Discussion

344 4.1. Land-use and edge effects

345 4.1.1. Soil texture

346 In the smooth landscapes of the PNRL plains, texture variations are first related to the occurrence of
347 silt deposits, overlying the clays of the Keuper and Oxfordian marlstone (Antoine *et al.*, 2003). We
348 showed that these textural variations are spatially organised along the slopes, ranging from coarser
349 texture upslope to finer texture downslope (Brubaker *et al.*, 1993; Hook and Burke, 2000) even at very
350 slight slopes (average slope: 2.3 %), and that they overlap with the distribution of forests (upslope) and
351 grasslands (downslope). This non-random correlation between soil texture gradients, slopes and
352 relative position of forests and grasslands in the landscape could be partly the result of past land-use
353 choices based on soil texture and topography. The main reason is that grasslands were always located
354 in the most favorable places in terms of soil moisture to promote fodder growth. However, the abrupt
355 textural change from the forest peripheries to the edges (at 15 meters apart) also suggests a possible
356 effect of land use. Several explanations for this land-use after-effect may be proposed: (1) the rapid
357 erosion mainly due to rainfall of the silty layer of grassland soils, but also adjacent forest edge soils, in
358 the years following logging and before the establishment of perennial herbaceous cover, (2) the

359 burrowing activity of earthworms, moles and boars in grasslands raises clay from deeper clayey
360 horizons to silty surface horizons and decrease the silt percentage, and (3) the possible illuviation of
361 clay particles in forest soils, although this is generally more common in sandy soils. The erosion of the
362 soil silt cover after forest clearing has indeed been demonstrated by various authors (since the bronze
363 or the Roman ages) in northeast France (Etienne, 2011), Belgium (Vanwalleghem *et al.*, 2006) and
364 south Germany (Houben, 2008), but still poorly reported within footslopes and valley bottoms at
365 forest-grassland interfaces. Besides, the impacts of the burrowing activity on soil surface texture in
366 grasslands are well known, while the slightly acid deciduous forest soils are known to favor the
367 illuviation of clay particles into a clay-enriched B horizon (Binkley and Fisher, 2013).

368

369 4.1.2. Soil acidity and soil nutrients content

370 Controlling for the effects of topography and soil texture, two main soil gradients related to soil
371 acidity and soil nutrients content were observed along the forest-grassland interfaces. Forest soils,
372 particularly in ancient forests, were slightly more acidic than grassland soils, possibly because of the
373 litter input on the topsoil (Glatzel, 1991) and cation leaching (von Oheimb *et al.*, 2008). However, we
374 found no land-use effect or edge effect on this soil acidity gradient. The small differences in pH and
375 humus (mainly mull type) that we found along the forest-grassland interfaces can probably explain
376 these results.

377 Besides, C, N, P and the C/N ratio showed a land-use effect associated with a strong edge effect in
378 grassland. The highest C and N observed in grassland topsoils (especially in grassland cores) are likely
379 due to a greater root biomass decomposition compared to forest soils (Hook and Burke, 2000). Indeed,
380 the annual turnover of soil organic matter derived from tree and shrubs in forests could be lower than
381 that from annual and perennial herbs in grasslands (Guo *et al.*, 2007; Guo and Gifford, 2002). The
382 unilateral edge effect for soil nutrients content, found only in grassland, probably also reflects the
383 influence of management practices such as grassland fertilization. This seems particularly true for P
384 (Brubaker *et al.*, 1993), which also exhibited a strong site effect which could be due to varying

385 fertilization levels among grasslands. It should be noted that other influences could also contribute to
386 these results, such as differences in net productivity between forest and grassland, bioturbation by ants
387 and worms higher in grassland, or even greater abundance of leguminous plants in grassland.

388 A land-use effect associated with a strong edge effect was also observed for Na with increasing values
389 from grassland edges to grassland cores. Unlike the previous ones, Na is the only element having
390 shown a topography effect. It likely highlights the well-known local presence of salt and gypsum
391 lenses within or near the clayey marls of our sites (BRGM 2001). In the lower lands, these salt lenses
392 can cause extensive salinization of near-surface soil horizons by dissolution within groundwater and
393 frequent rise of the water table. In the natural region of Saulnois, on the same geological substratum
394 and a few kilometers away from several of our study sites, these salt lenses are known to be at the
395 origin of resurgence of salt water and the presence of mainland salt meadows (Duvigneaud 1967). In
396 our sites, the extensive salinization observed in the grassland cores could thus be related to the
397 increasing soil moisture from forest to grassland, confirmed by plant bio-indication (Burst *et al.*,
398 2017), resulting from a lower topographic position and probably also from a lower evapotranspiration
399 of grasslands compared to forests (Nosetto *et al.*, 2007).

400 We also found a land use effect for K, but unlike the previous ones, it was mainly associated with a
401 strong edge effect in forest. This result could be due to a spillover of mineral fertilizers applied to
402 adjacent grasslands (Kepfer-Rojas *et al.*, 2015; Piessens *et al.*, 2006). Indeed, it was showed that many
403 elements can penetrate into forest edges during and after grassland fertilization, in particular by aerial
404 deposits (Thimonier *et al.*, 1992). However, we also observed a slight edge effect in grassland, and the
405 lowest K contents were found in grassland peripheries and cores. These results are possibly due to the
406 higher abundance of grasses within grassland peripheries and cores we observed (Burst *et al.* 2017),
407 which are better competitors and extractors of K compared to forbs (Bezemer *et al.*, 2006; Tilman *et*
408 *al.*, 1999). Because our grasslands are at least annually subject to the export of plant biomass by
409 mowing, K is probably here a limiting element, although this is sometimes not the case in low-
410 intensity grasslands (Swacha *et al.*, 2018). Another explanation is the increasing clay contents found in
411 soils towards grassland cores, which retain the exchangeable K. Note also that the highest K was

412 found in forest edge soils. This could be due to the fact that the trees and shrubs, whose highest
413 densities were found at the forest edges, can (i) affect the distribution of cations such as K through
414 uptake from deep soil horizons and transport to topsoil horizons by deposition of a cation-rich litter
415 (Jobbágy and Jackson, 2004), but also (ii) increase the K-rich throughfall inputs (Thimonier *et al.*,
416 1992).

417

418 4.2. Land-use change effects

419 4.2.1. Land-use legacy and gradual change of soil texture

420 After afforestation and deforestation, soil texture remained closer to the former land use than to the
421 new land use. Obviously, recent forests were established on grassland clayey soils while recent
422 grasslands resulted from the deforestation of silty forest soils. We assume that the afforestation could
423 not rapidly change the silt content of former clayey grassland soils, except if newly established forests
424 have preserve silt colluviation from upsloping ancient forests. Besides, in case of deforestation, the
425 fact that recent grasslands were enriched in clay compared to ancient forest cores may be related to a
426 rapid land-use effect after land-use change, here probably an erosional redistribution of silt and clay
427 particles after clearance of the forest.

428

429 4.2.2. Rapid change of soil properties after afforestation

430 In contrast to many studies showing the long-term persistence of soil properties after afforestation of
431 former cultivated parcels (Blondeel *et al.*, 2018; Dupouey *et al.*, 2002; Koerner *et al.*, 1997), no legacy
432 in soil properties was found from ancient grassland to recent forest after afforestation of grassland. As
433 observed in some previous studies (Graae *et al.*, 2003; von Oheimb *et al.*, 2008 for recent forest
434 developed on former grazed heathland), in recent forest, most soil properties became similar to those
435 found in ancient forest in less than two centuries. One explanation is that the differences in soil
436 properties would be lower between forest and grassland than forest and cultivated field (Janssen *et al.*,

437 2018; Joimel *et al.*, 2016). Another explanation would be that traditional management practices were
438 not as intensive as modern practices (Flinn and Marks, 2007; Kepfer-Rojas *et al.*, 2015). For example,
439 the fertilizer inputs in grassland, mainly from organic matter, were relatively low before the 20th
440 century (Green, 1990).

441 Besides, we showed higher density, C/N and K, as well as lower C, N, P and Na in recent forest
442 compared to ancient grassland. Because these values were mostly higher (density) or lower
443 (particularly P and Na) than those found in ancient forest, these results could likely reflect the
444 existence of a short period of cultivation before or after grassland use, followed by afforestation.
445 Indeed, cultivation generally increases soil density and greatly depletes former forest soils in organic
446 matter and nutrients (Murty *et al.*, 2002). We can also think of an increase of fertilizer inputs in
447 grassland in the modern period (supported by higher C/N and lower C, N and P in recent forest) and/or
448 a rapid decline in soil nutrients after afforestation. Indeed, after afforestation, soil nutrients are
449 progressively depleted through their sequestration within perennial plant tissues, in particular in the
450 tree and shrub trunks and roots (Laganière *et al.*, 2010). Regarding the decline of Na in recent forest, it
451 has also been shown that increased evapotranspiration due to trees and shrubs after afforestation could
452 result in a lowering of groundwater (Nosetto *et al.*, 2007) and consequently limit the intake of Na in
453 topsoils.

454

455 4.2.3. Land-use legacy or rapid change of soil properties after deforestation

456 After deforestation, density, C, N and Na have shown a legacy from forest past land use to grassland.
457 The higher soil density in recent grassland compared to ancient grassland could be the result of
458 forestry work during harvesting and a short period of tillage. Indeed, even if this higher density was
459 partly related to lower C, the density was also higher in recent grassland compared to ancient forest,
460 which is likely the result of soil compaction (Batey, 2009). Besides, the lower C and N in recent
461 grassland compared to ancient grassland can be explained by a relatively shorter period of
462 accumulation of nutrients and belowground organic matter and/or a better decomposition of organic

463 matter. Indeed, we showed that ancient grassland has the wettest soils (Burst *et al.*, 2017), which are
464 frequently and annually flooded. Consequently, the decomposition of organic matter is likely lower in
465 ancient grassland due to seasonal anaerobic periods (Hassink, 1997, 1994). Note that in their meta-
466 analysis, Guo and Gifford (2002) have reported an increase in organic matter and especially C in
467 recent grassland after deforestation. However, these authors found an increase in C in high-
468 precipitation zones, with 2000-3000 mm, but not in drier zones with less than 2000 mm. Finally, the
469 lower Na soil concentration in recent grassland compared to ancient grassland could be explained by a
470 deeper water table and consequently less frequent upwelling of dissolved salts in topsoils (Nosetto *et*
471 *al.*, 2007).

472 Unlike the previous soil properties, C/N, P and K have undergone a rapid change in recent grassland
473 compared to ancient forest. These results reflect the impact of management practices in recent
474 grassland. Indeed, the increase in P and the decline of C/N likely reflect the introduction of
475 fertilization in recent grassland. Besides, the decline in K demonstrates the impacts of both a higher
476 potassium content in the hay exported from grassland, in comparison with its very low level in wood,
477 and the higher rhythm of biomass exportation from grassland.

478

479 **Conclusions**

480 On a very short scale and despite low slopes, soil texture varied from silty clay to clay along the
481 forest-grassland gradient. This change was related to a thinning of the silty deposits overlaying the
482 clayey material derived from the marlstone. This may be due to initial land-use choices by farmers, to
483 land use after-effects, or both. Various arguments suggest that this gradient result from both.
484 Naturally, clayey soils downslope add value to the production of fodder while silty well-drained soils
485 upslope are better used for hardwood forest. However, ancient grasslands have been probably
486 repeatedly tilled, and tillage increase erosion. A strong effect of land use is suggested by (i) the
487 abruptness of the textural gradient compared to the very smooth topographic gradient, and (ii) the fact

488 that the soil texture of recent grasslands which developed for less than 150 years on former forests was
489 already enriched in clay compared to forests from which they derived.

490 In the study area, the present land use influenced soil acidity, soil organic content and concentration in
491 nutrients. Soil acidity decreased with the distance-to-edge in forest and increased in recent grassland
492 after deforestation. Soil organic content and nutrients generally increased with distance-to-edge in
493 grassland, decreased in case of afforestation and increased more or less rapidly after deforestation.
494 Besides these general trends, some exceptions were found, for example for K. For various reasons,
495 mainly related to differences in vegetation composition and management practices between forest and
496 grassland, soil properties change strongly according to both distance-to-edge and land-use change.
497 However, traces of former land uses persist on some soil properties. The legacy of ancient land-use
498 was stronger in case of deforestation than afforestation. The legacy of former forests in current
499 grasslands is a new result that has implications for conservancy of grasslands.

500

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510

511

512

513 **References**

- 514 Alstad, A.O., Damschen, E.I., Givnish, T.J., Harrington, J.A., Leach, M.K., Rogers, D.A., Waller,
515 D.M., 2016. The pace of plant community change is accelerating in remnant prairies. *Sci.*
516 *Adv.* 2, e1500975.
- 517 Antoine, P., Catt, J., Lautridou, J.-P., Sommé, J., 2003. The loess and coversands of northern France
518 and southern England. *J. Quat. Sci.* 18, 309–318.
- 519 Batey, T., 2009. Soil compaction and soil management – a review. *Soil Use Manag.* 25, 335–345.
- 520 Benjamini, Y., Hochberg, Y., 1995. Controlling the false discovery rate: A practical and powerful
521 approach to multiple testing. *J. R. Stat. Soc. Ser. B Methodol.* 57, 289–300.
- 522 Bezemer, T.M., Lawson, C.S., Hedlund, K., Edwards, A.R., Brook, A.J., Igual, J.M., Mortimer, S.R.,
523 Van Der Putten, W.H., 2006. Plant species and functional group effects on abiotic and
524 microbial soil properties and plant–soil feedback responses in two grasslands. *J. Ecol.* 94,
525 893–904.
- 526 Binkley, D., Fisher, R.F., 2013. *Ecology and management of forest soils*, 4th ed. Wiley, Hoboken, NJ.
- 527 Blondeel, H., Perring, M.P., Bergès, L., Brunet, J., Decocq, G., Depauw, L., Diekmann, M., Landuyt,
528 D., Liira, J., Maes, S.L., Vanhellefont, M., Wulf, M., Verheyen, K., 2018. Context-
529 dependency of agricultural legacies in temperate forest soils. *Ecosystems*.
- 530 Brêthes, A., 1976. *Catalogue des stations forestières du plateau lorrain*. ONF et INRA, internal
531 publication available to: <https://inventaire-forestier.ign.fr/spip.php?article783>.
- 532 BRGM, 2001. *Carte géologique de Château-Salins au 1/50 000*. BRGM Editions. Orléans.
- 533 Brubaker, S.C., Jones, A.J., Lewis, D.T., Frank, K., 1993. Soil properties associated with landscape
534 position. *Soil Sci. Soc. Am. J.* 57, 235–239.
- 535 Burst, M., Chauchard, S., Dupouey, J.-L., Amiaud, B., 2017. Interactive effects of land-use change
536 and distance-to-edge on the distribution of species in plant communities at the forest–
537 grassland interface. *J. Veg. Sci.* 28, 515–526.
- 538 Chantigny, M.H., 2003. Dissolved and water-extractable organic matter in soils: a review on the
539 influence of land use and management practices. *Geoderma* 113, 357–380.

540 Dupouey, J.-L., Dambrine, E., Laffite, J.-D., Moares, C., 2002. Irreversible impact of past land use on
541 forest soils and biodiversity. *Ecology* 83, 2978–2984.

542 Duvigneaud, J., 1967. Flore et végétation halophiles de la Lorraine orientale (Dép. Moselle, France).
543 *Mém. Soc. Roy. Bot. Belgique* 3, 1–122.

544 Ehrenfeld, J.G., Ravit, B., Elgersma, K., 2005. Feedback in the plant-soil system. *Annu. Rev. Environ.*
545 *Resour.* 30, 75–115.

546 Etienne, D., 2011. Les mardelles intra-forestières de Lorraine: origines, archives paléo-
547 environnementales, évolutions dynamiques et gestion conservatoire. PhD Thesis, University
548 of Nancy, FR.

549 Favre, C., Grel, A., Granier, E., Cosserat-Mangeot, R., Bachacou, J., Dupouey, J.-L., 2013.
550 Digitalisation des cartes anciennes. Manuel pour la vectorisation de l’usage des sols et le
551 géoréférencement des minutes 1:40000 de la carte d’Etat-Major. v.12.7.3. INRA Nancy.

552 Flinn, K.M., Marks, P.L., 2007. Agricultural legacies in forest environments: Tree communities, soil
553 properties, and light availability. *Ecol. Appl.* 17, 452–463.

554 Glatzel, G., 1991. The impact of historic land use and modern forestry on nutrient relations of Central
555 European forest ecosystems. *Fertil. Res.* 27, 1–8.

556 Graae, B.J., Sunde, P.B., Fritzboøger, B., 2003. Vegetation and soil differences in ancient opposed to
557 new forests. *For. Ecol. Manag.* 177, 179–190.

558 Green, B.H., 1990. Agricultural intensification and the loss of habitat, species and amenity in British
559 grasslands: a review of historical change and assessment of future prospects. *Grass Forage*
560 *Sci.* 45, 365–372.

561 Guo, L.B., Gifford, R.M., 2002. Soil carbon stocks and land use change: a meta analysis. *Glob.*
562 *Change Biol.* 8, 345–360.

563 Guo, L.B., Wang, M., Gifford, R.M., 2007. The change of soil carbon stocks and fine root dynamics
564 after land use change from a native pasture to a pine plantation. *Plant Soil* 299, 251–262.

565 Harrison, K.A., Bardgett, R.D., 2010. Influence of plant species and soil conditions on plant–soil
566 feedback in mixed grassland communities. *J. Ecol.* 98, 384–395.

567 Hassink, J., 1997. The capacity of soils to preserve organic C and N by their association with clay and
568 silt particles. *Plant Soil* 191, 77–87.

569 Hassink, J., 1994. Effects of soil texture and grassland management on soil organic C and N and rates
570 of C and N mineralization. *Soil Biol. Biochem.* 26, 1221–1231.

571 Hook, P.B., Burke, I.C., 2000. Biogeochemistry in a shortgrass landscape: Control by topography, soil
572 texture, and microclimate. *Ecology* 81, 2686–2703.

573 Houben, P., 2008. Scale linkage and contingency effects of field-scale and hillslope-scale controls of
574 long-term soil erosion: Anthropogeomorphic sediment flux in agricultural loess watersheds of
575 Southern Germany. *Geomorphology* 101, 172–191.

576 IUSS Working Group WRB, 2015. World Reference Base for soil resources 2014, update 2015.
577 International soil classification system for naming soils and creating legends for soil maps.
578 World Soil Resources Reports No. 106. FAO, Rome.

579 Jabiol, B., Brêthes, A., Brun, J.-J., Ponge, J.-F., Toutain, F., Zanella, A., Aubert, M., Bureau, F., 2009.
580 Typologie des formes d’humus forestières (sous climats tempérés). Association Française
581 pour l’Étude du Sol. *Référentiel pédologique 2008*, Quae, versailles, pp. 327-355.

582 Janssen, P., Bec, S., Fuhr, M., Taberlet, P., Brun, J.-J., Bouget, C., 2018. Present conditions may
583 mediate the legacy effect of past land-use changes on species richness and composition of
584 above- and below-ground assemblages. *J. Ecol.* 106, 306–318.

585 Jenny, H., 1941. *Factors of soil formation*. McGraw Hill, New York.

586 Jobbágy, E.G., Jackson, R.B., 2004. The uplift of soil nutrients by plants: Biogeochemical
587 consequences across scales. *Ecology* 85, 2380–2389.

588 Joimel, S., Cortet, J., Jolivet, C.C., Saby, N.P.A., Chenot, E.D., Branchu, P., Consalès, J.N., Lefort, C.,
589 Morel, J.L., Schwartz, C., 2016. Physico-chemical characteristics of topsoil for contrasted
590 forest, agricultural, urban and industrial land uses in France. *Sci. Total Environ.* 545–546, 40–
591 47.

592 Kepfer-Rojas, S., Verheyen, K., Johannsen, V.K., Schmidt, I.K., 2015. Indirect effects of land-use
593 legacies determine tree colonization patterns in abandoned heathland. *Appl. Veg. Sci.* 18,
594 456–466.

595 Koerner, W., Dupouey, J.L., Dambrine, E., Benoit, M., 1997. Influence of past land use on the
596 vegetation and soils of present day forest in the Vosges mountains, France. *J. Ecol.* 85, 351-
597 358.

598 Laganière, J., Angers, D.A., Paré, D., 2010. Carbon accumulation in agricultural soils after
599 afforestation: a meta-analysis. *Glob. Change Biol.* 16, 439–453.

600 Li, D., Niu, S., Luo, Y., 2012. Global patterns of the dynamics of soil carbon and nitrogen stocks
601 following afforestation: a meta-analysis. *New Phytol.* 195, 172–181.

602 McBratney, A.B., Mendonça Santos, M.L., Minasny, B., 2003. On digital soil mapping. *Geoderma*
603 117, 3–52.

604 McKenzie, N.J., Ryan, P.J., 1999. Spatial prediction of soil properties using environmental correlation.
605 *Geoderma* 89, 67–94.

606 Motzkin, G., Foster, D., Allen, A., Harrod, J., Boone, R., 1996. Controlling site to evaluate history:
607 vegetation patterns of a New England sand plain. *Ecol. Monogr.* 66, 345–365.

608 Mucina, L., Bültmann, H., Dierßen, K., Theurillat, J.-P., Raus, T., Čarni, A., Šumberová, K., Willner,
609 W., Dengler, J., García, R.G., Chytrý, M., Hájek, M., Di Pietro, R., Iakushenko, D., Pallas, J.,
610 Daniëls, F.J.A., Bergmeier, E., Santos Guerra, A., Ermakov, N., Valachovič, M., Schaminée,
611 J.H.J., Lysenko, T., Didukh, Y.P., Pignatti, S., Rodwell, J.S., Capelo, J., Weber, H.E.,
612 Solomeshch, A., Dimopoulos, P., Aguiar, C., Hennekens, S.M., Tichý, L., 2016. Vegetation of
613 Europe: hierarchical floristic classification system of vascular plant, bryophyte, lichen, and algal
614 communities. *Appl Veg Sci* 19, 3–264.

615 Murty, D., Kirschbaum, M.U.F., Mcmurtrie, R.E., MCGilvray, H., 2002. Does conversion of forest to
616 agricultural land change soil carbon and nitrogen? a review of the literature. *Glob. Change*
617 *Biol.* 8, 105–123.

618 Nakagawa, S., Schielzeth, H., 2013. A general and simple method for obtaining R^2 from generalized
619 linear mixed-effects models. *Methods Ecol. Evol.* 4, 133–142.

620 Nosetto, M.D., Jobbágy, E.G., Tóth, T., Bella, C.M.D., 2007. The effects of tree establishment on
621 water and salt dynamics in naturally salt-affected grasslands. *Oecologia* 152, 695–705.

622 Peña, E. de la, Baeten, L., Steel, H., Viaene, N., Sutter, N.D., Schrijver, A.D., Verheyen, K., 2016.
623 Beyond plant–soil feedbacks: mechanisms driving plant community shifts due to land-use
624 legacies in post-agricultural forests. *Funct. Ecol.* 30, 1073–1085.

625 Piessens, K., Honnay, O., Devlaeminck, R., Hermy, M., 2006. Biotic and abiotic edge effects in highly
626 fragmented heathlands adjacent to cropland and forest. *Agric. Ecosyst. Environ.* 114, 335–
627 342.

628 Ponge, J.-F., Jabiol, B., Gégout, J.-C., 2011. Geology and climate conditions affect more humus forms
629 than forest canopies at large scale in temperate forests. *Geoderma* 162, 187–195.

630 Schimel, D., Stillwell, M.A., Woodmansee, R.G., 1985. Biogeochemistry of C, N, and P in a soil
631 catena of the shortgrass steppe. *Ecology* 66, 276–282.

632 Schmidt, M., Jochheim, H., Kersebaum, K.-C., Lischeid, G., Nendel, C., 2017. Gradients of
633 microclimate, carbon and nitrogen in transition zones of fragmented landscapes – a review.
634 *Agric. For. Meteorol.* 232, 659–671.

635 Seibert, J., Stendahl, J., Sørensen, R., 2007. Topographical influences on soil properties in boreal
636 forests. *Geoderma* 141, 139–148.

637 Solon, J., Degórski, M., Roo-Zielińska, E., 2007. Vegetation response to a topographical-soil gradient.
638 *CATENA* 71, 309–320.

639 Swacha, G., Botta-Dukát, Z., Kački, Z., Pruchniewicz, D., Żołnierz, L., 2018. The effect of
640 abandonment on vegetation composition and soil properties in *Molinion* meadows (SW
641 Poland). *PLOS ONE* 13, e0197363.

642 Thimonier, A., Dupouey, J.L., Timbal, J., 1992. Floristic changes in the herb-layer vegetation of a
643 deciduous forest in the Lorraine Plain under the influence of atmospheric deposition. *For.*
644 *Ecol. Manag.* 55, 149–167.

645 Tilman, E.A., Tilman, D., Crawley, M.J., Johnston, A.E., 1999. Biological weed control via nutrient
646 competition: Potassium limitation of dandelions. *Ecol. Appl.* 9, 103–111.

647 Vanwallegem, T., Bork, H.R., Poesen, J., Dotterweich, M., Schmidtchen, G., Deckers, J., Scheers, S.,
648 Martens, M., 2006. Prehistoric and Roman gullying in the European loess belt: a case study
649 from central Belgium. *The Holocene* 16, 393–401.

- 650 von Oheimb, G., Härdtle, W., Naumann, P.S., Westphal, C., Assmann, T., Meyer, H., 2008. Long-
651 term effects of historical heathland farming on soil properties of forest ecosystems. *For. Ecol.*
652 *Manag.* 255, 1984–1993.
- 653 Wang, J., Fu, B., Qiu, Y., Chen, L., 2001. Soil nutrients in relation to land use and landscape position
654 in the semi-arid small catchment on the loess plateau in China. *J. Arid Environ.* 48, 537–550.

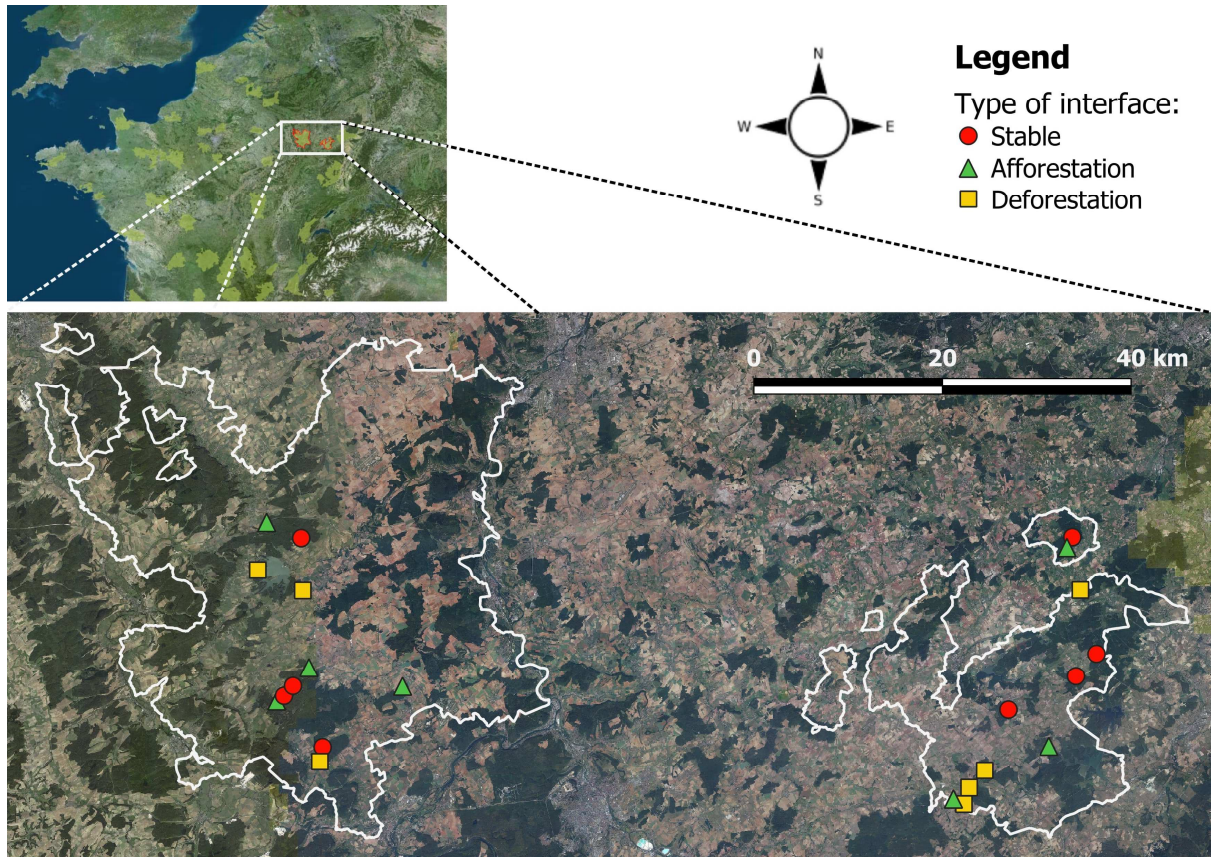


Fig. 1. Study area and location of the three types of forest-grassland interfaces studied. All the study sites (forest-grassland interfaces) were selected within the Lorraine Regional Natural Park in northeast France (inside the white lines). The red circles show the stable interfaces (ancient forest vs. ancient grassland, 8 sites), the green triangles show the afforestation interfaces (recent forest vs. ancient grassland, 7 sites) and the yellow squares show the deforestation interfaces (ancient forest vs. recent grassland, 7 sites).

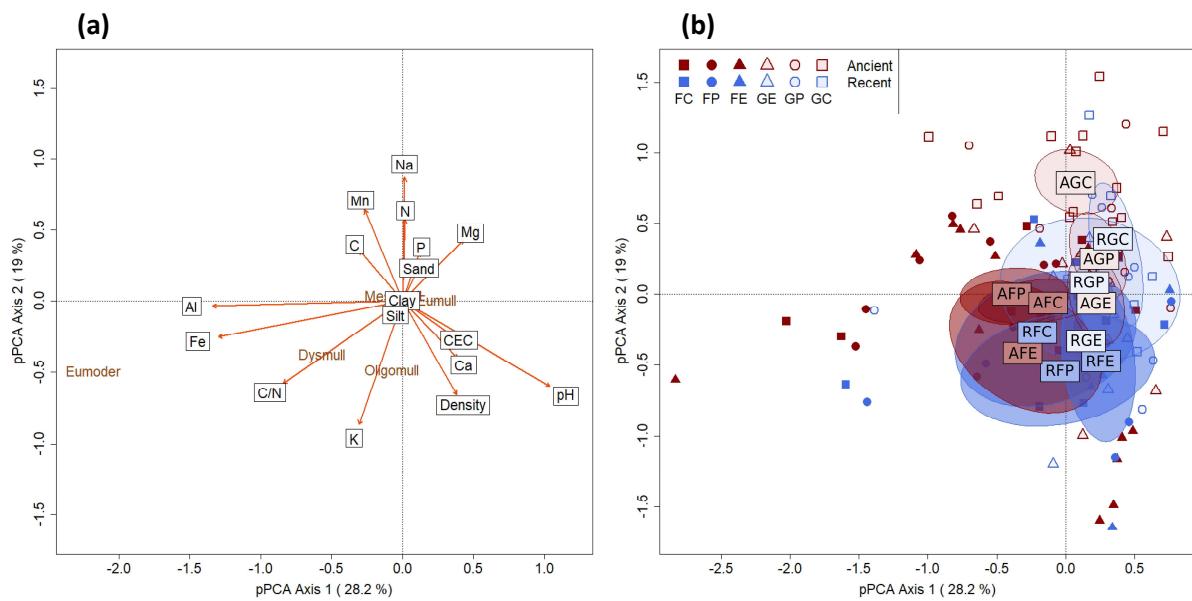


Fig. 2. Partial principal component analysis of soil properties. The effects of topography (factor with 4 levels) and soil texture (*i.e.* soil clay content) were partialled out before PCA. Projection of soil properties (see appendix B for their list) is shown in figure (a) and the humus types have been projected as supplementary variables. Plots are shown in figure (b) and dispersion ellipses are added for the interaction of the history with the position (12 levels) using a confidence interval of 0.95. Ellipse labels are the combination of history and plot position and were sometimes slightly shifted for better readability. History: A (ancient) and R (recent). Plot position: FC (forest core), FP (forest periphery), FE (forest edge), GE (grassland edge), GP (grassland periphery) and GC (grassland core).

Table 1. Goodness-of-fit of linear mixed models (LMMs) and analyses of variance of fixed effects for soil properties.

		LMM goodness-of-fit		Position		History:Position		Topography		Clay content	
		Marginal r ²	Conditional r ²	chi ²	p-value	chi ²	p-value	chi ²	p-value	chi ²	p-value
LMM1											
	Clay	0.32	0.84	25.1	***	28.9	***	16	**		
	Silt	0.3	0.83	23.2	***	35.5	***	13.9	**		
	Sand	0.08	0.81	4.9	ns	3.5	ns	10.1	*		
LMM2											
	Density	0.5	0.62	10.5	.	19.6	**	7	.	22.1	***
	C/N	0.61	0.79	134.7	***	13.6	*	1.2	ns	6.7	**
	C	0.64	0.73	10	.	14.4	*	2	ns	68.7	***
	N	0.78	0.82	45.7	***	13.4	*	3	ns	109.6	***
	P	0.25	0.76	35.2	***	6.3	ns	0.7	ns	7.2	**
	CEC	0.79	0.88	16.8	**	3.9	ns	3.4	ns	192.2	***
	Ca	0.61	0.82	13.3	*	2.8	ns	2.4	ns	101.7	***
	Mg	0.19	0.91	1.5	ns	2.5	ns	7.7	.	54.4	***
	Na	0.71	0.81	89.3	***	32.9	***	12.8	**	16.3	***
	K	0.5	0.57	79.2	***	8.2	ns	2.9	ns	25.5	***
	Fe	0.14	0.36	12.3	*	8.4	ns	4.1	ns	1	ns
	Mn	0.13	0.65	4.5	ns	6.5	ns	7.5	.	5.6	*
	Al	0.18	0.39	11.5	*	9.3	ns	2.1	ns	0.8	ns
	pH	0.26	0.5	15.6	**	3.9	ns	3.5	ns	10.3	**

Marginal r²: fixed effects only. Conditional r²: fixed and random (site in sector) effects. *** $P < 0.001$, ** $0.001 \leq P < 0.01$, * $0.01 \leq P < 0.05$ and '.' $0.05 \leq P < 0.1$.

Table 2. Contrast analyses of the land-use, edge and land-use change effects for 14 soil properties. The contrasts were performed between selected levels of the interaction of the history with the position (12 levels), which was one of the fixed effects of the linear mixed models (LMMs). The number of plots used in each contrast analysis is indicated in the 'Number of plots' line.

		Site effect	Land-use effect	Edge effect		Land-use change effect			
		Marginal r^2 / Conditional r^2	AFC-AGC	Forest	Grassland	Afforestation (recent forest)		Deforestation (recent grassland)	
				AFC-AFE	AGC-AGE	Change	Legacy	Change	Legacy
						RFCP-AGCP	RFCP-AFCP	RGCP-AFCP	RGCP-AGCP
Number of plots		15-15	15-15	15-15	14-30	14-30	14-30	14-30	
LMM1									
	Clay	62 %	***	***	***	*	**	*	***
	Silt	64 %	***	***	***	.	***	*	***
	Sand	90 %	ns	ns	ns	ns	ns	ns	ns
LMM2									
	Density	19 %	.	ns	**	**	ns	ns	**
	C/N	23 %	***	ns	**	***	.	***	ns
	C	12 %	*	ns	***	*	ns	ns	.
	N	5 %	***	ns	***	***	ns	ns	*
	P	67 %	*	ns	*	***	.	*	ns
	CEC	10 %	ns	ns	ns	ns	ns	ns	ns
	Ca	26 %	ns	ns	ns	ns	ns	ns	ns
	Mg	79 %	ns	ns	ns	ns	ns	ns	ns
	Na	12 %	***	ns	***	***	ns	.	***
	K	12 %	**	**	.	***	ns	**	ns
	Fe	61 %	.	ns	ns	ns	ns	*	ns
	Mn	80 %	ns	ns	ns	ns	ns	ns	ns
	Al	54 %	ns	ns	ns	ns	ns	ns	ns
	pH	48 %	ns	.	ns	ns	ns	*	ns

Planned contrasts. Land use effect: ancient forest core vs. ancient grassland core (AFC-AGC). Edge effect: ancient forest core vs. ancient forest edge (AFC-AFE) and ancient grassland core vs. ancient grassland edge (AGC-AGE). Land-use change effect: recent forest core and periphery vs. ancient grassland core and periphery (RFCP-AGCP), recent forest core and periphery vs. ancient forest core and periphery (RFCP-AFCP), recent grassland core and periphery vs. ancient forest core and periphery (RGCP-AFCP) and recent grassland core and periphery vs. ancient grassland core and periphery (RGCP-AGCP). *** $P < 0.001$, ** $0.001 \leq P < 0.01$, * $0.01 \leq P < 0.05$ and '.' $0.05 \leq P < 0.1$.

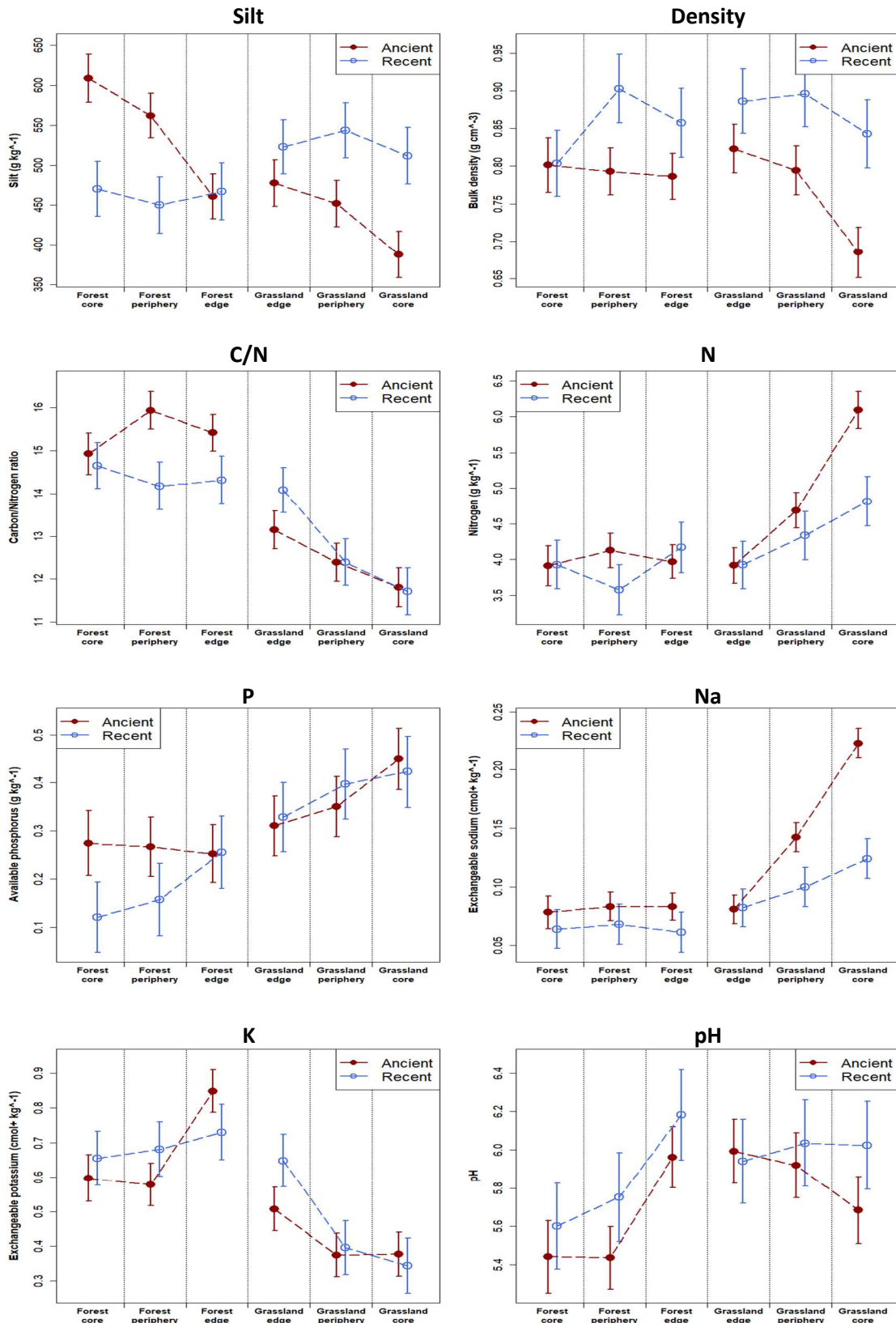


Fig. 3. Least-squares means and standard errors of eight soil properties by position along forest-grassland interfaces for each land-use history. Least-squares means, which are means adjusted for the means of others factors, were calculated using only the topography as a categorical control variable (factor with 4 levels) for the silt content (LMM1), and in addition using the soil texture *i.e.* the soil clay content as a numerical control variable for other soil properties (LMM2).