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1 **The role of mobilisation and delivery processes on contrasting dissolved**
2 **nitrogen and phosphorus exports in groundwater fed catchments**

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16 **Abstract**

17 Diffuse transfer of nitrogen (N) and phosphorus (P) in agricultural catchments is controlled by
18 the mobilisation of sources and their delivery to receiving waters. While plot scale
19 experiments have focused on mobilisation processes, many catchment scale studies have
20 hitherto concentrated on the controls of dominant flow pathways on nutrient delivery. To
21 place mobilisation and delivery at a catchment scale, this study investigated their relative
22 influence on contrasting nitrate-N and soluble P concentrations and N:P ratios in two shallow
23 groundwater fed catchments with different land use (grassland and arable) on the Atlantic
24 seaboard of Europe. Detailed datasets of N and P inputs, concentrations in shallow
25 groundwater and concentrations in receiving streams were analysed over a five year period
26 (October 2010 – September 2015). Results showed that nitrate-N and soluble P
27 concentrations in shallow groundwater give a good indication of stream concentrations,
28 which suggests a dominant control of mobilisation processes on stream exports. Near-
29 stream attenuation of nitrate-N (-30%), likely through denitrification and dilution, and
30 enrichment in soluble P (+100%), through soil-groundwater interactions, were similar in both
31 catchments. The soil, climate and land use controls on mobilisation were also investigated.
32 Results showed that grassland tended to limit nitrate-N leaching as compared to arable land,
33 but grassland could also contribute to increased P solubilisation. In the context of land use
34 change in these groundwater fed systems, the risk of pollution swapping between N and P
35 must be carefully considered, particularly for interactions of land use with soil chemistry and
36 climate.

37 **Keywords**

38 Agricultural catchments, land use, climate, nutrients, nitrate, phosphorus

39 **1. Introduction:**

40 Land-to-water transfer of nitrogen (N) and phosphorus (P) is a major concern worldwide as
41 excessive concentrations of these two nutrients cause eutrophication in freshwater and

42 marine ecosystems (Conley et al., 2009). Diffuse emissions from agricultural origin can
43 represent a significant contribution to annual N and P loads in rivers (Dupas et al., 2015a).

44 Mitigation measures to decrease N and P emissions from agricultural landscapes must rely
45 on underpinning research to understand mobilisation and delivery mechanisms in order to be
46 effective and several mechanisms have been identified (Lloyd et al., 2016; Mellander et al.,
47 2012; Outram et al., 2016). It is generally accepted that N is prone to vertical leaching from
48 the soil, through the unsaturated zone down to the saturated zone. It can then be transferred
49 to surface waters, mainly as nitrate, via groundwater (Legout et al., 2007; Mellander et al.,
50 2014; Molenat et al., 2008). Nitrate mobilised below the rooting zone may be subject to
51 denitrification, which generally takes place within anoxic groundwaters, riparian wetlands and
52 the hyporheic zone (Anderson et al., 2014; McAleers et al., 2017; Oehler et al., 2007).
53 Mobilised nitrate may also have long time lags between mobilisation and emergence, due to
54 the potentially long transit time of water in the unsaturated and saturated zone (Fovet et al.,
55 2015; Hrachowitz et al., 2010). Phosphorus has a higher adsorption affinity with the soil
56 compared to nitrate; as a result it is less soluble and is less prone to leaching. A general
57 understanding is that P is typically transferred to surface waters via surface pathways, either
58 in soluble or particulate form (Sharpley et al., 2008). However, several studies have
59 highlighted the possibility of soluble P leaching into shallow groundwater and subsequent
60 lateral transfer to surface water as a dominant mechanism in groundwater-fed catchments
61 (e.g. Dupas et al., 2015c; Haygarth et al., 1998; Holman et al., 2010; Mellander et al., 2016;
62 van der Salm et al., 2011).

63 Nitrate and soluble P leaching is controlled by several factors, some of which are
64 manageable and some of which are inherent properties of soils and climate. Nitrate leaching
65 is controlled by i) the balance between fertiliser inputs and crop uptake on an annual basis; ii)
66 soil mineralisation, which increases when the soil C:N ratio is low and when favourable
67 moisture and temperature conditions are met (Rodrigo et al., 1997); iii) temporal mismatches
68 between crop uptake capacity and high nitrate concentration in the soil combined with high

69 drainage potential (particularly in autumn, Dupas et al., 2015d). In temperate regions,
70 grasslands have been shown to have a better capacity to take up N from the soil throughout
71 the year, particularly in the autumn period, compared to cropland (McDowell et al., 2014;
72 Moreau et al., 2012), although urine patches in grazed grassland can be hotspots of nitrate
73 leaching.

74 Soluble P leaching is controlled by i) the soil P content, as determined by soil P tests; ii) soil
75 chemistry, particularly the abundance of iron (Fe) and aluminium (Al) oxides, which are
76 important adsorption sites in acidic soils (Daly et al., 2015; Schoumans and Chardon, 2015);
77 iii) temporal variation of soil pH, redox state (Henderson et al., 2012) and drying-rewetting or
78 freezing-thawing cycles (Blackwell et al., 2010); iv) organic matter (OM) content, which may
79 control the formation of Fe-bound organic P colloids that are more mobile than truly soluble P
80 (Granger et al., 2007), and which influences the concentration of dissolved organic matter
81 (DOM) that can compete with P for adsorption sites (Kang et al., 2009); v) land use,
82 particularly the presence of grassland, which is often associated with high OM content in soil
83 and which may release root exudates that stimulate the microbial biomass and bring P into
84 solution (Roberts et al., 2013). Grassland may also favour preferential flow paths through
85 macropores, enhancing nitrate and soluble P transport (Djordjic et al., 2004; Gachter et al.,
86 1998).

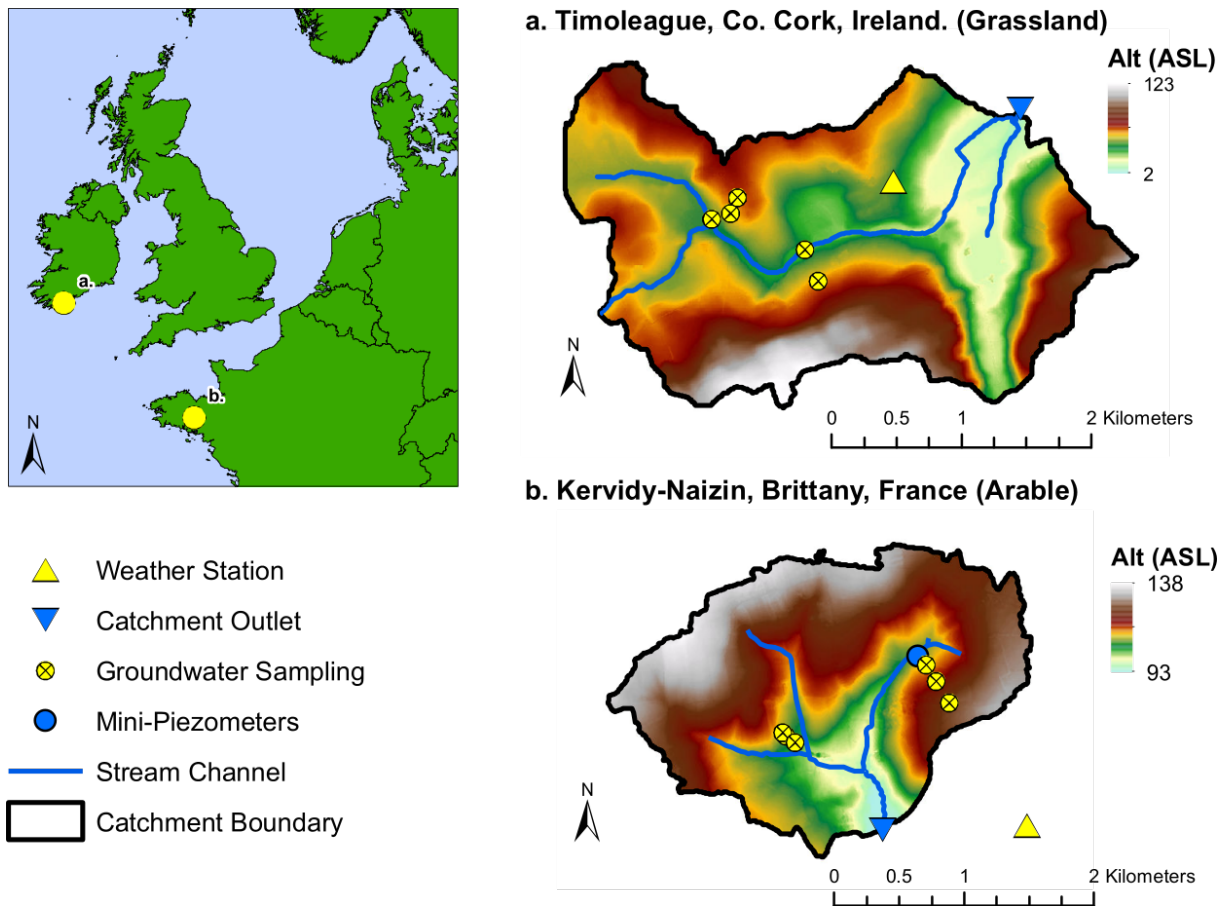
87 Several studies comparing the export behaviour of N and P in agricultural catchments along
88 the Atlantic seaboard of Europe have highlighted the crucial role played by dominant flow
89 pathways on export patterns (e.g. Lloyd et al., 2016; Mellander et al., 2012; Outram et al.,
90 2016). In general, well-drained catchments are prone to nitrate leaching and subsequent
91 transfer via groundwater, whereas poorly-drained catchments can transfer large masses of P
92 via surface flow pathways (Jordan et al., 2012; McDowell et al., 2014). However, few studies
93 have compared catchments with similar dominant flow pathways to investigate the role of
94 mobilisation and delivery mechanisms at this small catchment scale.

95 The main aim of this paper, therefore, was to investigate mobilisation processes (i.e.
96 leaching) as an important role in the transfer continuum of nitrate-N and soluble P in
97 groundwater-fed catchments. This was undertaken by monitoring the nutrient transfer
98 continuum, from source to mobilisation and from mobilisation to delivery (Haygarth et al.,
99 2005), in two intensively studied agricultural catchments (with similar groundwater fed
100 systems) on the Atlantic seaboard of Europe. A detailed survey of N and P input and content
101 in soil was used to characterise the sources, nitrate-N and soluble P concentration in shallow
102 groundwater was used to characterise the amount of nutrient being mobilised below the
103 rooting zone and comparison of groundwater and stream nitrate-N and soluble P
104 concentration was used to characterise delivery. The specific objectives were to i) investigate
105 the respective influence of mobilisation and delivery mechanisms on contrasting nitrate-N
106 and soluble P concentrations and N :P ratios in the two catchments and ii) analyse the
107 respective controls of soil properties, climate and land use on nitrate-N and soluble P
108 mobilisation. We use this analysis to explore some implications for the future trajectory (i.e.
109 evolution under changing external forces) of diffusion pollution in these and similar
110 catchment types under land use and climate change scenarios.

111 **2. Materials and methods**

112 **2.1. Study areas**

113 This investigation took place in two intensively farmed catchments in the pedo-climatic zone
114 of the Atlantic seaboard of Europe, in Western France and South-western Ireland (Fig. 1).



115

116 **Figure 1: Catchment site map. Location of the Timoleague catchment in Ireland and**
 117 **the Kervidy-Naizin catchment in France and location of the monitoring sites.**

118 The French catchment, Kervidy-Naizin, belongs to the AgrHyS environmental research
 119 observatory, where the impacts of agriculture and climate change on water quality are
 120 studied (Aubert et al., 2013). Kervidy-Naizin is a 5 km² catchment drained by a stream of
 121 second Strahler order. Climate is temperate oceanic, with mean ± standard deviations of
 122 annual rainfall, temperature and discharge of 938 ± 218 mm, 11.3 ± 0.5 °C and 359 ± 206
 123 mm, respectively (October 2010 – September 2015). Topography is gentle, with an elevation
 124 range of 93-135 m above sea level. The bedrock consists of impervious, locally fractured
 125 Brioverian schists and is capped by several metres of unconsolidated weathered material
 126 and silty, loamy soils. The hydrological behaviour is dominated by the development of a
 127 water table that varies seasonally along the hillslope. In the upland domain, consisting of
 128 well-drained soils (86% of the catchment area), the water table remains below the soil

129 surface throughout the year, varying in depth from 1 m to more than 8 m. In the wetland
130 domain, developed near the stream and consisting of hydromorphic soils (14% of the
131 catchment area), the water table is shallower, remaining near the soil surface generally from
132 October to April each year (Molenat et al., 2008). Artificial drainage represents <10% of the
133 surface area but is generally ineffective at lowering the water table in the wetland area
134 because the drains are in disrepair. The baseflow index (Institute of Hydrology, 1980) during
135 the October 2010 – September 2015 study period was 0.7. The land use is mostly
136 agriculture, specifically arable crops and confined animal production (dairy cows and pigs). A
137 farm survey conducted in 2013 led to the following land use zones: 71% winter cereals and
138 maize, 16% grassland and 13% other crops (rape seed, vegetables). Soil tillage (10 cm to 40
139 cm) is practised for all crops (including temporary grassland) one to three times a year;
140 cultivation primarily takes place in spring for summer crops and autumn for winter crop; all
141 fields are ploughed every one to three years. Annual nutrient inputs were typically (on
142 average for the crop rotation present in 2013, i.e., for a 1 to 5 year period) 212 kg N ha⁻¹
143 (71% organic, 29% chemical) and 62 kg P ha⁻¹ (83% organic, 17% chemical). Direct inputs
144 from grazing livestock were included in the organic N and P input term, and atmospheric
145 deposition primarily comes from fertiliser volatilisation, thus it was not included in the N input
146 term to avoid double-counting. Pig slurry dominates organic effluent application, followed by
147 cow manure; spreading period stretches from February to May and organic effluent
148 application is prohibited from July to January (March for maize).

149 The Irish catchment, Timoleague, is part of the Agricultural Catchment Programme, which is
150 a European Union Nitrates Directive evaluation experiment established to monitor
151 agricultural practices under water quality policies in Ireland (Wall et al., 2011). Timoleague is
152 a 7.6 km² catchment drained by a stream of second Strahler order. Climate is temperate
153 oceanic, with mean ± standard deviations of annual rainfall, temperature and discharge of
154 1047 ± 92 mm, 10.1 ± 0.5 °C and 589 ± 106 mm, respectively (October 2010 – September
155 2015). The topography is rolling to flat with an elevation range of 17-127 m above sea level.

156 The lithology consists of Old Red Sandstone and mudstone of the Castlehaven formation.
157 Belowground flow paths are likely concentrated in the high permeability layers and along the
158 contacts of different layer types and in possible fractures or faults. The aquifer is unconfined
159 and classified as productive with a secondary permeability. In the upland domain, consisting
160 of well-drained brown earths (Cambisols 87% of the catchment area), the water table
161 remains below the soil surface throughout the year, varying in depth from 2 m to more than
162 10 m. In the wetland domain, developed near the stream and consisting of poorly drained
163 Gleysols, alluvials and peat soils (13% of the catchment area), the water table is shallower,
164 remaining near the soil surface generally from October to April each year (Mellander et al.,
165 2014). The baseflow index (Institute of Hydrology, 1980) during the October 2010 –
166 September 2015 study period was 0.7. According to the Irish Soil Information System the soil
167 texture is primarily loam. The land use is mainly dairy production with an average livestock
168 density of 1.9 livestock units (LU) ha⁻¹. The animals are housed over winter and graze over
169 the spring to autumn period. The surface area is represented by 77% grassland and 11%
170 mixture of cereals (e.g. spring barley, spring wheat and winter oilseed rape), forage maize
171 and root crops (e.g. fodder beet). Based on 2010 and 2011 nutrient use census, the average
172 annual input rates were approximately 310 kg N ha⁻¹ (41% organic and 59% chemical) and
173 54 kg P ha⁻¹ (92% organic, 8% chemical). As in the Kervidy-Naizin catchment, direct inputs
174 from grazing livestock were included in the organic N and P input term, and atmospheric
175 deposition was not included. Timings of these fertiliser inputs were mainly skewed to spring,
176 when growth demands were highest, as 73% of N and 75% of P was applied by the end of
177 June. Peak application rates (average 60kg N ha⁻¹ and 24 kg P ha⁻¹ in the available organic
178 form and 51 kg N ha⁻¹ and 2 kg P ha⁻¹ in the mineral form) were in April, which is required at
179 this time to match the nutrient demand as grass growth typically peaks in Ireland during this
180 period.

181 **2.2. Soil, stream and groundwater monitoring**

182 In the Kervidy-Naizin catchment, soil samples were taken using a triangular network
183 sampling method at 89 sampling points in 2013 (Matos-Moreira et al., 2017). Seven soil
184 cores (0-15 cm) were taken within a 1 m radius around each of the 89 sampling point and
185 were composited. Samples were air-dried, sieved (2 mm mesh) and stored at room
186 temperature before analysis. Dyer P was determined by using 20 g l⁻¹ citric acid with a
187 soil:solution ratio of 1:5 (NF X 31-160). Oxalate extractable Al and Fe were determined after
188 extraction with 0.0866 mmol l⁻¹ oxalic acid + 0.1134 mmol l⁻¹ ammonium oxalate using a
189 soil:solution ratio of 1:40 (w:v), in the dark and at pH 3 (Tamm, 1922). Organic matter,
190 nitrogen, and carbon contents were determined by dry combustion (1000°C, NF ISO 13878,
191 NF ISO 10694).

192 Also in the Kervidy-Naizin catchment, stream discharge was measured every minute at a
193 gauging station with a float operated sensor upstream of a rectangular weir and a data logger
194 (OTT Thalimedes). The weather station (Cimel Enerco 516i) was located 1.1 km from the
195 catchment outlet (Fig. 1b) and recorded hourly rainfall and temperature. The stream chemical
196 monitoring consisted of a daily sampling performed manually at approximately the same time
197 (17:00 local time). For each sample, two aliquots were filtered directly on-site for nitrate-N
198 analysis (0.22 µm) and soluble reactive P (SRP) analysis (0.45 µm). Nitrate was determined
199 as N by ionic chromatography (DIONEX DX 100), with a precision of ± 2.5%. Soluble
200 reactive P was determined colorimetrically by reaction with ammonium molybdate, with a
201 precision of ± 0.004 mg l⁻¹. For SRP, samples were analysed every 6 days from October
202 2010 to September 2013 and every day (i.e., like nitrate) from October 2013 to September
203 2015 (Minaudo et al., in review). Other N and P species were not considered in this study
204 because they represent minor fractions of dissolved N and P: nitrite and ammonium
205 concentrations determined in grab samples were generally below the detection limits,
206 respectively < 0.07 mg l⁻¹ and < 0.04 mg l⁻¹ (n=147, unpublished data). Soluble reactive P
207 represented >80 % of total dissolved P (n=25, Dupas et al., 2015b).

208 Shallow groundwater samples were collected every 3 - 4 months in 10 wells along two
209 transects (Fig. 1b). Screening depths were 1.5 – 3 m for the wells located downslope and 4 –
210 8 m or 6 – 10 m for those upslope, i.e. in the groundwater fluctuation zone. At these depths,
211 groundwater is typically aerobic, poor in organic compounds or other electron donors, limiting
212 significant denitrification (Molénat et al., 2002). This is in contrast to deeper pathways,
213 whereby anaerobic conditions promote the dissolution of solid phase bacterial energy
214 sources (Mn^{2+} , Fe^{2+} , S^-), which can in turn drive autotrophic denitrification (Pauwels et al.,
215 1998). A shallow groundwater comparison is therefore appropriate to characterise nitrate
216 mobilisation (i.e. leaching) rather than nitrate reduction (i.e. denitrification) processes
217 (McAleer et al., 2017). Only nitrate-N was analysed in the groundwater samples. Soluble
218 reactive P was not investigated in these wells because it was assumed to be low in this
219 catchment. To test the latter assumption and to investigate SRP mobilisation below the
220 rooting zone, one shallow piezometer (screening depth = 1 m) was placed at the footslope of
221 one transect; it was sampled weekly during the 2013-2014 water year.

222 Finally, the Kervidy-Naizin catchment was equipped with mini-piezometers placed within the
223 soil in triplicate (screening depth: 4 – 5 cm) in two adjacent plots in a cropland field and a
224 grassland plot. The mini-piezometers were made from 15 cm Polyvinyl chloride tubes
225 (diameter = 5 cm) closed at the bottom, with three slits at 4 cm, 4.5 cm and 5 cm below the
226 soil surface. These mini-piezometers were sampled with a syringe, five times during the
227 2013-2014 water year, when water was present in the mini-piezometers (from January to
228 March). Samples were analysed for SRP to compare the effect of land use on SRP
229 mobilisation with similar land and climate characteristics.

230 In the Timoleague catchment, soil samples were taken using a regular grid sampling scheme
231 at 27 sampling points in 2012. As in the Kervidy-Naizin catchment, a grid sampling scheme
232 was chosen to cover the variability of factors influencing soil properties (such as soil
233 classification, land use and topography). Forty soil cores (0-10 cm) were taken from a 25 m x
234 25 m area around each sampling point and composited. Samples were oven-dried (40°C),

235 sieved (2 mm mesh) and stored at room temperature before analysis. Influence of oven-
236 drying of Timoleague soils as compared to air-drying of Kervidy-Naizin was assumed to have
237 a minor influence on subsequent analyses given the low temperature used. A modified
238 Mehlich (Mehlich, 1984) method was used to extract P, Al and Fe from a 2 g sub-sample of
239 soil with Mehlich3 (M3) reagent (0.2 M CH₃COOH + 0.25 M NH₄NO₃ + 0.015 M NH₄F + 0.13
240 M HNO₃ + 0.001 M EDTA) at a 1:10(w/v) soil to solution ratio for 5 min. According to Sarr et
241 al. (2007), this method extracts 1.5 times less P from the soil than the Dyer method used in
242 the Kervidy-Naizin catchment. Organic matter was estimated as the loss-on-ignition from 4 g
243 soil sub-samples (500 °C for 16h). Total carbon and N were determined by dry combustion
244 using a CN LECO FP2000 analyser (LECO Corporation, St. Joseph, MI, USA),

245 In the Timoleague catchment, stream discharge was measured every 10 minutes with a
246 vented-pressure instrument (OTT Orpheus Mini) installed in a stilling well adjacent to non-
247 standard Corbett flat-v weirs. The weather station (Campbell Scientific BWS200) was located
248 in the central part of the catchment (Fig. 1a) and recorded rainfall and temperature on a 10
249 min basis. For both catchments, daily mean values of discharge and weather data are
250 presented. Also in Timoleague, the river outlet was equipped with bankside analysers that
251 monitored total oxidized nitrogen (TON) and total reactive P (TRP) concentrations. TON was
252 measured in situ by a UV probe (Hach-Lange Nitratax), with a measuring range of 0.1 – 50
253 mg l⁻¹. It was assumed that TON was equivalent to nitrate-N (Melland et al., 2012). The P
254 instrumentation (Hach-Lange Phosphax-Sigma) measured TRP every 20 min by colorimetry
255 using the molybdate–antimony method (DIN EN ISO 6878), with a measuring range of 0.01 –
256 5 mg l⁻¹ (Wall et al., 2011). It was also assumed that TRP was approximately equivalent to
257 SRP since the flow-weighted mean SRP was previously reported to account for 98–99% of
258 the flow weighted mean TRP in another (similar groundwater-fed) Irish grassland catchment
259 (Shore et al., 2014). Similar to the Kervidy-Naizin catchment, nitrite and ammonium
260 concentrations determined in grab samples were low, respectively 0.023 ± 0.003 mg l⁻¹ and

261 0.183 ± 0.017 mg l⁻¹ (n=256 and 298, unpublished data). Soluble reactive P represented >70
262 % of total dissolved P (n=187, unpublished data).

263 Groundwater was sampled monthly in six multilevel monitoring wells along two transects
264 (Mellander et al., 2014; McAleer et al., 2017; Fig. 1a). The wells consisted of three
265 piezometers that screened different depths; only the piezometers that screened comparable
266 depths to the arable catchment (i.e. in the groundwater fluctuation zone: 2.8 – 5.5 m for the
267 wells located downslope and 5.5 – 12 m for those upslope) were considered here. All
268 samples were filtered immediately after sampling (0.45 µm) and analysed for nitrate-N and
269 SRP.

270 **2.3. Data analysis**

271 This analysis focuses on the dissolved N and P form because previous studies have shown
272 that particulate transfer involved different mechanisms from the input-groundwater-stream
273 continuum investigated here, e.g. stream bank erosion, surface runoff, etc. (Dupas et al.,
274 2015b; Sherriff et al., 2015). Despite the monitoring protocols differing slightly between the
275 two catchments, it was possible to extract similar metrics from each, and perform a
276 comparison of the factors controlling nitrate-N and soluble P mobilisation and delivery, based
277 on the following assumptions and data treatments:

- 278 - It was assumed that the soil sampling protocols were comparable in both catchments.
279 Because tillage > 15 cm was performed at least once a year in each field in the
280 Kervidy-Naizin catchment, which homogenises the soil properties over this depth, a
281 soil sampling depth of 10 cm in Timoleague and 15 cm in Kervidy-Naizin can be
282 compared.
- 283 - Based on previous studies, it was assumed that TON and nitrate-N were comparable
284 (Melland et al., 2012) and that TRP and SRP were similarly comparable (Shore et al.,
285 2014).

286 - Sub-hourly stream data from the Timoleague catchment were averaged every hour
 287 and then sub-sampled at the time of sampling in the Kervidy-Naizin catchment (17:00
 288 local time) to standardise the time-series between the catchments.

289 - Average (\pm standard deviations) values were considered for groundwater data as
 290 differences between catchments were larger than spatial variability within each
 291 catchment. For comparison of nitrate-N concentrations, five upslope piezometers
 292 were selected in the Kervidy-Naizin catchment and three upslope piezometers were
 293 selected in the Timoleague catchment. The reason for excluding downslope
 294 piezometers was because they were placed in vegetated buffer strips in the Kervidy-
 295 Naizin catchment (whereas this study aimed to investigate mobilisation below
 296 cropland fields in this catchment) and also placed in denitrifying riparian zones in both
 297 catchments (whereas this study aimed to investigate mobilisation prior to riparian
 298 attenuation). For comparison of SRP concentrations, one shallow piezometer placed
 299 below the rooting zone at the footslope of a cropland field was selected in the
 300 Kervidy-Naizin catchment and five piezometers located at similar topographic
 301 positions were selected in the Timoleague catchment. The piezometers selected
 302 were representative of each catchment in the sense that they were placed in cropland
 303 fields in Kervidy-Naizin and in grassland fields in Timoleague, and they had similar
 304 screening depths and topographic positions along the hillslopes.

305 The data analyses were fourfold:

306 - Comparing the inter- and intra-annual variability of hydroclimatic variables and nitrate-
 307 N and soluble P loads in the two catchments. Annual nitrate-N and soluble P loads
 308 (L) were estimated with the discharge-weighted concentration method (Moatar et al.,
 309 2013):

$$L = K * \left(\frac{\sum_{i=1}^n C_i * Q_i}{\sum_{i=1}^n Q_i} \right) * \bar{Q}$$

310 where C_i is the instantaneous nitrate-N or soluble P concentration, Q_i is the mean
311 daily discharge associated with C_i , \bar{Q} is mean annual discharge, K is a conversion
312 factor to obtain load in $\text{kg ha}^{-1} \text{ yr}^{-1}$ and n represents the number of (C_i, Q_i) pairs for
313 each water year .

314 - Comparing time-weighted nitrate-N and soluble P concentrations in groundwater and
315 flow-weighted nitrate-N and soluble P concentrations in stream water between the
316 two catchments. In addition, variability in concentration and discharge data was
317 quantified using the coefficient of variation, defined as the ratio of the standard
318 deviation to the mean. This metric has been used in previous studies to characterise
319 the chemodynamic or chemostatic character of solute export (according to the
320 relative dispersion of concentration data compared to discharge; Musolff et al., 2015).
321 The coefficient of variation was calculated both on an annual mean and daily basis, to
322 describe both inter-annual and intra-annual variability.

323 - Calculating N:P ratios for fertiliser inputs, groundwater concentrations and stream
324 flow-weighted concentrations to investigate the respective control of the sources, the
325 mobilisation mechanisms and delivery on contrasting nitrate-N and soluble P
326 concentrations and load in the two catchments. To relate groundwater and stream
327 concentrations to current inputs (averaged over a five year period) we assumed
328 steady state conditions in the two catchments, which is acceptable here as transit
329 times are < 10 years in both catchments and there has been no major change in land
330 use and/or N and P inputs in the recent period (Molenat et al., 2008).

331 - Comparing soil, land use and climate characteristics to investigate their respective
332 effect of nitrate-N and soluble P mobilisation/leaching.

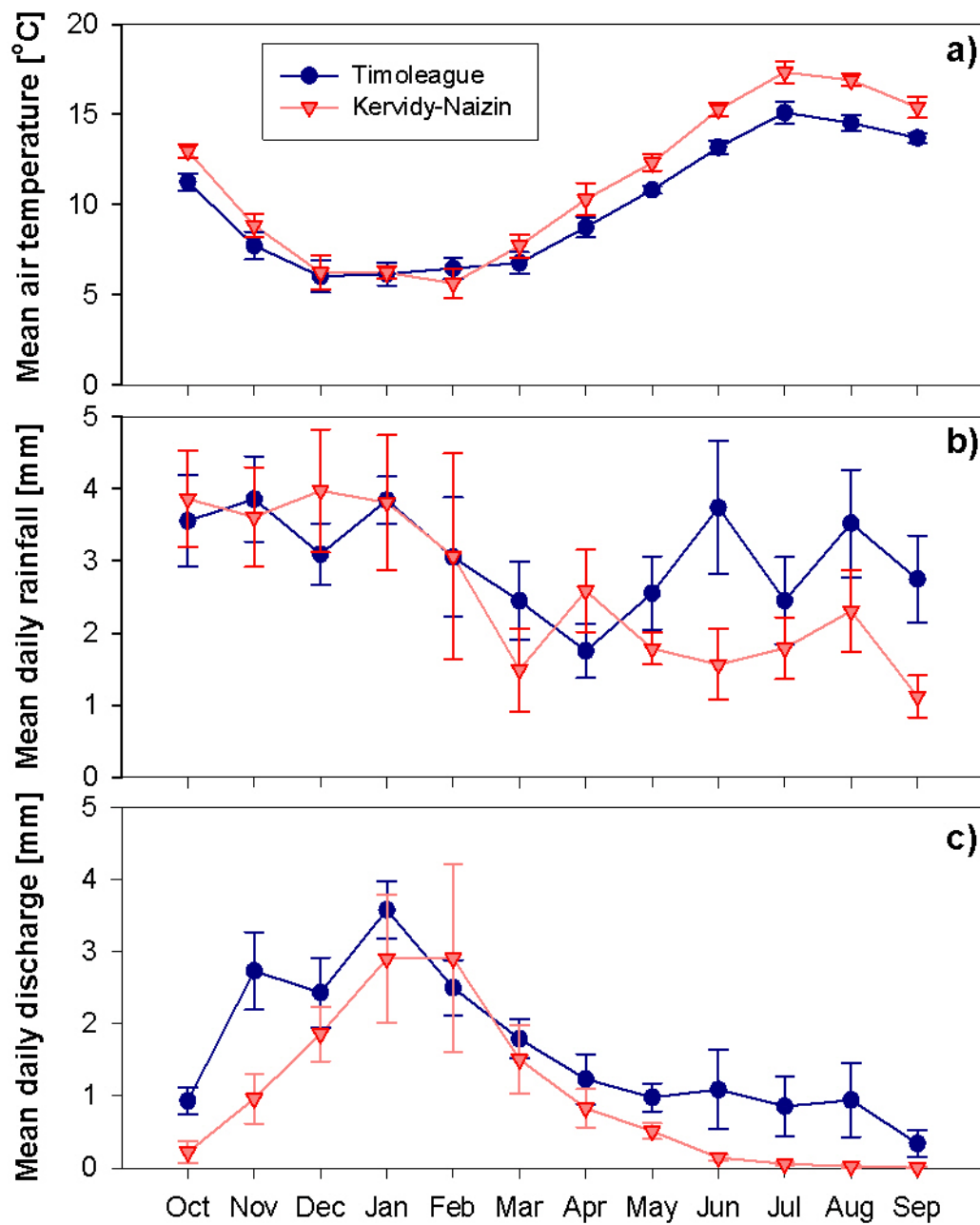
333 All graphical figures and statistical tests were performed with SigmaPlot (Systat Software,
334 San Jose, CA). Unless stated otherwise, two-tailed t-tests were employed for comparisons,
335 and normal distributions were checked using a Shapiro-Wilk test.

336 **3. Results and discussion**

337 **3.1. Temporal variability in the hydroclimate, nitrogen and phosphorus loads**

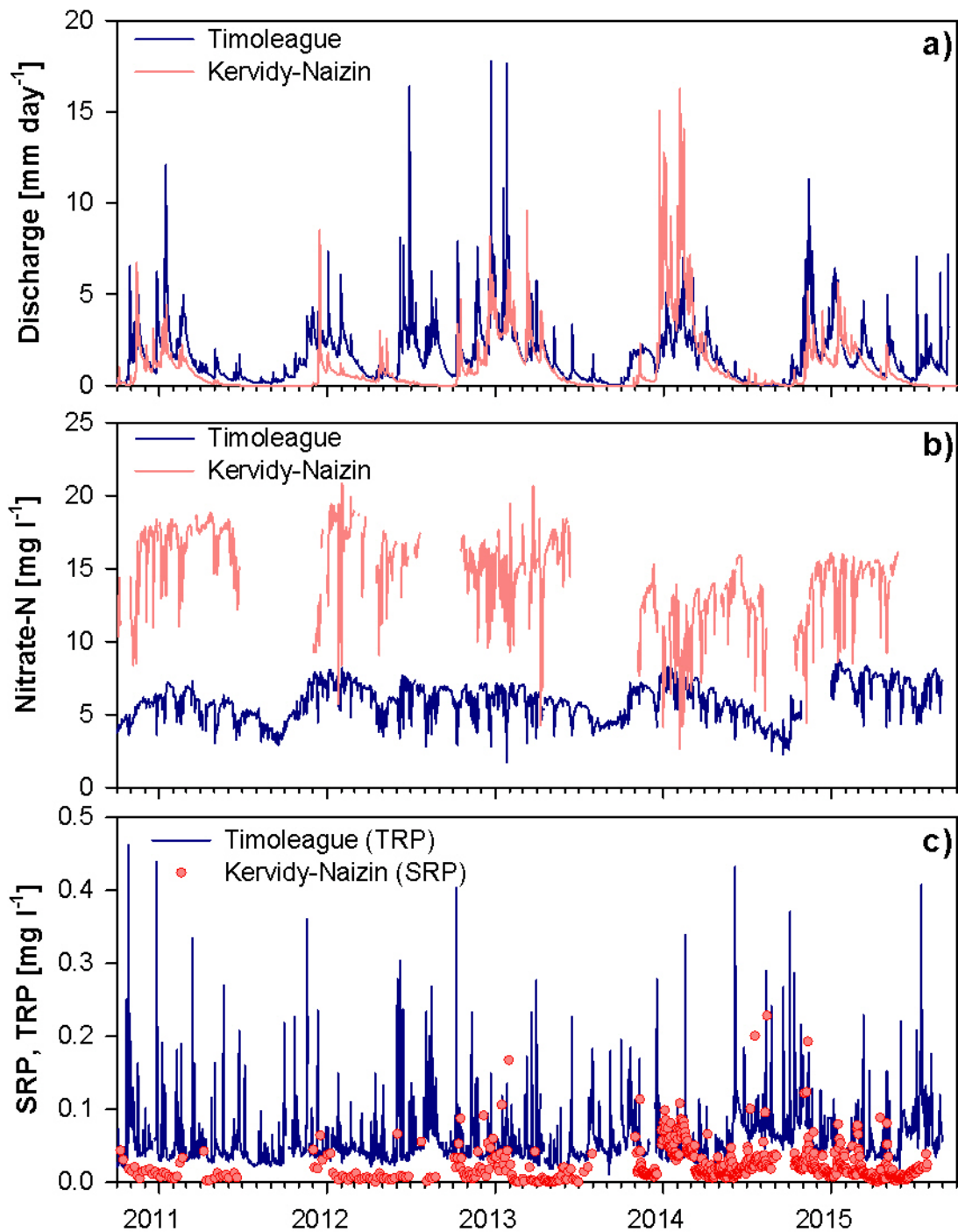
338 The Kervidy-Naizin and the Timoleague catchments appeared to be affected by similar
339 seasonal weather patterns but different inter-annual weather patterns. Both catchments were
340 characterised by mild temperatures and the occurrence of rainfall throughout the year (Table
341 1; Fig. 2 a and b); mean annual temperature was slightly but significantly higher in Kervidy-
342 Naizin (11.3 ± 0.5 °C versus 10.1 ± 0.5 °C, $p < 0.05$) but the annual cumulated rainfall was
343 not significantly different in both catchments (938 ± 218 mm in Kervidy-Naizin versus $1047 \pm$
344 92 mm in Timoleague, $p > 0.05$). In both catchments, stream discharge showed a strong
345 seasonality with high flow during the winter period and, for the most part, low flow during
346 summer. In the Timoleague catchment, the stream flowed throughout the year while the
347 stream was dry during one to two months in the summer period every year in the Kervidy-
348 Naizin catchment (Fig. 2c and 3a). This difference is due to higher summer
349 evapotranspiration in Kervidy-Naizin (3.7 ± 0.3 mm day⁻¹ versus 2.7 ± 0.3 mm day⁻¹ from
350 June to August, $p < 0.05$) and higher summer rainfall in Timoleague (3.2 ± 1.2 mm day⁻¹
351 versus 1.8 ± 0.6 mm day⁻¹ from June to August, $p < 0.05$). Annual discharge was on average
352 lower, but more variable, in the Kervidy-Naizin catchment than in the Timoleague catchment
353 (359 ± 206 mm versus 589 ± 106 mm); the larger inter-annual variability in the Kervidy-Naizin
354 catchment is reflected by the larger estimated coefficient of variation (57% versus 18%).

355 **[Please insert Table 1 here]**



356

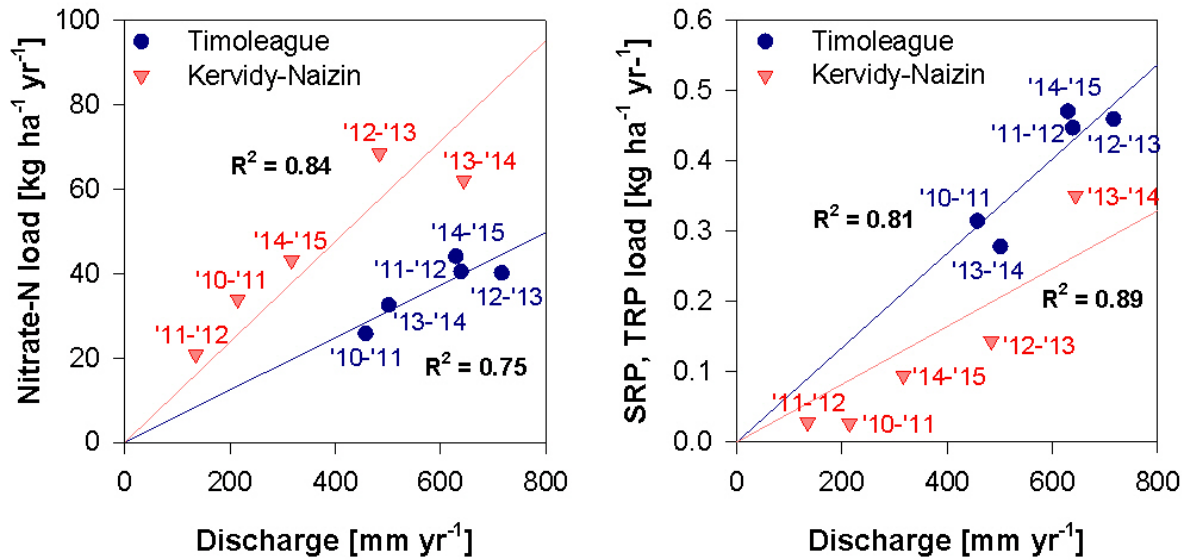
357 **Figure 2: Monthly averaged air temperature (a), daily rainfall (b) and discharge (c) over**
 358 **the period October 2010 – September 2015 in Timoleague and Kervidy-Naizin**
 359 **catchments. Error bars represent one standard deviation (n = 5 years).**



360

361 **Figure 3: Time series of daily discharge (a), nitrate concentration (b), and soluble**
 362 **phosphorus concentrations (c) at the outlet of Timoleague and Kervidy-Naizin**
 363 **catchments. Soluble phosphorus concentration in Kervidy-Naizin was only measured**
 364 **on a 6 day frequency from October 2010 to September 2013 and every day from**
 365 **October 2013 to September 2015.**

366 Annual soluble P loads were, on average, also higher in the Timoleague catchment ($0.39 \pm$
367 $0.09 \text{ kg TRP ha}^{-1} \text{ yr}^{-1}$ versus $0.13 \pm 0.13 \text{ kg SRP ha}^{-1} \text{ yr}^{-1}$ in the Kervidy-Naizin catchment),
368 but annual nitrate-N loads were on average higher in the Kervidy-Naizin catchment (46 ± 20
369 $\text{ kg N ha}^{-1} \text{ yr}^{-1}$ versus $37 \pm 7 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ in the Timoleague catchment). Similar to annual
370 discharge, nutrient annual loads were more variable in the Kervidy-Naizin catchment than in
371 the Timoleague catchment (Fig. 4); this is reflected by the larger estimated coefficient of
372 variation in the Kervidy-Naizin catchment (43% versus 20% for nitrate-N load and 103%
373 versus 23% for soluble P load). Only the difference between soluble P annual loads was
374 statistically significant for the five years of study ($p < 0.05$, $n = 5$). Flow-weighted mean
375 annual soluble P concentration was significantly higher in the Timoleague catchment ($66.6 \pm$
376 $7.2 \mu\text{g P l}^{-1}$ versus $29.3 \pm 15.8 \mu\text{g P l}^{-1}$ in the Kervidy-Naizin catchment, $p < 0.05$, $n = 5$) while
377 flow-weighted mean annual nitrate-N concentration was significantly higher in the Kervidy-
378 Naizin catchment ($13.7 \pm 2.5 \text{ mg N l}^{-1}$ versus $6.2 \pm 0.6 \text{ mg N l}^{-1}$ in the Timoleague catchment,
379 $p < 0.05$, $n = 5$). In both catchments, the coefficients of variation estimated for flow-weighted
380 mean annual concentration were lower than that for annual discharge: 53% and 11% for
381 soluble P in Kervidy-Naizin and Timoleague, respectively; and 18% and 10% for nitrate-N in
382 Kervidy-Naizin and Timoleague respectively. The lower variability of concentrations as
383 compared to discharge reveals a biogeochemical stationarity in these catchments, also
384 termed chemostasis (Basu et al., 2010). This chemostatic character observed in many
385 managed catchments worldwide has been previously attributed to the legacy of
386 anthropogenic N and P inputs accumulated within the catchments (Basu et al., 2010; Musolff
387 et al., 2015).

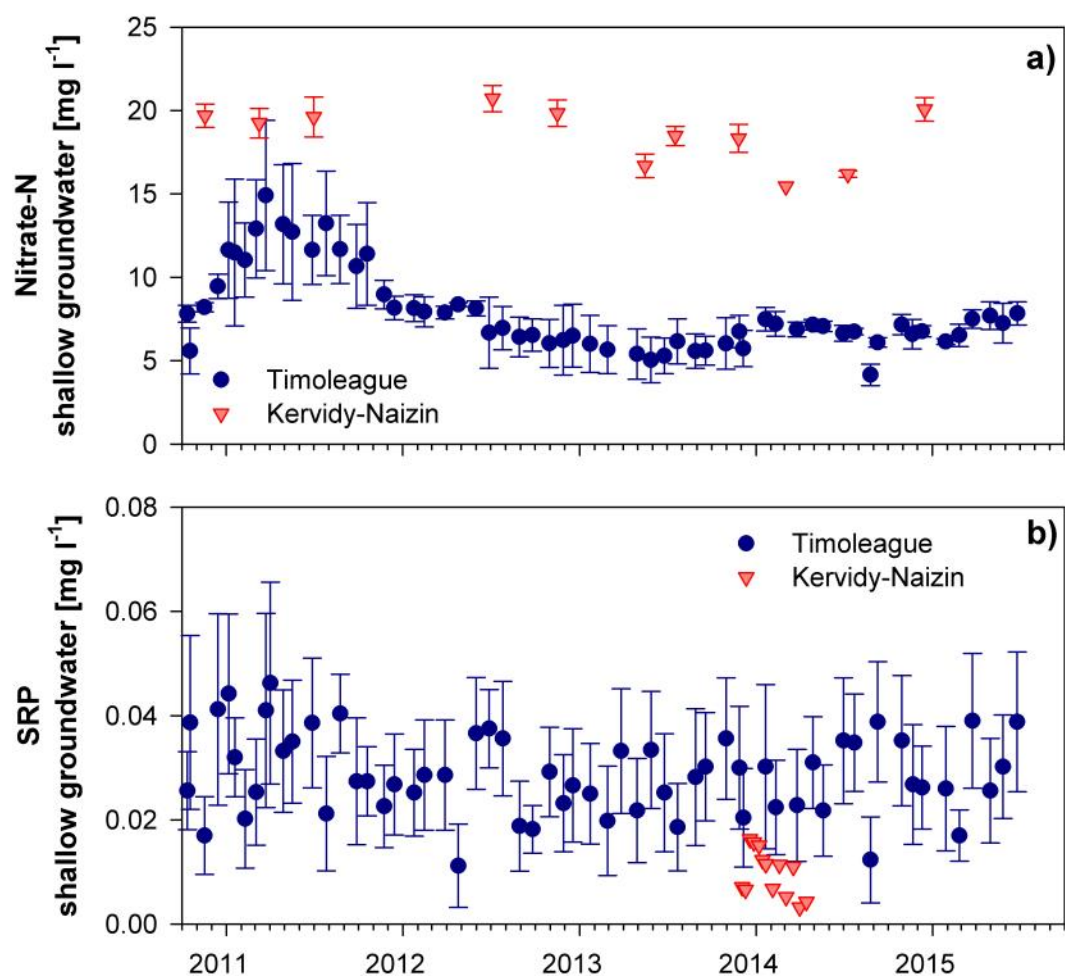


388

389 **Figure 4: Total annual nitrate (a) and soluble phosphorus (b) loads versus total annual**
 390 **discharge for Timoleague and Kervidy-Naizin catchments. The slope coefficients are**
 391 **significantly different from zero ($p < 0.05$).**

392 The positive linear relationship between annual discharge and nitrate-N and soluble P loads
 393 showed that, on an annual basis, exports from the two catchments were transport-limited
 394 processes (Basu et al., 2010; Mellander et al., 2014; Molénat et al., 2008; Musolff et al.,
 395 2015). The transport limitation evidenced from annual discharge – annual load plots is
 396 consistent with the chemostatic character of the catchment, because both transport limitation
 397 and chemostasis imply large nutrient storage within the catchments. The slope of the
 398 discharge-load relationship was steeper in the Kervidy-Naizin catchment for nitrate-N and
 399 steeper in the Timoleague catchment for soluble P, which reflects the contrasting flow-
 400 weighted mean concentrations. The transport limitation evidenced from annual loads and the
 401 difference in slopes between the two catchments imply larger nitrate-N storage in Kervidy-
 402 Naizin and larger storage of mobile P in the Timoleague catchment, which was confirmed by
 403 groundwater concentrations presented in Fig. 5 (mean concentrations: 20.8 mg N l⁻¹ versus
 404 8.5 mg N l⁻¹, and 12.4 µg P l⁻¹ versus 33.1 µg P l⁻¹ in Kervidy-Naizin and Timoleague,
 405 respectively). The storage of nitrate-N and soluble P in the aquifers, i.e. “anthropogenic

406 legacy of accumulated nutrient sources” (Basu et al., 2010), is likely the result of long term
407 leaching processes from the soil (Musolff et al., 2015). As expected, groundwater
408 concentrations were relatively stable in time, because they integrate nutrient leaching from
409 several fields within the hillslopes where the piezometer transects were located, and because
410 they integrate the response of cropping systems over multiannual time scales. An exception
411 to these stable groundwater concentrations was observed in the nitrate-N time series in the
412 Timoleague catchment, which exhibited higher concentration than usual during the autumn
413 2010 - autumn 2011 period (Fig. 5a). A previous study showed that this was the result of
414 grassland ploughing and reseeded in the vicinity of a monitored piezometer (Mellander et al.,
415 2014), and the effect of this management operation appeared visible until mid-2013. In
416 contrast, groundwater nitrate-N and soluble P dynamics in the Kervidy-Naizin catchment and
417 soluble P dynamics in the Timoleague catchment could not be related to known management
418 events.

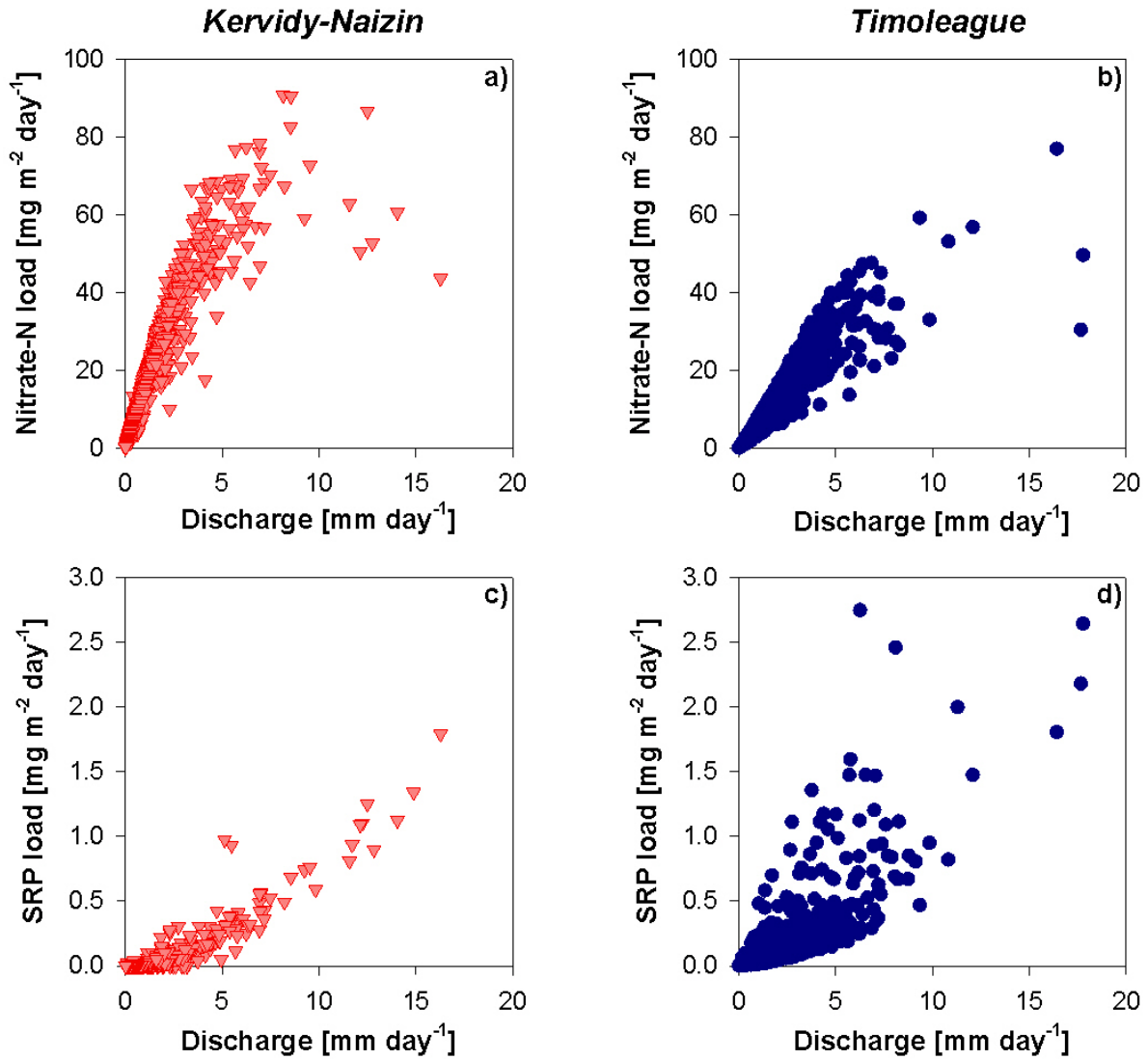


419

420 **Figure 5: Time series of nitrate (a) and soluble phosphorus (b) concentration in**
 421 **groundwater in Timoleague and Kervidy-Naizin catchments. Nitrate concentrations**
 422 **were measured in upslope piezometers (n=3 in Timoleague and n=5 in Kervidy-Naizin).**
 423 **Soluble phosphorus concentrations were measured in downslope shallow**
 424 **piezometers (n=5 in Timoleague and n=1 in Kervidy-Naizin). Error bars represent one**
 425 **standard deviation.**

426 On a daily basis, the coefficients of variation for soluble P concentration (109% in Kervidy-
 427 Naizin and 76% in Timoleague) and nitrate-N concentration (20% in Kervidy-Naizin and 22%
 428 in Timoleague) were lower than for discharge (173% in Kervidy-Naizin and 102% in
 429 Timoleague), similar to observations of annual mean discharge and concentration. All

430 coefficients of variation calculated on a daily basis were higher than those calculated on an
431 annual basis, most likely because intra-annual weather variability is larger than inter-annual
432 weather variability. Also on a daily basis, an inflexion in the discharge-load relationship
433 indicates a degree of supply-limitation for nitrate-N when daily discharge exceeded
434 approximately 6 mm d^{-1} in both catchments (Fig. 6 a and b). In contrast, the discharge-load
435 plots showed an increase in soluble P supply when daily discharge exceeded approximately
436 6 mm d^{-1} in the Kervidy-Naizin catchment, and a large variation in the Timoleague catchment
437 (Fig. 6 c and d). This change in discharge-loads relationship when discharge exceeded 6 mm
438 d^{-1} may be explained by the increasing activity of shallow and overland flow pathways during
439 storm events, which connect compartments of the catchment rich in soluble P and poor in
440 nitrate-N (Heathwaite and Dils, 2000). It is a primary cause of the dilution events observed in
441 the nitrate-N time series and the accretion events observed in the soluble P time series for
442 both catchment (Fig. 3 b and c).



443

444 **Figure 6: Daily nitrate (a, b) and soluble phosphorus (c, d) loads versus daily**
 445 **discharge for Timoleague and Kervidy-Naizin catchments.**

446

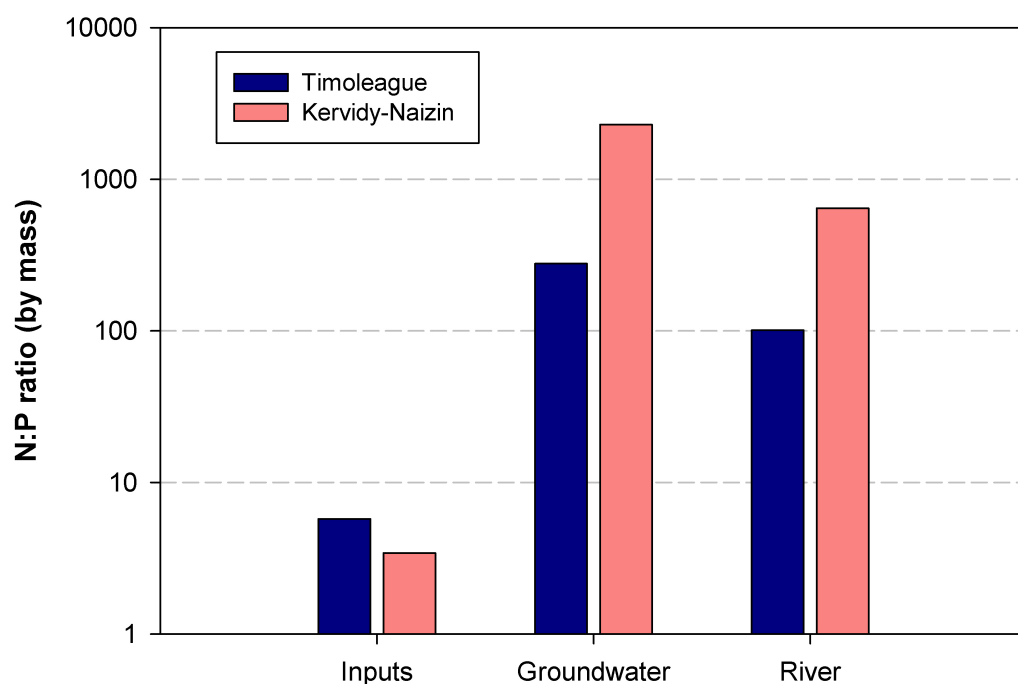
3.2. Nitrogen and phosphorus monitoring along the nutrient transfer continuum

447 Nitrogen chemical and organic inputs were 46% higher in the Timoleague catchment
 448 compared to the Kervidy-Naizin catchment, but flow-weighted mean nitrate-N concentration
 449 in the stream was 61% lower. Additionally, P chemical and organic inputs were 15% higher in
 450 the Kervidy-Naizin catchments, but flow-weighted mean soluble P concentration was 60%
 451 lower. Thus there was no direct relationship between nutrient total inputs and exports in
 452 soluble forms. Similar to P inputs, soil P content was higher in the Kervidy-Naizin catchment

453 than in the Timoleague catchment: the average Dyer P in the former was 265.2 mg P kg⁻¹
454 and the average Mehlich 3 P in the latter was 76.4 mg P kg⁻¹, i.e. approximately 114.6 mg P
455 kg⁻¹ Dyer P, considering that the Mehlich 3 method extracts 1.5 times less P than the Dyer
456 method (Sarr et al., 2007). Thus the discrepancy previously observed between P sources
457 and P exports was still valid if the soil P content is considered a source factor in addition to P
458 inputs.

459 This discrepancy between nutrient inputs and soluble exports was further highlighted by N:P
460 ratios. The N:P ratio of inputs was 1.7 times higher in the Timoleague catchment compared
461 to the Kervidy-Naizin catchment, while the N:P ratio of soluble exports in the stream was 6.4
462 times lower (Fig. 7). However, the decrease in the N:P ratio between groundwater and the
463 stream was similar in both catchments, by a factor of 3.6 and 2.7 in the Kervidy-Naizin and
464 the Timoleague catchments, respectively. This reflected a decrease of nitrate-N
465 concentration by 34% and 27%, respectively, between the groundwater and the river and an
466 increase of soluble P concentration by 136% and 101%.

467 In summary, the catchment with the highest nitrate-N concentration in the groundwater also
468 had the highest nitrate-N concentration in the stream, but a lower rate of N inputs, and
469 similarly the catchment with the highest soluble P concentration in the groundwater also had
470 the highest soluble P concentration in the stream, but a lower rate of P input. Therefore, no
471 direct relationship could be evidenced between nutrient inputs and exports in soluble form.
472 The decrease in nitrate-N concentration between groundwater and the stream might be
473 explained by attenuation processes such as dilution in non-fertilised near-stream areas
474 and/or denitrification within riparian wetlands and the hyporheic zone (Anderson et al., 2014;
475 Oehler et al., 2007). The increase in soluble P concentration between the groundwater and
476 the stream might be explained by additional transfer to the river via overland flow and/or via
477 the interception of shallow groundwater with the soil in riparian wetlands (Dupas et al.,
478 2015b, 2016).



479

480 **Figure 7: N:P ratios along the nutrient transfer continuum. “Inputs” were calculated**
 481 **from chemical and organic N and P fertiliser application. “Groundwater” was**
 482 **calculated from mean nitrate-N and soluble P groundwater concentration. “River” was**
 483 **calculated from nitrate-N and soluble P river loads (October 2010 – September 2015).**

484 **3.3. Comparison of factors controlling N and P leaching**

485 Table 2 summarises the soil-climate-land use factors that were compared in this study to
 486 explain the different amounts of nitrate-N and soluble P leaching in Kervidy-Naizin and
 487 Timoleague. All factors (except source factors) that are known to increase nitrate-N leaching
 488 were higher in the Kervidy-Naizin catchment and all factors (except source factors) that are
 489 known to increase soluble P leaching were higher in the Timoleague catchment, and this is
 490 reflected in the catchment outputs respectively.

491 In temperate climates the autumn-winter conditions are critical in determining nitrate-N
 492 leaching because of the risk of a temporal mismatch between the presence of high
 493 concentrations in the soil and plant uptake capacity (Moreau et al., 2012) combined with high

494 drainage potential. The soil organic N represents an important N storage pool (Sebilo et al.,
495 2013) which can subsequently release nitrate-N through mineralisation, especially when the
496 soil C:N ratio is low. Although the soil C:N ratio was similar in the two catchments, higher
497 autumn temperature in Kervidy-Naizin may lead to higher mineralisation rates than in
498 Timoleague (Rodrigo et al., 1997; Table 2). The main crops in the Kervidy-Naizin catchment,
499 winter cereals and maize, do not take up N efficiently in the autumn – winter period (Dupas et
500 al., 2015d). Furthermore, grasslands in the Timoleague catchment have a seemingly better
501 capacity to take up N from the soil throughout the year, particularly in the autumn – winter
502 period (McDowell et al., 2014) and reduce N leaching compared to cropland (Moreau et al.,
503 2012; Rode et al., 2009), although urine patches in grazed grassland can be hotspots of
504 nitrate leaching..

505 The abundance of Al oxide is an indicator of a soil's adsorption capacity with P (Daly et al.,
506 2015; Schoumans and Chardon, 2015). Extractable Al in the Kervidy-Naizin catchment
507 (using oxalate extraction) was 3.4 times higher than in Timoleague (using Mehlich 3
508 extraction), while previous studies have shown that the difference in extraction rates between
509 the two methods was less than 2 (Sims et al., 2002; Zhang et al., 2005). Hence the
510 maximum P sorption capacity of soils in the Kervidy-Naizin catchment was probably higher
511 than that of soils in the Timoleague catchment. The abundance of Fe oxide has also been
512 used as an indicator of a soil's adsorption capacity (e.g. Schoumans and Chardon, 2015) but
513 it can also favour the formation of colloidal Fe-P forms, which can increase leaching (Daly et
514 al., 2015; Mellander et al., 2016; Pautler and Sims, 2000). Extractable Fe in the Kervidy-
515 Naizin catchment (using oxalate extraction) was 10.5 times higher than in Timoleague (using
516 Mehlich 3 extraction), while previous studies have shown that the difference in extraction
517 rates between the two methods varied between ca. 5 - 12 (Sims et al., 2002; Zhang et al.,
518 2005). Hence the influence of Fe oxides in soils on P leaching between the two catchments
519 is ambiguous and it is difficult to conclude which catchment has the highest abundance of
520 extractable iron due to different extraction methods. The abundance of high organic matter

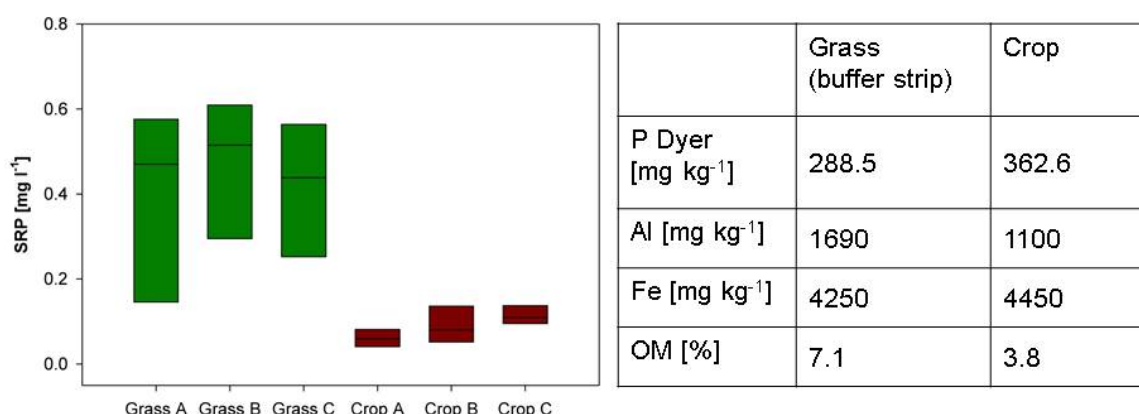
521 content in soils may increase P leaching by competing with P adsorption sites (Kang et al.,
522 2009) and favour the formation of easily leachable organic colloids (Granger et al., 2007).
523 These may be destabilised as rainwater with a lower ionic strength than the soil water
524 reaches the subsoil (Ilg et al., 2005). Organic matter content in the Timoleague catchment
525 was 74% higher than in Kervidy-Naizin. Therefore, with the soil data available, Timoleague
526 exhibited a higher risk of P leaching regarding all the soil characteristics available for this
527 comparative study. Subsoil properties have also been found to be critical for P leaching, and
528 particularly in sandy soils with high P sources, high degrees of P saturation and low sorption
529 capacities (Andersson et al., 2015). The subsoils have not been investigated in this study but
530 may have a role in the higher P leaching in Timoleague.

531 The dominance of grassland in the Timoleague catchment could also increase the risk of P
532 leaching compared to the mainly arable Kervidy-Naizin catchment. Presence of grassland
533 increases organic matter content, which was previously identified as increasing the risk of P
534 leaching (Djodjic et al., 2004; Gachter et al., 1998; Haygarth et al., 1998). In particular, a
535 review by Darch et al. (2014) has highlighted the potentially high contribution of soluble
536 organic P to total soluble P leaching in temperate agricultural soils (up to 80%), especially
537 when a rainfall occurs after slurry application. Although soluble organic P was not considered
538 in the current study, it may have contributed to the SRP measured in the groundwater and
539 the stream (after hydrolysis). Grassland may also favour the formation of macropores, which
540 increases the risk of preferential flow, because grassland is not frequently disturbed by tillage
541 thus biologically formed macropores (such as by roots or earthworms) are preserved (Simard
542 et al., 2000). Furthermore, root exudates from grassland may stimulate the activity of soil
543 microbes increasing P solubilisation (Roberts et al., 2013; Stutter et al., 2009) and drying-
544 rewetting cycles may release soluble organic and inorganic P from the microbial biomass
545 (Blackwell et al., 2010; Dupas et al., 2015c; Gu et al., 2017).

546 Because soil, climate and land use factors covary in the two catchments studied, Kervidy-
547 Naizin data from mini-piezometers placed in two adjacent plots with contrasting land uses but

548 similar soil chemistry and climate (Fig. 8) help to interpret the specific effect of land use on P
 549 solubilisation. The two plots exhibited similar soil P content and extractable Al and Fe and
 550 sampling points were at similar positions along the hillslope. However, organic matter content
 551 in the grassland plot was 87% higher than in the arable cropland plot. Results showed that P
 552 solubilisation was 4.7 times higher in the grassland plot compared to the cropland plot
 553 (Mann-Whitney test, $p < 0.05$; Fig. 8), either because of a direct effect of the plant type or via
 554 an increase in organic matter content. In the Timoleague catchment, where inputs of organic
 555 P are applied during grazing and spreading, leaching of soluble P is possibly even higher
 556 (Darch et al., 2014).

557 **[Please insert Table 2 here]**



558
 559 **Figure 8: Soluble Reactive Phosphorus (SRP) concentration in the soil solution**
 560 **sampled from two adjacent plots (grass and crop) at 4-5 cm (mini-piezometers in**
 561 **triplicate: A, B, C) in the Kervidy-Naizin catchment. Box-plots represent the median**
 562 **and 10th-90th percentile of five sampling dates during the 2013-2014 water year for each**
 563 **replicate. The table provides the mean value for each plot of soil Dyer P, aluminium,**
 564 **iron and organic matter contents.**

565
 566 **3.4. Catchment trajectories**

567 The catchment monitoring programmes both in Western France and South-western Ireland
568 aim to provide underpinning knowledge on N and P transfer processes to support
569 sustainable agricultural practices in a context of adaptive management, climate and land use
570 changes. In Western France, research and management efforts to reduce nutrient pollution
571 have long been directed towards N rather than P. This is because Western France is a
572 coastal region where the risk of green algae development in several bays (where N is
573 deemed to be the limiting nutrient) is the main concern for policy makers (Levain et al.,
574 2015). Conversion of cropland to grassland has been identified previously as an efficient way
575 of reducing nitrate-N leaching while maintaining good economic performance (Moreau et al.,
576 2012). The present comparative study with a grassland catchment in South-western Ireland
577 tends to confirm the beneficial role of grassland in decreasing nitrate-N losses. However, this
578 comparison also raises the question of a risk of pollution swapping, with a possible increase
579 of P mobilisation if arable land is converted into (intensive) grassland. In this regard, the
580 influence of organic matter on Al/Fe complexes will be important to determine. The
581 ambivalent role of (intensive) grassland in controlling N and P mobilisation has recently been
582 highlighted at the field and farm scale in the United-Kingdom (Peukert et al., 2014). In a
583 scenario where P mobilisation and transfer increase because of management measures that
584 only target nitrate-N mitigation and reduction of coastal green tides, eutrophication of
585 adjacent freshwaters is a possibility.

586 The Timoleague catchment also drains into estuarine waters with both N and P deemed to
587 be limiting nutrients in the Irish context (Longphuir et al., 2015). Studies in this catchment
588 have highlighted how reductions in mean flow-weighted P concentrations have occurred in
589 the quickest (i.e. shallowest) flow pathways due to reduced P source pressures, while
590 maintaining milk production yields at the highest level (Murphy et al., 2015). While this
591 indicates an increased catchment resilience, most likely related to reductions in the highest
592 soil P concentrations and conforming to regulations on nutrient management (Wall et al.,
593 2011), care is required on the trajectory of changes that might be observed under future land

594 use changes. For example, scenarios to increase milk production under grassland and also
595 under existing regulations may require grasslands to increase grass yield.

596 The potential risk of increasing nutrient losses under land use change or intensification may
597 be exacerbated by climate change in both Western France and South-western Ireland. With
598 increasing mean annual temperature, and particularly autumn temperature, N mineralisation
599 is expected to increase (Rodrigo et al., 1997) causing higher nitrate-N leaching if plant
600 uptake rates and timing are not adequately adapted. Concerning soluble P, the increasing
601 occurrence of wetter periods following drier periods (Ockenden et al., 2016) is also expected
602 to increase solubilisation (Blackwell et al., 2010; Dupas et al., 2015c; 2016) and potentially
603 leaching.

604 **4. Conclusion**

605 In two groundwater-driven and intensively farmed catchments on the Atlantic seaboard of
606 Europe (Western France and Southwest Ireland), mobilisation processes appeared to play a
607 dominant role in the contrasting export of nitrate-N and soluble P from land to the receiving
608 streams. Nitrate-N leaching was mainly controlled by mineralisation and the timing of plant
609 uptake: grasslands have a longer growth period and better utilisation of N than cropland,
610 hence grasslands reduced nitrate-N leaching. Soluble P leaching was mainly controlled by
611 the soil adsorption capacity and land use, with an increased P solubilisation in grassland
612 compared to arable cropland. Near-stream attenuation of nitrate-N, likely through
613 denitrification and dilution, and enrichment in soluble P, through soil-groundwater
614 interactions, were similar in both catchments. This study highlighted the potential risk of
615 pollution swapping between dominant nitrate-N transfer and dominant soluble P transfer in a
616 context of land use change in groundwater fed catchments. Further research will be required
617 to better quantify the risk of increased nitrate-N and soluble P transfer when switching from
618 arable land to grassland, or the other way around, and to take account of interactions with
619 soil chemistry and climate in similar catchments.

620 **Acknowledgement**

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