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1 **Albedo and LAI estimates from FORMOSAT-2 data for crop** 2 **monitoring**

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12 13 **ABSTRACT**

14
15 This paper aimed at estimating albedo and Leaf Area Index (LAI) from FORMOSAT-2 satellite that
16 offers a unique source of high spatial resolution (eight meters) images with a high revisit frequency
17 (one to three days). It mainly consisted of assessing the FORMOSAT-2 spectral and directional
18 configurations that are unusual, with a single off nadir viewing angle over four visible – near infra
19 red wavebands. Images were collected over an agricultural region located in South Eastern France,
20 with a three day frequency from the growing season to post-harvest. Simultaneously, numerous
21 ground based measurements were performed over various crops such as wheat, meadow, rice and
22 maize. Albedo and LAI were estimated using empirical approaches that have been widely used for
23 usual directional and spectral configurations (i.e. multidirectional or single nadir viewing angle over
24 visible – near infrared wavebands). Two methods devoted to albedo estimation were assessed,
25 based on stepwise multiple regression and neural network (NNT). Although both methods gave
26 satisfactory results, the NNT performed better (relative RMSE =3.5% versus 7.3%), especially for

27 low vegetation covers over dark or wet soils that corresponded to albedo values lower than 0.20.
28 Four approaches for LAI estimation were assessed. The first approach based on a stepwise multiple
29 regression over reflectances had the worst performance (relative RMSE = 65%), when compared to
30 the equally performing NDVI based heuristic relationship and reflectance based NNT approach
31 (relative RMSE \approx 34%). The NDVI based neural network approach had the best performance
32 (relative RMSE =27.5%), due to the combination of NDVI efficient normalization properties and
33 NNT flexibility. The high FORMOSAT-2 revisit frequency allowed next replicating the dynamics
34 of albedo and LAI, and detecting to some extents cultural practices like vegetation cuts. It also
35 allowed investigating possible relationships between albedo and LAI. The latter depicted specific
36 trends according to vegetation types, and were very similar when derived from ground based data,
37 remotely sensed observations or radiative transfer simulations. These relationships also depicted
38 large albedo variabilities for low LAI values, which confirmed that estimating one variable from the
39 other would yield poor performances for low vegetation cover with varying soil backgrounds.
40 Finally, this empirical study demonstrated, in the context of exhaustively describing the
41 spatiotemporal variability of surface properties, the potential synergy between 1) ground based
42 web-sensors that continuously monitor specific biophysical variables over few locations, and
43 2) high spatial resolution satellite with high revisit frequencies.

44

45 **1 Introduction**

46

47 Several applications may exploit remote sensing estimates of land surface variables. They include
48 agriculture, forestry or land management as well as hydrological and meteorological forecasts. Due
49 to the high level of spatial heterogeneity, the size of agricultural fields or natural vegetation patches,
50 and the large dynamic of vegetation, observations made frequently at high spatial resolution are
51 desirable. However, because of technological and economical constraints, only high spatial

52 resolution sensors (10-30 m) with low revisit frequency (15-30 days) or low spatial resolution (250-
53 1000 m) with high revisit frequency (between one and three days) have been available over the last
54 three decades. With the continuous technological advances that reduce mass and cost of sensors and
55 satellite platforms, new missions are planned or launched such as FORMOSAT-2 (Chern et al.,
56 2006), Rapid-eye (Scherer and Kriscke 2001), Venus (Dedieu et al., 2006) and Sentinel-2
57 (Martimort et al., 2007).

58

59 Among the several surface variables that are accessible from remote sensing observations, Leaf
60 Area Index (LAI) and albedo are key players involved in the main processes that drive soil-plant-
61 atmosphere exchanges and biomass accumulation, including light and rain interception, evaporation
62 and transpiration, as well as photosynthesis and respiration. These two variables are part of the
63 essential climate variables identified by the Global Climate Observing system (GCOS, 2006). LAI
64 is defined as half the developed leaf area per unit of horizontal ground area, where leaf area
65 includes both leaf faces (Chen and Black 1992). Albedo is defined as the fraction of solar irradiance
66 (diffuse and direct fractions) reflected by the surface in the upper hemisphere, and integrated over a
67 given broad spectral domain (Jacob and Oliosio 2005). For energy balance related applications, the
68 integration domain is $[0.3-3.0] \mu\text{m}$ to derive solar albedo, although this domain is often split for
69 meteorological modeling into shortwave ($[0.3-0.7] \mu\text{m}$) and infrared ($[0.7-3.0] \mu\text{m}$) sub-domains, to
70 derive visible and near-infrared albedos respectively. While surface albedo is partly influenced by
71 LAI, several other variables have to be included for its computation, in addition to illumination
72 conditions: canopy structure (leaf area density, vegetation height, clumping and orientation...), and
73 leaf and soil optical properties (Jacobs and van Pul 1990; Oliosio 1992; Weiss et al., 1999). For this
74 reason no simple and general relationship between albedo and LAI is a priori expected.

75

76 Estimating LAI from remote sensing may be achieved using empirical relationships or radiative
77 transfer model based algorithms, as recently reviewed by Baret and Buis (2008). Empirical

78 relationships may either use bidirectional reflectance measurements along with parametric inverse
79 models such as multiple linear regression (Eklundh et al., 2003) and neural networks (Baret et al.,
80 1995), or combine bands into vegetation indices used as input for parametric modeling (Baret and
81 Guyot, 1991; Walthall et al., 2004). Such approaches are limited by data dependence and therefore
82 have a limited extrapolation capacity. Radiative transfer model inversion is potentially a very
83 powerful approach (Goel, 1989, Weiss et al., 2000; Meroni et al., 2004; Schlerf and Atzberger,
84 2006; Darvishzadeh et al., 2008), able to explicitly account for available information such as known
85 peculiarities of targeted canopies thanks to a priori knowledge. However, radiative transfer model
86 inversion is often limited by the realism of canopies structure description, and is known to be
87 severely ill-posed, which induces equifinality problems (Combal et al., 2002; Durbha et al., 2007).

88

89 Albedo can also be estimated from remote sensing by inverting radiative transfer models, but the
90 difficulties are the same than those faced when retrieving LAI. Another possibility is using
91 statistical relationships to perform a spectral extrapolation, so called Narrowband-To-Broadband
92 (NTB) conversion. NTB conversion can be applied to nadir bidirectional reflectance (Brest and
93 Goward, 1987; Russell et al., 1997; Liang et al., 1999; Song and Gao 1999; Liang, 2001; 2003;
94 Liang et al., 2005; Susaki et al., 2007) or, more adequately, on hemispherical reflectance (Wanner
95 et al., 1997; Weiss et al., 1999; Jacob et al., 2002a; van Leeuwen and Roujan, 2002; Greuell and
96 Oerlemans, 2004), where the latter is derived from Bidirectional Reflectance Distribution Function
97 (BDRF) parametric modeling along with directional integration over the upper hemisphere (Lucht,
98 1998; Lucht and Roujean, 2000; Jacob et al. 2002b; Pokrovsky and Roujean, 2003). Although
99 proved efficient when the directional dimension is sufficiently sampled, BRDF modeling can not be
100 applied to sensors such as FORMOSAT-2 or Venùs because the latter have single viewing
101 configurations. Besides, applying NTB conversion to FORMOSAT-2 or Sentinel data is not
102 straightforward regarding limited spectral samplings but especially single off nadir viewing angles.
103 Indeed, NTB conversion has been widely used for multidirectional observations or single nadir

104 viewing directions as abovementioned, but never for single off nadir viewing angles.

105

106 In the context of the high spatial resolution and revisit frequency offered by the FORMOSAT-2
107 sensor, the objective of this study was to explore its potentialities for the retrieval of solar albedo
108 and LAI, regarding the specificities of its spectral and directional configurations. Rather than
109 elaborating complex approaches based on radiative transfer model inversion, we concentrated on
110 empirical approaches based on continuous ground measurements over few representative fields.
111 Remote sensing and ground based datasets were collected over the Crau-Camargue experimental
112 site. The latter embraced a wide range of vegetation and soil moisture conditions, from dry wheat
113 and bare soil fields to irrigated meadow and flooded rice plantations. This allowed investigating the
114 robustness of the empirical methods considered for albedo and LAI retrieval. Furthermore, the high
115 revisit frequency of the collected FORMOSAT-2 dataset provided unique and detailed information
116 about evolution of land surface variables, over a large period ranging from the growing season to
117 post-harvest of major crops. From these chronicles of spatialized measurements, it was therefore
118 interesting to characterize the dynamics of albedo and LAI over different crop cycles, to analyze
119 possible links between these two variables, and to detect possible characteristic trends for different
120 vegetation covers according to anthropogenic forcing such as agricultural practices.

121

122 The paper is structured as follows. Section 2 describes the experiment and the resulting data set.
123 Section 3 presents the empirical methods developed to retrieve albedo and LAI from FORMOSAT-
124 2 data. Section 4 provides the performances of albedo and LAI retrieval methods, and discusses the
125 temporal dynamics of both variables for the several crops we considered. The best performing
126 methods for albedo and LAI retrieval are next selected and applied to the whole image set. This
127 allows analyzing the possible relationships between albedo and LAI for several crops throughout
128 their cultural cycles. Finally, conclusions are drawn with due attention to potential applications and
129 need for further investigations.

130

131 **2 Materials**

132 **2.1 The Crau-Camargue experimental site**

133

134 The study area was located in the “Crau-Camargue” area, South Eastern France (43.53°N; 4.66°E;
135 3 m above sea level, see Figure 1). Climate was typically Mediterranean, with irregular
136 precipitations, long dry periods in spring and summer, and strong winds. This flat area was
137 characterized by a wide range of soil and irrigation practices. The experiment took place in 2006,
138 including intensive ground measurements simultaneously collected with satellite data on various
139 crop types (Courault et al., 2008). Low cumulative precipitation was observed in 2006 (456 mm) as
140 compared to the regional average (548 mm in 2005). The weather was especially dry from April 1st
141 to mid-September, with one rainfall event only occurring in early June (28 mm).

142

143

[Figure 1]

144

145 The land cover was classified using a maximum likelihood supervised classification. The latter
146 relied on 1/ the four FORMOSAT-2 wavebands, 2/ five images distributed along the experimental
147 period, and 3/ reference areas of known field occupation within the study area. Image number and
148 repartition was driven by the temporal dynamics of vegetation cover, especially that resulting from
149 anthropogenic forcing such as meadow cuts throughout the cultural cycle (three to four cuts per
150 season, leaving 10 cm vegetation height). Twelve classes were identified, which included the main
151 vegetation covers, free water and bare soil surfaces. Although this map was improved with ground
152 based information, it was undeniably subject to uncertainties, with a 20% training residual error.
153 Figure 1 displays the major vegetation classes that represented about 40% of the whole study area.

154 The remaining 60% included natural marshlands, the Rhone River and other minor cultures.

155

156 **2.2 The five sampled fields**

157

158 Five fields that corresponded to the major crops within the region of interest (Figure 1) were chosen
159 for intensive ground based measurements. The two wheat fields (#1 and #2) were sown on
160 November 11th and December 15th, and harvested on June 27th and July 4th, respectively. They were
161 not irrigated, and turned to bare soils after harvest around DOY 179 (28th of June). The meadow
162 field (#3) was flooded every 11 days according to a regional water management designed for high
163 yields and quality. Three cuts were performed during the growing season, on May 5th, July 7th, and
164 August 11th. The maize field (#4) was sown on May 5th, was intermittently irrigated by sprinklers
165 depending on weather conditions, and was finally harvested on August 8th. This field was located on
166 poor soils that were very stony at some locations, which induced large heterogeneities in vegetation
167 cover that affected the representativeness of ground based measurements. The rice field (#5) was
168 sown on dry soil on April 27th, then continuously submerged from May 5th till October 6th with a
169 0.10 ± 0.05 m water height, and finally harvested on October 18th. Due to stem lodging within the
170 rice field after August 30th caused by strong winds, LAI measurements and FORMOSAT-2
171 reflectances were biased. Therefore, the data collected afterwards were discarded from analysis.

172

173 **2.3 Ground based measurements**

174

175 Albedo was measured with Kipp & Zonen CM7 sensors mounted between 1.5 m and 2 m above top
176 of canopy. Measurements were averaged over 10 minute periods throughout vegetation cycles. The
177 measurement footprints were circular, with radii between 25 and 35 m. Albedometers were
178 calibrated to measure incoming radiation over the whole solar spectrum (300 to 3000 nm).

179

180 Leaf Area Index (LAI) was derived from hemispherical images that provided Effective LAI. The
181 latter was closer to remote sensing estimates than the true LAI because of leaf clumping (Weiss
182 et al., 2004). The hemispherical images were collected with time intervals of about 10 days, for
183 capturing canopy structure dynamics. In order to represent field average LAI, a cross-pattern
184 sampling protocol was adopted. It consisted of 50 hemispherical images acquired within each study
185 field at each date of measurement. Table 1 gives the main characteristics of the ground
186 measurements performed within the different above mentioned fields.

187

188

[Table 1]

189

190 The CAN-EYE software (http://www.avignon.inra.fr/can_eye/page5.php) was used to process the
191 hemispherical images. CAN-EYE allowed computing the gap fraction from a series of RGB color
192 images through a simple interactive supervised classification process. LAI was then derived from
193 the resulting gap fractions, using look-up-table techniques based on the Poisson model (Nilson,
194 1971). A strong correlation ($R^2=0.98$) was observed when comparing those estimates against
195 planimetry based destructive LAI measurements that were simultaneously collected over few
196 locations. However a systematic underestimation of LAI due to leaf clumping was observed, with a
197 clumping coefficient of 0.68. This was consistent with the study from Demarez et al. (2008) who
198 reported a value of 0.71 over wheat, maize and sunflower.

199

200 **2.4 FORMOSAT-2 Data**

201

202 FORMOSAT-2 is a high spatial resolution satellite that collects images with an 8 m nadir spatial
203 resolution over a 24 km swath, in four 90 nm width wavebands centered at 488, 555, 650 and

204 830 nm. The orbital cycle is completed within one day. The sensor may deviate from nadir in order
205 to point at sites close to the ground track. Therefore, accessible locations at Earth's surface are
206 observed under a unique viewing direction. In our case, the Crau-Camargue site was targeted with
207 zenith (relative to nadir) and azimuth (relative to north) viewing angles around 41° and 239° ,
208 respectively. Images were collected every three to four days at 10:30 UTC from March to October
209 2006. They were processed for geolocation, radiometric calibration and atmospheric perturbations
210 following Hagolle et al. (2008). Clouds and related shadows were discarded following Baillarin et
211 al. (2004). Over the 36 images collected between March and October, 30 images were cloudless,
212 with a temporal gap spanning from April 17th to May 14th because of cloudy conditions.

213

214 **2.5 Matching ground measurements with FORMOSAT-2 data**

215

216 To consistently calibrate and validate the empirical methods we considered; it was necessary
217 performing spatial and temporal matching between ground based and FORMOSAT-2 data.

218

219 For albedo, the four FORMOSAT-2 pixels included in each albedometer footprint were extracted
220 and averaged. Albedo values acquired at 10:30 UTC were selected for comparisons, since they
221 matched satellite overpasses. According to the starting and ending dates of data collection that
222 varied from one field to another, the resulting dataset included 130 ground samples. Table 1 shows
223 very low values for the albedo coefficient of variation (CV, equal to the ratio of standard deviation
224 to mean value) derived from FORMOSAT-2 retrievals (retrieving method explained in Section 3)
225 over the albedometer footprints. Albedo could therefore be considered quite homogeneous, even for
226 the maize field that depicted an albedo CV twice larger than those depicted by the other fields.

227

228 For LAI, ground based measurements and FORMOSAT-2 pixels were collected or selected in order

229 to consider representative values at the field scale. Ground samples were collected within each field
 230 according to a cross-pattern protocol, and next averaged. For each field, FORMOSAT-2 overlaying
 231 pixels were selected by excluding borders, and the corresponding waveband reflectances were
 232 averaged. Table 1 shows the LAI CV over the field extensions, derived from FORMOSAT-2
 233 retrievals on a pixel basis (retrieving methods explained in Section 3). The low LAI CV values for
 234 the wheat, meadow and rice fields confirmed that comparing ground based and satellite data was
 235 consistent. However, the LAI CV for the maize field was significantly larger than those for the
 236 other fields, due to large heterogeneities for soil properties (Section 2.1.2). Calibration and
 237 validation results for this field were therefore carefully analyzed (Section 4 and 5). Overall, because
 238 of low heterogeneities for four fields over the fives, we expected calibration and validation results
 239 would not be affected by non linearity between LAI and reflectance values (Garrigues et al., 2007).

240

241 The quite low 10 day frequency of LAI ground based measurements induced a temporal
 242 interpolation was necessary, in order to obtain concurrent ground and satellite LAI estimates. For
 243 this purpose, the LAI dynamic model proposed by Koetz et al. (2005) was applied:

$$244 \quad LAI = k \cdot \left[\frac{1}{1 + e^{-b(T - T_i)}} - e^{-a(T - T_s)} \right] \quad (1)$$

245 T is the cumulated daily mean air temperature above vegetation zero, starting from the sowing date.
 246 The growth period is defined by a logistic equation which parameter b is the relative growth rate at
 247 the inflexion point T_i . The senescence is determined by an exponential equation which parameter a
 248 is the relative growth rate at the cumulated temperature T_s when all leaves are senescent. The k
 249 parameter is the maximal leaf area. These parameters were estimated using the simplex iterative
 250 optimization method (Nelder and Mead, 1965), with a 5% residual calibration error. For irrigated
 251 meadow, this empirical model did not allow representing the cuts. Therefore, a simple linear
 252 interpolation was applied, benefiting from more frequent ground based data on this field. Figure 9
 253 (later in the result section) illustrates LAI chronicles for the five fields. According to the starting
 254 and ending dates of data collection that varied from one field to another, the resulting dataset

255 included 72 ground based estimates that matched FORMOSAT-2 retrievals, completed with 25 data
256 for which the background was bare for sure, corresponding to LAI=0.

257

258 A summary on the available albedo and LAI ground measurement datasets is given in Table 2. It is
259 shown albedo values were within the range of usual values reported in the literature, for both bare
260 soil and vegetation cover conditions. Values for LAI were up to almost seven, which yielded us
261 expecting saturation problems when retrieving LAI from FORMOSAT-2 reflectances, especially
262 for these specific situations of large vegetation cover.

263

264 [Table 2]

265

266 **3 Methods**

267

268 When choosing the methods to be implemented for the retrieval of albedo and LAI from
269 FORMOSAT-2 data, we concentrated on empirical approaches based on continuous ground based
270 measurements over few representative fields. The motivations for choosing empirical approaches
271 were multiple. In the context of exhaustively describing the spatiotemporal variability of surface
272 properties, the FORMOSAT-2 spatial and temporal configurations allowed assessing the potentials
273 of synergy between 1/ ground based web-sensors that continuously monitor specific biophysical
274 variables over few locations, and 2/ high revisit frequency and high spatial resolution satellite
275 images. Second, although using deterministic approaches has more portability, it required first
276 inversion strategies that face the ill-posed problem and related equifinality troubles. Finally, the
277 FORMOSAT-2 spectral and especially directional configurations allowed enlarging the assessment
278 of empirical approaches that have been widely used for usual configurations (i.e. nadir viewing).

279

280 For albedo, the considered empirical approaches were the Narrowband-To-Broadband (NTB)
 281 conversion and the NNT based method. NTB conversion has been widely used for recovering
 282 albedo. NNT based method has been extensively employed for the retrieval of biophysical variables
 283 (fraction cover, chlorophyll content...) as reviewed by Baret and Buis (2008), and was therefore
 284 considered as a candidate method for albedo retrieval. When dealing with LAI, we considered
 285 multiple linear regressions similar to NTB conversion and recently proposed by Eklundh et al.
 286 (2003) for application over forests. We also assessed NNT based methods that have been widely
 287 used for retrieving LAI. In this last case, we considered to ways for applying NNT, either from
 288 reflectances to LAI, or from NDVI to LAI given NDVI has efficient normalization properties.
 289

290 **3.1 Stepwise multiple regression method**

291

292 For a given sun direction Ω_s , albedo $a(\Omega_s)$ over a considered spectral range can be approximated as
 293 the weighed summation of hemispherical reflectances $\rho_j^h(\Omega_s)$ (Jacob and Olioso, 2005):

$$294 \quad a(\Omega_s) = \gamma_0(\Omega_s) + \sum_{j=1}^n \gamma_j(\Omega_s) \cdot \rho_j^h(\Omega_s) \quad (2)$$

295 A specific spectral band amongst n is labeled j . The weighting coefficients $\gamma_j(\Omega_s)$ may vary as a
 296 function of Ω_s , and have to be adjusted through stepwise multiple regressions. As explained in
 297 Introduction, hemispherical reflectances $\rho_j^h(\Omega_s)$ may be derived from multidirectional observations
 298 through the parametric modeling of BRDF. When observations are available in a single viewing
 299 direction only, several studies proposed directly applying Equation 2 on bidirectional
 300 reflectances $\rho_j(\Omega_s, \Omega_o)$, where the latter are collected with close nadir viewing:

$$301 \quad \alpha(\Omega_s) = \beta_0(\Omega_s) + \sum_{j=1}^n \beta_j(\Omega_s) \cdot \rho_j(\Omega_s, \Omega_o) \quad (3)$$

302 This relies on assuming hemispherical and nadir bidirectional reflectances are linearly related,
303 either because spatial variabilities do not significantly affect the ratio of nadir to hemispherical
304 reflected radiances, or because it is possible considering a constant BRDF shape that is scaled to the
305 observed bidirectional reflectance. This strong assumption was directly validated over a limited
306 range of environmental conditions only (Russell et al., 1997; Weiss et al. 2002a; Schaaf et al. 2002).
307 However, indirect validations were performed over a large range of environmental conditions,
308 through the comparisons of albedo products derived from Equation 3 against reference estimates
309 (Brest and Goward, 1987; Russell et al., 1997; Liang et al., 1999; Song and Gao 1999; Liang, 2001;
310 Liang et al., 2002; Liang, 2003; Liang et al., 2005; Susaki et al., 2007). For the current study, the
311 additional difficulty was using off nadir bidirectional reflectance (zenith and azimuth angles of 41°
312 and 239° , respectively), which yielded foreseeing NTB conversion may perform poorly.

313

314 The stepwise multiple linear regression method has been applied by several authors for different
315 spectral and angular configurations. Given no coefficient set were available for the FORMOSAT
316 specific configuration over the Crau-Camargue study site, we assessed the performances of a
317 coefficient set calibrated over the collected dataset, and compared against existing coefficient sets
318 proposed by Weiss et al. (1999), Liang et al. (1999), and Jacob et al. (2002a). These coefficients
319 sets were obtained by considering different spectral configurations (waveband locations and
320 widths), different directional configurations (using Equation 2 and 3 for hemispherical or nadir
321 reflectances respectively), and by considering or not the diffuse component of solar irradiance.

322

323 LAI was derived from multiple linear regression applied over bidirectional reflectances by using the
324 same formalism as Equation 3. As compared to Eklundh et al. (2003) who validated the concept
325 with Landsat data over coniferous and deciduous forest canopies, assessing this approach was
326 interesting with regards to the large differences in experimental conditions, whether it was the
327 considered biomes (agricultural lands versus forests), the directional configuration (off nadir versus

328 nadir viewing), or the spectral configuration (no shortwave infrared band with FORMOSAT-2).

329

330 For both LAI and albedo retrievals, we suspected a significant limitation when using the coefficient
331 set calibrated over the whole Crau-Camargue dataset, with regards to changes in solar direction at
332 the time of satellite overpass. Indeed, solar zenith (respectively azimuth) angle varied throughout
333 the experimental period from 25° to 45° (respectively from 135° to 160°). Equation 3 was therefore
334 adjusted for each individual date, using the five pairs of matching ground albedo measurements /
335 FORMOSAT-2 bidirectional reflectances. Results displayed in Figure 2 showed that coefficients
336 $\beta_j(\Omega_s)$ devoted to albedo did not exhibit specific features as a function of solar direction, indicating
337 these coefficients could be assumed independent on Ω_s . The same independency was observed for
338 the coefficients devoted to LAI (results not shown for sake of brevity). Therefore, β coefficients
339 were also assumed independent on Ω_s for LAI in Equation 3.

340

341 [Figure 2]

342

343 **3.2 Neural network based method (NNT)**

344

345 Neural networks enabled relating the FORMOSAT-2 Green, Red and NIR waveband reflectances to
346 either ground based albedo values or interpolated LAI (labeled NNT_{Ref}). The Blue waveband was
347 omitted because it considerably degraded the performances of both albedo and LAI estimations
348 (results not reported here). Moreover, Jiang et al. (2008) demonstrated that, for a large range of
349 vegetation conditions at the global scale, the inclusion of the Blue band does not significantly
350 improve the vegetation characterization. We used the feed-forward back-propagation algorithm
351 detailed in Hagan and Menhaj (1994). It was made of made of a single hidden layer with two
352 tangent-sigmoid neurons, and one output layer with a single linear neuron (Figure 3). Prior to

353 training, inputs and outputs were normalized by their minimum and maximum values. The learning
 354 process was achieved using the Levenberg-Marquardt back-propagation method. Twenty random
 355 initializations were tested and the one providing the best performance was selected. Implementing
 356 the neural network required 11 coefficients to be tuned over the training data set (eight weights and
 357 three biases, Figure 3a). Hyper specialization was not evaluated due to the lack of independent data.
 358 It was however reduced by the minimal architecture selected for the network.

359

360

[Figure 3]

361

362 Neural networks were also used for relating NDVI to interpolated LAI (labeled NNT_{NDVI}). In this
 363 case, the implementation of the neural network required seven coefficients only to be tuned over the
 364 training data set (four weights and three biases, Figure 3b). Indeed, this implementation was based
 365 on a unique NDVI independent variable, whereas the previous NNT design was based on three
 366 independent variables as inputs, i.e. the Green, Red and NIR reflectances (Figure 3a).

367

368 **3.3 Exponential law based method**

369

370 Among the various relationships between NDVI and LAI proposed in the literature, we considered
 371 for the current study the exponential law derived from the studies of Asrar et al. (1984), Baret and
 372 Guyot (1991) and Wilson and Meyer (2007), and which has been widely used.

$$373 \quad LAI = -\left(\frac{1}{K_{LAI}}\right) \ln\left(\frac{NDVI - NDVI_{\infty}}{NDVI_s - NDVI_{\infty}}\right) \quad (4)$$

374 $NDVI_{\infty}$ is the asymptotic value of NDVI when LAI tends towards a maximum value, $NDVI_s$ is the
 375 bare soil NDVI value and K_{LAI} is an extinction coefficient. The simplex optimization approach was
 376 used to adjust parameters $NDVI_{\infty}$, $NDVI_s$, and K_{LAI} by minimizing the Root Mean Square Error
 377 (RMSE) between measured and estimated LAI.

379 **3.4 Calibration and validation procedures and performance metrics**

380

381 The experimental data set (sample number N=130 for albedo and 97 for LAI) was not large enough
 382 to be split into independent calibration and validation datasets. Therefore, a “leave-one-out” cross-
 383 validation method (Stone 1974; Geisser 1975) was used for validation. It consisted in calibrating
 384 over n-1 data, and validating over the remaining “left-out” data. This process was repeated N times
 385 to cover the whole dataset. Then, performances were assessed using standards metrics:

- 386 • Absolute Mean Error (ME_A), was the bias between measured (M_i) and estimated (E_i) values,

$$387 \quad ME_A = \frac{1}{n} \sum_{i=1}^n (E_i - M_i) \quad (5)$$

- 388 • Absolute Root Mean Square Error ($RMSE_A$) quantified the scatter between measured and
 389 estimated values, and Relative Root Mean Square error ($RMSE_R$) was the ratio of $RMSE_A$ to
 390 the mean of measured values $\langle M_i \rangle$.

$$391 \quad RMSE_A = \sqrt{\frac{1}{n} \sum_{i=1}^n (E_i - M_i)^2} \quad RMSE_R = \frac{RMSE_A}{\langle M_i \rangle} \quad (6)$$

392

393 **3.5 LAI versus Albedo**

394

395 The large dataset of albedo and LAI, that included spatialized estimates spreading over entire crop
 396 cycles, allowed analyzing a possible relationship between these two biophysical variables. To be
 397 able to compare and understand such a possible relationship, radiative transfer simulations were
 398 carried out for conditions similar to those prevailing during the experiment. LAI and albedo values
 399 were simulated using the canopy reflectance model PROSAIL (Jacquemoud et al. 2008) for three
 400 types of ground surfaces (water, dry soil and wet soil), and a range of leaf inclination angles (20°,
 401 58° and 70°) under a constant zenith angle of 30°. These values corresponded to averaged values

402 that were representative of the experimental conditions. Diffuse fraction and spectral irradiance
403 were simulated using the 6S atmosphere radiative transfer model (Vermote et al., 1997) with mid-
404 latitude summer atmosphere and clear conditions (aerosol optical thickness at 550 nm set to
405 0.2347). All the other variables used for the simulations corresponded to typical average values
406 representative of soil and vegetation conditions similar to the Crau-Camargue conditions (Table 3).

407

408 [Table 3]

409

410 **4 Results and discussion**

411

412 This section presents the results we obtained when retrieving albedo (§ 4.1) and LAI (§ 4.2), as well
413 as when analyzing a possible relationship between both (§ 4.3). For the retrieving of albedo and
414 LAI, results are reported by separating method performances (§ 4.1.1 for albedo, § 4.2.1 for
415 reflectance based LAI, § 4.2.2 for NDVI based LAI) and analysis of chronicles captured by both
416 ground based and FORMOSAT-2 observations (§ 4.1.2 for albedo and § 4.2.3 for LAI).

417

418 **4.1 Albedo estimates from FORMOSAT-2 data**

419 **4.1.1 Comparison of the different method performances**

420

421 Table 4 displays the results we obtained for the retrieval of albedo, when calibrating NTB
422 conversion and NNT method over the Crau-Camargue dataset (Set 1 and NNT), and when
423 validating coefficient sets proposed by previous studies for other sensors with different spectral and
424 direction configurations. Set 2 and 3 were calibrated over a simulated dataset, and designed for
425 hemispherical reflectances collected within generic wavebands (Weiss et al., 1999). Set 4 was

426 calibrated over a measured dataset, and designed for hemispherical reflectances collected within
427 POLDER wavebands (Jacob et al., 2002a). Set 5 was calibrated over a measured dataset, and
428 designed for nadir bidirectional reflectances collected within MISR waveband (Liang et al., 1999).

429

430 [Table 4]

431

432 The stepwise multiple regression (Equation 3) calibrated over the Crau-Camargue dataset was
433 obtained by selecting positive and statistically significant bands only (Set 1). Best performances
434 were obtained when using the Red and NIR bands only, with a corresponding offset equal to 0.
435 Thus, absolute bias ME_A was almost negligible and Relative Root Mean Square Error $RMSE_R$ was
436 acceptable, around 7.5%. These validation results were comparable to calibration residual errors
437 reported by Weiss et al. (1999), Liang et al. (1999) and Jacob et al. (2002a), and were close to
438 relative accuracy of albedometer measurements and FORMOSAT-2 corrected data (around 5%).

439

440 Similarities in performances for Set 1 versus Set 2 and 4 were explained by similarities in
441 coefficient values, the latter varying of about 8 and 14% in relative for the red and near infrared
442 bands, respectively. When applying coefficients Set 3 and 5 that included the Green band;
443 performances were even worse as compared to Set 2 that included Red and NIR bands only. This
444 indicated the Green band could have added more noise than information in albedo estimation.

445

446 We could not discriminate performances according to the consideration of hemispherical or nadir
447 bidirectional reflectances, whereas the best performances were observed with the data set of
448 FORMOSAT-2 off nadir bidirectional reflectance. These both elements contributed to strengthen
449 the directional approximations formulated in Section 3.1, when assuming NTB conversion could
450 also be applied to off nadir bidirectional reflectances. However, this positive report ought to be
451 moderated because of additional complexities when comparing the different coefficient sets, such as

452 the combined effects between differences in spectral configurations and land surface properties.

453

454 When dealing with the coefficient Set 1 that was calibrated over the Crau-Camargue dataset, the
455 sum of the coefficients appeared to be almost equal to one (column $\Sigma\beta$ in Table 4). This was in
456 agreement with Brest and Goward (1987), and Jacob et al. (2002a). Indeed, the whole solar
457 spectrum could be split into nominal spectral intervals, and supposedly fully scanned through the
458 corresponding wavebands. Then, the associated weighting coefficients corresponded to the fractions
459 of solar irradiance over these intervals. However, this physical assumption might be far from reality
460 when characterizing the whole spectral domain by using a visible and a near infrared band only.
461 The Blue band was never significant in the considered coefficient sets: its weight might be
462 relatively small due to the low radiation level. Further, it might also be considerably disturbed by
463 residual atmospheric effects due to inappropriate aerosol corrections. Moreover, this band might not
464 provide additional information. Indeed, Jiang et al. (2008) reported strong correlations with the Red
465 band for a for a large range of vegetation conditions at the global scale.

466

467 Inspecting performances from one field to another showed that those were lower over rice and
468 meadow, with $RMSE_R$ of 10.5% and 10.1% respectively. Figure 4a emphasizes the difficulties for
469 the regression to fit the scattering induced by very different situations between meadow and rice,
470 with an overestimation (respectively underestimation) for rice albedo (respectively freshly cut
471 meadow albedo). These poor performances could be explained by the lack of water sensitive
472 shortwave infrared (SWIR) wavebands within the FORMOSAT-2 configuration. Indeed, inclusion
473 of such bands might improve albedo estimation under conditions of wet soil and free water
474 background. To our knowledge, very few studies tried to estimate rice albedo from remote sensing,
475 probably because of additional difficulties induced by the presence of water background. Very
476 recently, Susaki et al. (2007) obtained a $RMSE_R$ of 15.1 % with ASTER data over rice cultures in
477 Japan, when applying the appropriate coefficient set proposed by Liang et al. (2001). This error was

478 larger than that obtained here, indicating our calibrated coefficient set was acceptable, despite 1/ the
479 consideration of complex surfaces that combined water and vegetation, 2/ the absence of
480 FORMOSAT-2 SWIR waveband, and 3/ the consideration of off nadir bidirectional reflectances.

481

482 [Figure 4]

483

484 Table 4 also indicates the NNT method had the best performance. The improvement mainly
485 occurred for meadow ($RMSE_R$ decreasing from 10.3% to 4.3%), and in a lesser extent for rice and
486 maize (Figure 4b). Performance improvement systematically corresponded to albedo values lower
487 than 0.2, i.e. for crops with low amount of vegetation and wet or dark soil background. As
488 compared to the simple linear multiple regression, the NNT were more flexible, thanks to both their
489 non linear character and their larger degree of freedom (number of coefficients to be tuned, see last
490 column of Table 4). This could explain the better performances observed in complex situations,
491 with variable background properties and low amount of vegetation. Even if the obtained $RMSE_R$
492 was 3.5%, it could be set to 5%, which corresponded to the accuracy of both albedometer
493 measurements and FORMOSAT-2 corrected data. It is worth noting this NNT based empirical
494 approach has never been applied to albedo estimation, and such results are quite encouraging.
495 Further applications on other datasets would be necessary for additional validations.

496

497 **4.1.2 Dynamics of Albedo as estimated from FORMOSAT-2 data**

498

499 Figure 5 displays albedo dynamics captured throughout the study period from ground based
500 measurements and NNT based FORMOSAT-2 retrievals, when considering all fields apart from
501 meadow (Field #3). It confirms the close agreement previously observed on Figure 4b, since we
502 selected estimates from the best performing retrieval method. Dynamics of ground based albedo
503 values measured at 10:30 UTC showed large variations, mainly due to daily changes in diffuse

504 fraction of solar irradiance, and in a lesser extent to changes in surface properties. Conversely,
505 albedo estimates from FORMOSAT-2 depicted a reduced variability, because they corresponded to
506 data collection under clear sky conditions with a low diffuse component of solar irradiance.

507

508 Both wheat fields had comparable albedo dynamics, while bare soils exhibited contrasted time
509 courses after harvesting (DOY 179), and especially after DOY 220 (pointed by an arrow in bottom
510 left subplot of Figure 5). This was mainly due to green vegetation re-emergence in wheat Field #1.
511 Maize field dynamics showed limited amplitudes of variation, the few large changes being ascribed
512 to irrigation practices. Dynamics of albedo for rice showed a significant increase when the canopy
513 developed. This was in agreement with observations from Maruyama et al. (2007) over rice crops.
514 Indeed, they reported a first period with low albedo values (around 0.10) that corresponded to low
515 vegetation cover overlaying a water or very wet soil background, and followed by an albedo
516 increase during vegetation growth, until it stabilized around a value of 0.18 to 0.20.

517

518

[Figure 5]

519

520 Figure 6 displays albedo dynamics captured throughout the study period from ground based
521 measurements and NNT based FORMOSAT-2 retrievals, when considering Field #3 (meadow)
522 only. It is shown the good agreement between ground based observations and remotely sensed
523 estimates. The amplitude of variation was limited, with almost neither seasonal trend nor large
524 variations. This was mainly explained by irrigation or rainfall events, which induced a systematic
525 albedo drop (between 0.01 and 0.04) that vanished after two or three days. Note that the cuts did not
526 induce large albedo variation, probably because a significant fraction of green vegetation was kept.

527

528

[Figure 6]

529

530 **4.2 LAI estimates from FORMOSAT-2 data**

531 **4.2.1 LAI as a function of individual reflectances**

532

533 Table 5 displays the parameters we obtained when calibrating the various empirical methods to be
534 considered when retrieving LAI from FORMOSAT-2 data. Are also indicated the corresponding
535 performances in terms of absolute bias, absolute and relative root mean square errors. We recall
536 LAI references were obtained from ground based measurements through a temporal interpolation
537 (§2.1.5). Statistical analysis through stepwise regression retained the Red and NIR wavebands only.
538 As illustrated by Figure 7a that displays the “leave-one-out” cross-validation, the corresponding
539 performances were significantly poor. Further, the $RMSE_R$ values displayed in Table 5 were larger
540 than those obtained by Fassnacht et al. (1995), Eklundh et al. (2003) and Jensen and Binford (2004).

541

542 [Table 5]

543

544 [Figure 7]

545

546 As compared to the previous studies abovementioned, the lower performances we observed were
547 explained by differences in land surface properties and remotely sensed information. Indeed, these
548 former studies were devoted to the monitoring of forests from Landsat Thematic Mapper (TM), thus
549 benefiting from both lower spatial heterogeneities, and from additional spectral information through
550 shortwave infrared (SWIR) wavebands. For the current study, the significantly poor performances
551 were ascribed to several factors. The first one was the large variabilities of canopy structure and soil
552 background properties, to be both taken into account with three freedom degrees only (Table5).
553 Second, the absence of FORMOSAT-2 SWIR waveband may contribute to the poorer performances
554 in LAI estimation. Indeed, Eklundh et al. (2003) obtained a significant contribution of the SWIR

555 wavebands while estimating the LAI over a forest using the multiple regression approach.

556

557 Rice was the unique field for which albedo was reasonably well estimated, all the other fields
558 corresponding to large scattering between measured and estimated values. Although stepwise
559 multiple linear regression could have been independently applied over each of the four land cover
560 classes, the restricted data set would have prevented from obtaining robust relationships.
561 Additionally, the maize field was quite heterogeneous, and the ground sampling was probably too
562 small to obtain a representative value of the field LAI.

563

564 Training of neural networks over individual reflectances (NNT_{Ref}) showed significant
565 improvements of retrieval performances. The RMSE values were indeed twice lower (Table 5 first
566 line as compared to second one, and Figure 7c as compared to Figure 7a). This was ascribed to the
567 larger number of coefficients to be tuned (Table 5). Regardless of the considered field, the
568 scattering between estimated and measured LAI values (Figure 7c) was similar.

569

570 **4.2.2 LAI as a function of NDVI**

571

572 Table 5 and Figure 7b show that both the NDVI based heuristic formulation (exponential shape)
573 approach and the NDVI based neural network approach NNT_{NDVI} performed better than the
574 reflectances based approaches, whether it was NTB conversion or NNT.

575

576 A unique set of parameters for the heuristic formulation (Equation 4) was adjusted over the whole
577 set of (ground based LAI, FORMOSAT-2 NDVI) pairs. Value of $NDVI_{\infty}$ (Table 5) was very
578 comparable to those obtained by Weiss et al. (2002b) and Wilson and Meyers (2007) over similar
579 vegetation types but with different viewing angles (nadir looking with these former studies, against

580 41° zenith / 239° azimuth with the current study), and probably different solar directions. This may
581 explain why K_{LAI} value obtained in the current study (0.71) was slightly larger than that obtained
582 (0.67) by Weiss et al. (2002b). The NDVI normalization properties appeared to be efficient. With
583 three freedom degrees only, better performances were obtained than with the 11 freedom degree of
584 the reflectance based NNT method, the $RMSE_R$ dropping down by 20% in relative.

585

586 Training a neural network with measured LAI and NDVI values allows more flexibility in the shape
587 of the relationship between these two variables. Seven parameters had indeed to be tuned, as
588 compared to three parameters when using heuristic formulation (Equation 4). Results show this
589 method performed best, with a $RMSE_R$ value of 27.54% (Table 5 and Figure 7d). These good
590 performances were ascribed to the combined effect of NDVI efficient normalization properties and
591 the NNT flexibility. However, saturation problems still were observed for LAI values larger than 4,
592 a problem that may result from the saturation of the remotely sensed signal over the optical domain.

593

594 Figure 8 illustrates the LAI – NDVI relationship generated by both the NNT_{NDVI} and the heuristic
595 formulation (Equation 4). Both methods followed a smooth exponential trend in compliance with
596 Equation 4. The figure also shows how the heterogeneous, row-planted maize crop exhibited a
597 different behavior than the other homogeneous vegetation covers. This heterogeneity in structure
598 and field cover could have made the LAI – NDVI relationship quite different over this crop.
599 Consequently, the latter was not well characterized through the multi-crop calibrated relationship.

600

601

[Figure 8]

602

603 **4.2.3 Dynamics of LAI as estimated from FORMOSAT-2 data**

604

605 Dynamics of FORMOSAT-2 LAI retrieved from the best performing NNT_{NDVI} approach was very

606 smooth, as expected (Figure 9). Dynamics showed a classical temporal pattern for wheat, maize and
607 rice, with increase LAI during the vegetative growth, followed by a relatively rapid senescence.
608 Rice and maize crops depicted bare soil conditions (LAI=0) at the beginning. Wheat fields showed
609 significant amount of green vegetation after harvest (LAI \approx 0.4), presumably corresponding to weeds
610 that developed after the rainfall events observed around DOY 255. Dynamics of meadow shows the
611 three cuts and the re-emergence just afterwards. Note that since LAI was around one just after the
612 cuts, this explains why the albedo dynamics was not affected by these cuts.

613

614 [Figure 9]

615

616 **4.3 LAI versus Albedo**

617

618 NNT based FORMOSAT-2 retrievals of both LAI and albedo over the whole study area were
619 matched for all vegetation types and for all acquisition dates. Figure 10 displays the resulting
620 boxplots along with the corresponding statistics (mean value, lower and upper quartiles), and the
621 ground based measurements. It is shown when LAI increases, albedo follows almost the same trend
622 for wheat, meadow and maize. However, rice behaves differently, and especially for LAI values
623 lower than 2.5. The general trend observed over the whole area from FORMOSAT-2 retrievals was
624 similar to that observed over the five fields where measurements were collected. However, most of
625 the field data were outside the 95% confidence interval of albedo variation within each LAI class.
626 This indicated although ground based data were representative of each field as shown by the low
627 CV values in Table 1, the fields were not representative of the land use classes. Although this was
628 not expected, consequences on the empirical models calibrated for estimating LAI and albedo might
629 be marginal. Indeed, calibrations were performed by including all crops and dates, while LAI and
630 albedo measured values covered most of the ranges observed across the study area and period.

631

632

[Figure 10]

633

634 Figure 11 displays the LAI – albedo relationship simulated from radiative transfer modeling (§ 3.5).

635 For the investigated crops, these simulations agree very well with observations displayed on

636 Figure 10. The soil background played a major role for low to moderate LAI values ($LAI < 3$), which

637 could explain the differences observed between the four different crop types. For larger LAI values,

638 albedo tends towards an asymptotic value that may depend on canopy architecture (leaf inclination)

639 as well as on other variables not investigated here like chlorophyll content and mesophyll structure

640 parameter. The largest albedo variability within a LAI class, depicted in Figure 10, corresponded to

641 intermediate LAI values ($1 < LAI < 2.5$), where both soil background and vegetation influence the

642 reflected radiation. The albedo ranges were slightly lower for LAI lower than 1, because the main

643 influence was due to different soil surface reflectance properties. For LAI larger than 2.5, the albedo

644 range kept on decreasing with increasing LAI where the vegetation reflectance properties dominate.

645

646

[Figure 11]

647

648 **5 Conclusion**

649

650 This study demonstrated that it is possible accurately retrieving albedo and LAI from the specific

651 FORMOSAT-2 observations, along with empirical approaches based on restricted spatial sampling

652 but continuous monitoring. Although calibrations were performed by including all data collection

653 dates and all crops (wheat, maize, rice and meadow), very good performances were achieved.

654 Further, the advantages of the neural network techniques (NNT) over linear multiple regressions or

655 heuristic formulations was demonstrated. For the first time an NNT based method was used to

656 retrieve albedo, and the results were quite encouraging.

657

658 Due to the relatively small sample size, the empirical models established for retrieving LAI and
659 albedo probably need additional independent evaluation about their robustness, with emphasis on
660 the sampling strategy that optimizes the number, locations and dates of the ground based
661 measurements. When applied over orchards and vineyards, attention should also be paid to row
662 orientations relatively to solar or viewing directions. Furthermore, these relationships were
663 calibrated under the specific FORMOSAT-2 viewing conditions that are latitude / longitude
664 dependent (here, zenith angle = 41° and azimuth angle = 239°). Application to other conditions may
665 require adaptations, either by using radiative transfer models if well calibrated over the considered
666 surfaces, or by replicating the whole experimental process under these new conditions.

667

668 Alternative approaches based on radiative transfer model inversion were not considered in this
669 study, and should require further efforts. This might be possible under conditions of well defined
670 prior information, given single multi-date and multi-crop calibration of empirical approaches
671 yielded accurate estimates of LAI and albedo, and this in spite of limited information provided by
672 the FORMOSAT-2 samplings (three wave bands and a single off nadir viewing direction).

673

674 Another original output of this study was the possibility to investigate the spatial correlation
675 between albedo and LAI. We observed great consistencies when comparing albedo – LAI
676 relationships derived from ground based data, remotely sensed observations, or radiative transfer
677 simulations; with specific trends according to vegetation types. However, we observed large albedo
678 variabilities within most LAI classes, particularly for low leaf area indices. It appeared therefore
679 that estimating one variable from the other would yield poor performances, particularly for low LAI
680 values under varying soil background conditions. Conversely, albedo may be estimated to be
681 around 0.2 for LAI values larger than 4, at least for the canopy we considered.

682

683 This study also highlighted the interest of frequent observations at high spatial resolution for
684 vegetation monitoring. It allowed obtaining detailed features of the dynamics from which several
685 information could be derived; in relation with either brutal changes resulting from cultural practices
686 (cuts, irrigation under certain conditions) or more smooth evolutions resulting from canopy
687 phenology and functioning. In the context of exhaustively describing the spatiotemporal variability
688 of surface properties, this study finally demonstrated the potentials of the synergy between
689 1/ ground based web-sensors that continuously monitor specific biophysical variables over few
690 locations, and 2/ high revisit frequency and high spatial resolution satellite images.

691

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693

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701

702

703

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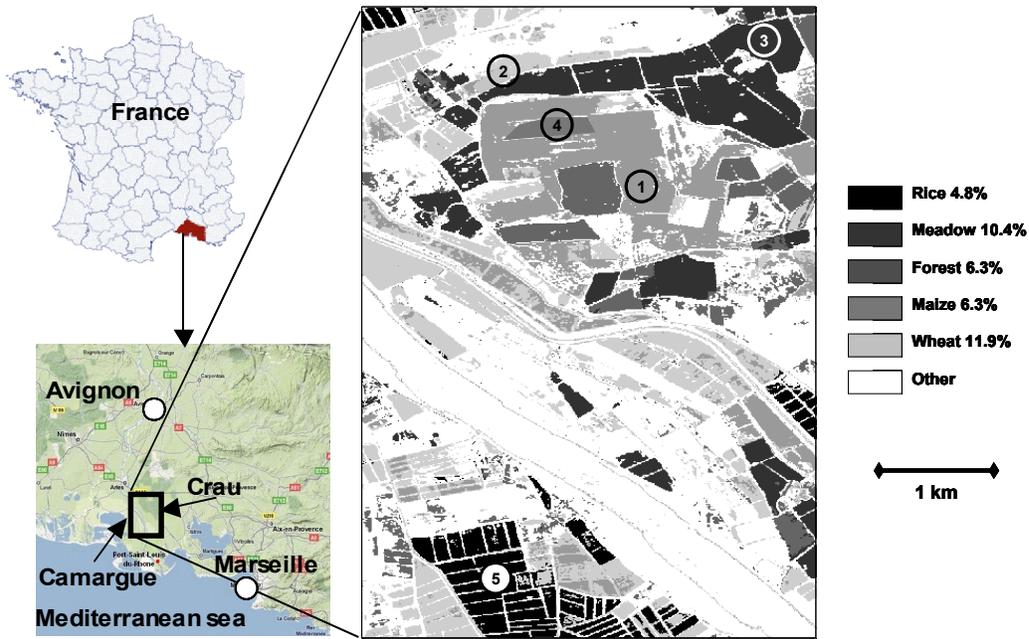
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886 **Figure 1: Land use map of the “Crau-Camargue” study area. The major cultivations are represented with their**
887 **occupation occurrence over the experimental area. Ground measurements took place from March to October**
888 **2006 over fields that were numbered #1 and #2 for wheat (turning to bare soils at the end of June), #3 for**
889 **meadow, #4 for maize and #5 for rice.**

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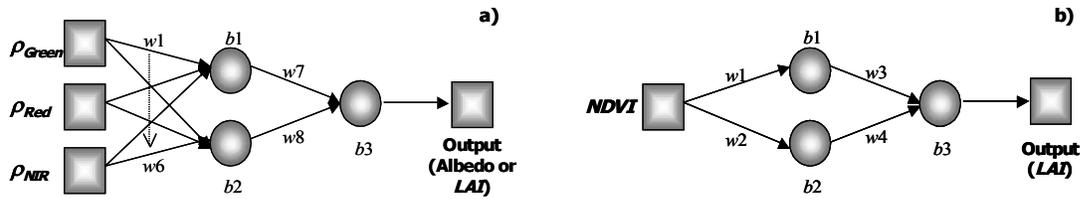
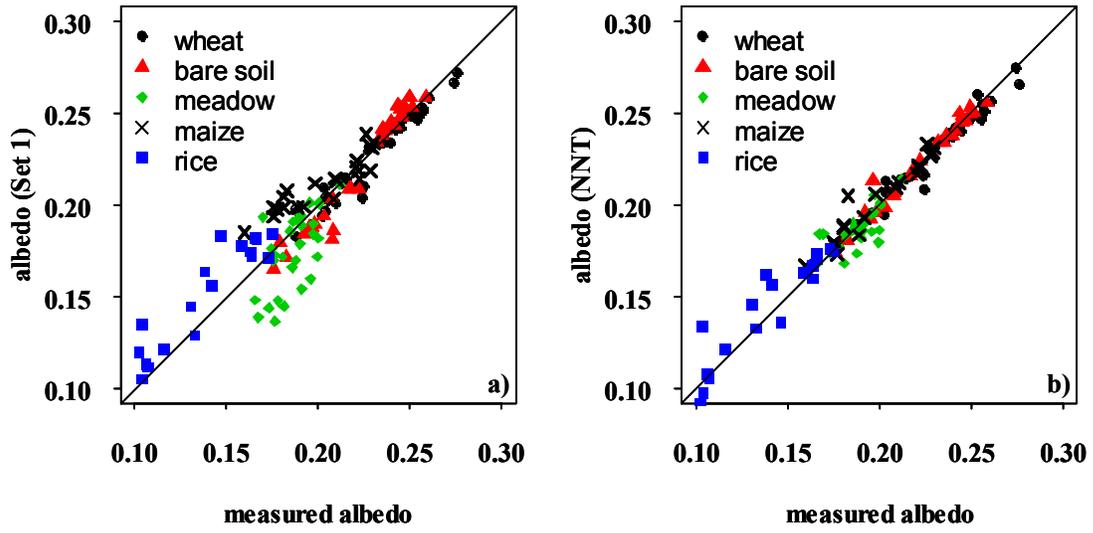


Figure 3: Neural network architecture when inputs were a) the green, red, and near infrared reflectances (ρ) and b) NDVI only. The squares represent the input and output variables, and the circles represent both the two tangent sigmoid neurons of the hidden layer and the single linear neuron of the output layer. The network structure required for a) 11 coefficients to be tuned: eight weights ($w1...w8$) and three biases ($b1...b3$) and for b) seven coefficients to be tuned: four weights ($w1...4$) and three biases ($b1...b3$).



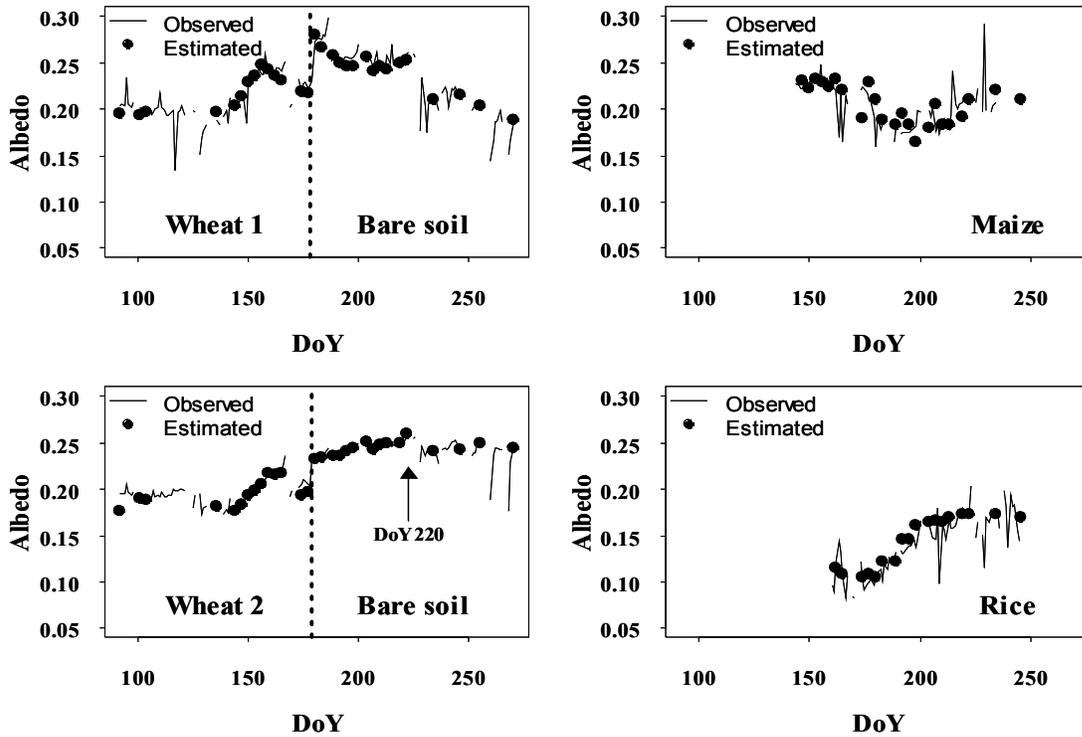
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912 **Figure 4: Comparison, over the four considered crops, between ground based measured albedo and**
 913 **FORMOSAT-2 retrievals based on (a) the coefficient Set 1 and (b) the NNT. Scatterplots correspond to the**
 914 **leave-one-out validation data set. The 5 fields are identified by different symbols.**

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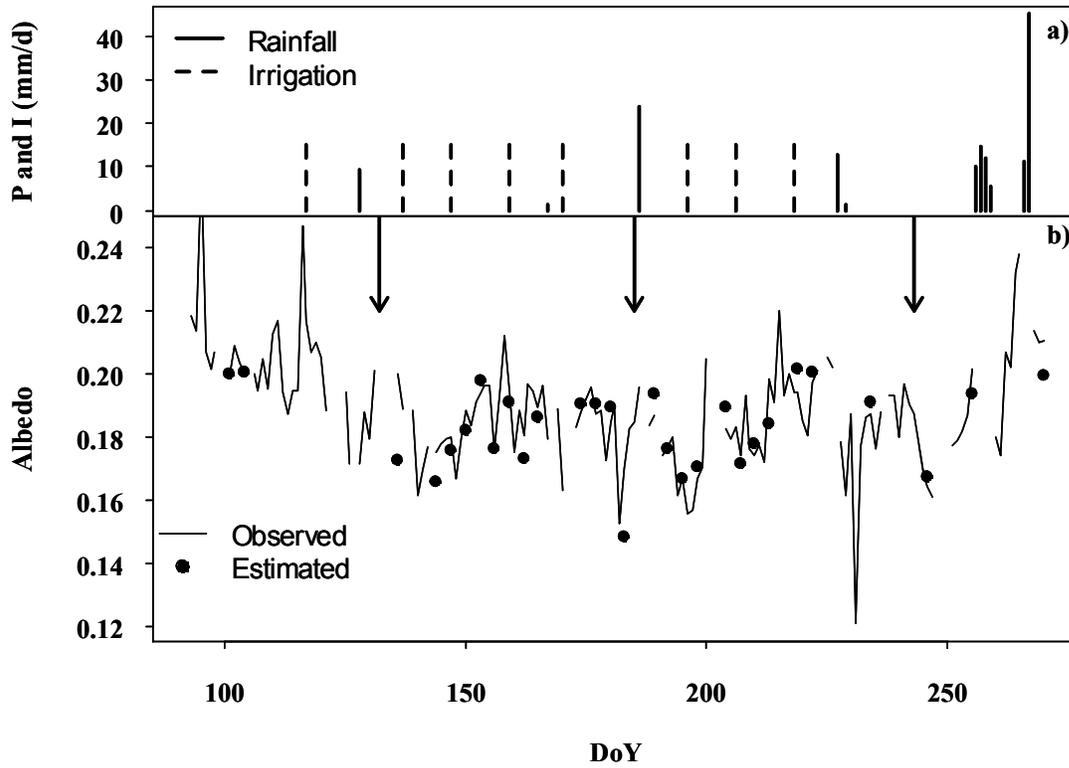
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920 **Figure 5: Albedo dynamics captured throughout the study period from ground based measurements (labeled**
 921 **observed) and FORMOSAT-2 retrievals using the neural network method (labeled estimated), when considering**
 922 **wheat / bare soil (Field #1 and 2), maize (Field #4) and rice (Field #5). Dotted vertical lines indicate the harvest**
 923 **dates. DOY 220 discussed in the text is pointed by an arrow on the Wheat 2 figure (bottom left subplot).**

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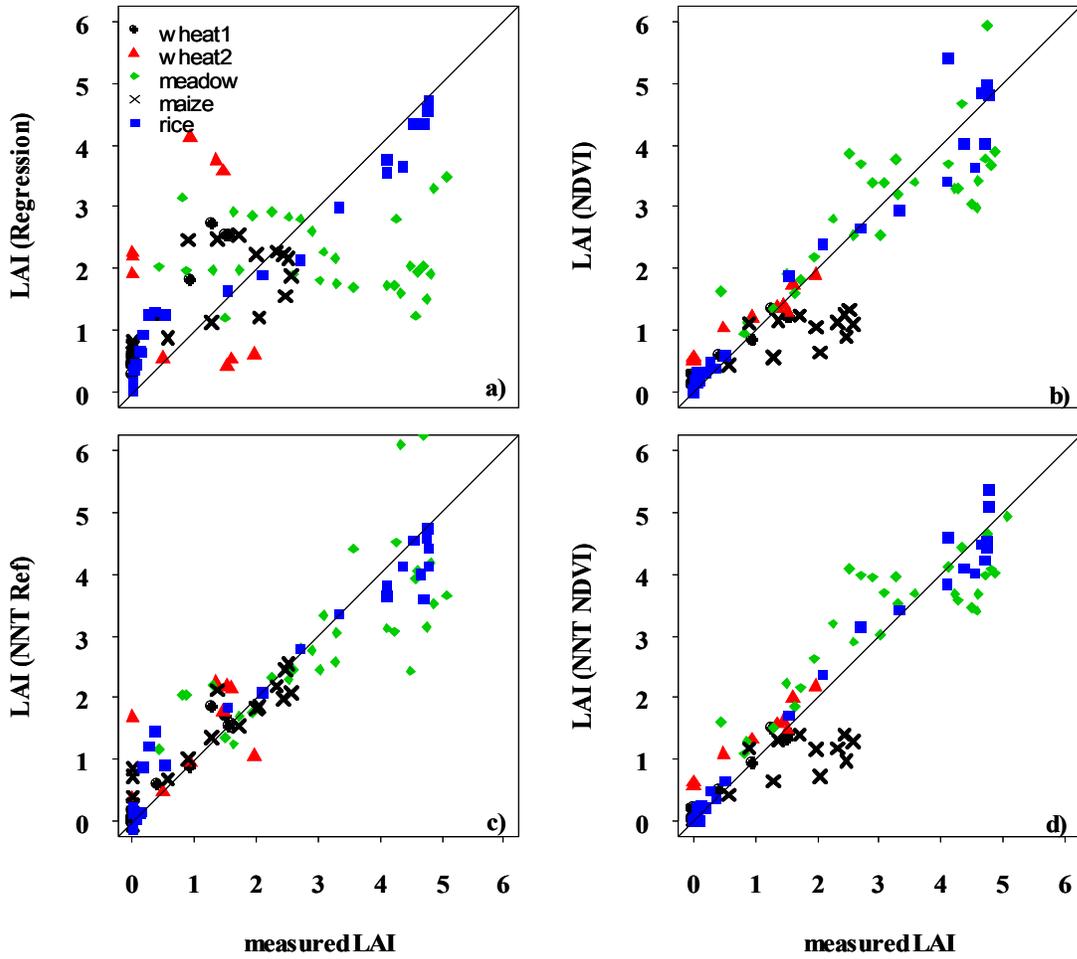
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929 **Figure 6: Albedo dynamics captured throughout the study period from ground based measurements (labeled**
 930 **observed) and FORMOSAT-2 retrievals using the neural network method (labeled estimated), when considering**
 931 **irrigated meadow (Field #3). For analysis, albedo dynamics (bottom) is plotted with rainfall and irrigation events**
 932 **(top). Meadow cuts are illustrated by vertical arrows.**

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938 **Figure 7: Leave-one-out cross-validation, over the four considered crops, of FORMOSAT-2 LAI retrievals**

939 **against ground based measurements, where the latter were temporally interpolated through Equation 1.**

940 **Estimates are derived from a) multiple regression, b) neural network technique (NNT) based method with**

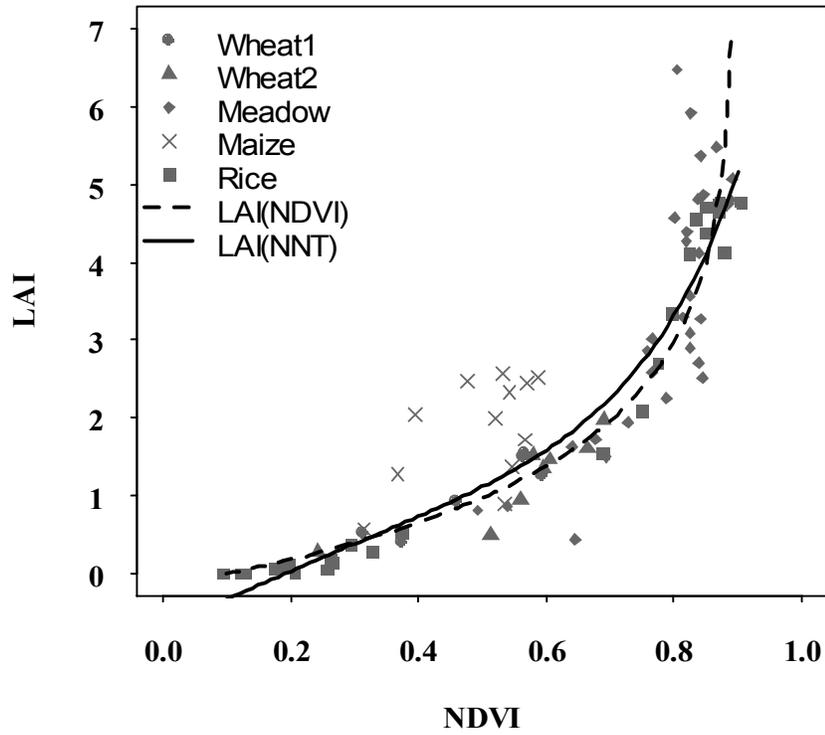
941 **reflectances as inputs, c) NDVI based heuristic formulation (Equation 4), and d) NNT based method with NDVI**

942 **as input.**

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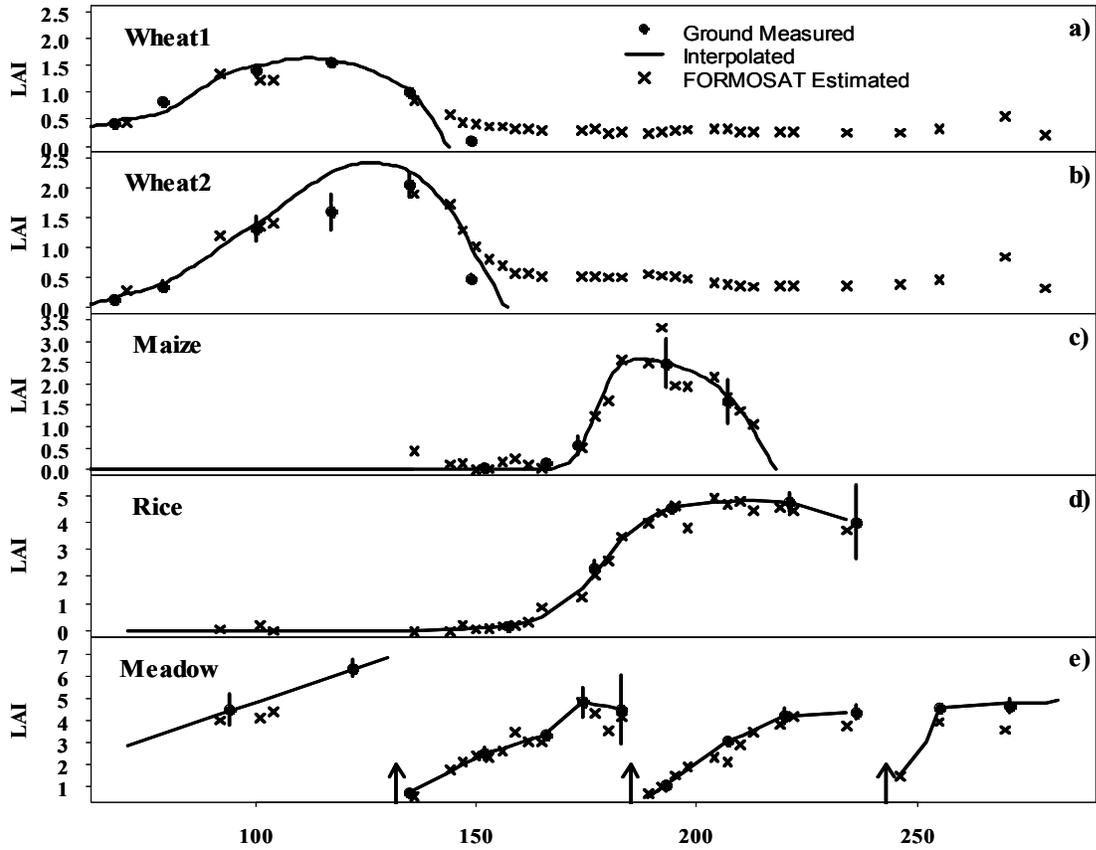
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948 **Figure 8: Comparison between the LAI versus NDVI relationships, as obtained by the NDVI exponential law**
949 **method (LAI_{NDVI}) or by the neural network method (LAI_{NNT}).**

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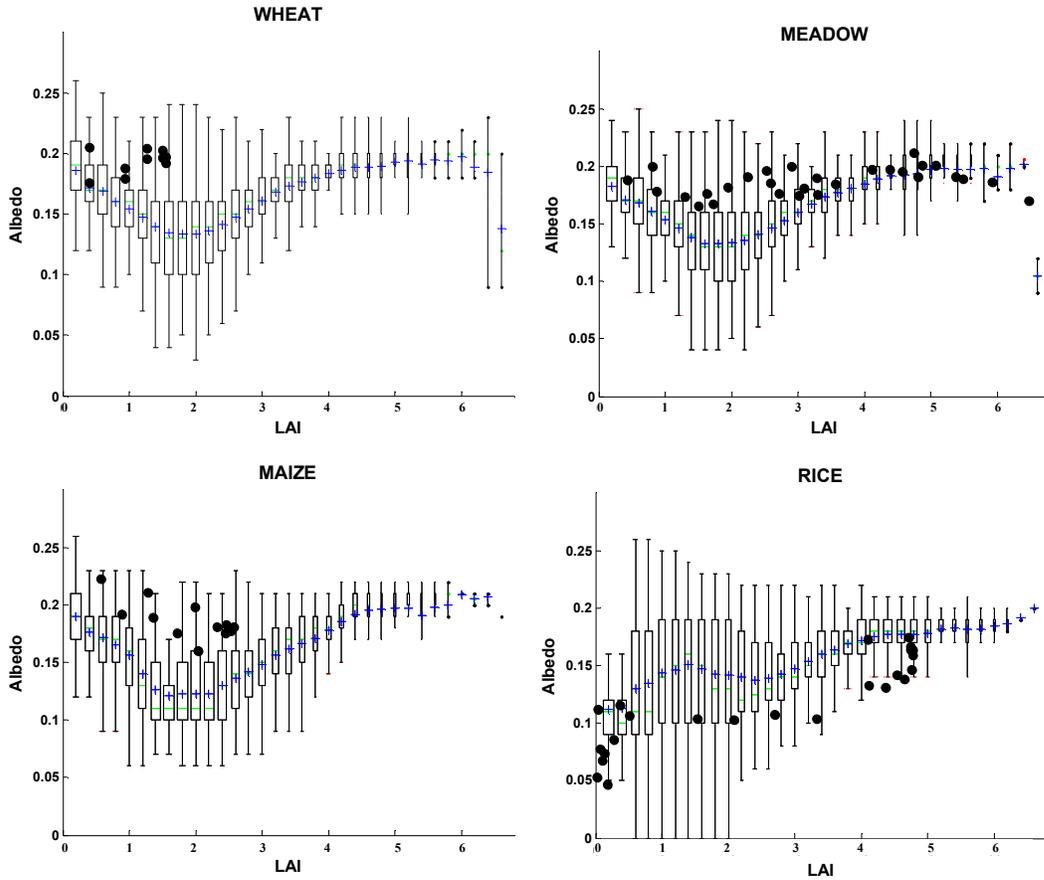
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955 **Figure 9: Dynamics of LAI over a) wheat1, b) wheat2, c) maize, d) rice, and e) meadow. Values of LAI are either**
 956 **measured (dots), interpolated between measurements (solid line), or derived from FORMOSAT-2 data by using**
 957 **NNT_{NDVI} approach (x). Vertical bars illustrate the 95% confidence interval around the ground based**
 958 **measurements mean. Vertical arrows in e) represent meadow cutting dates.**

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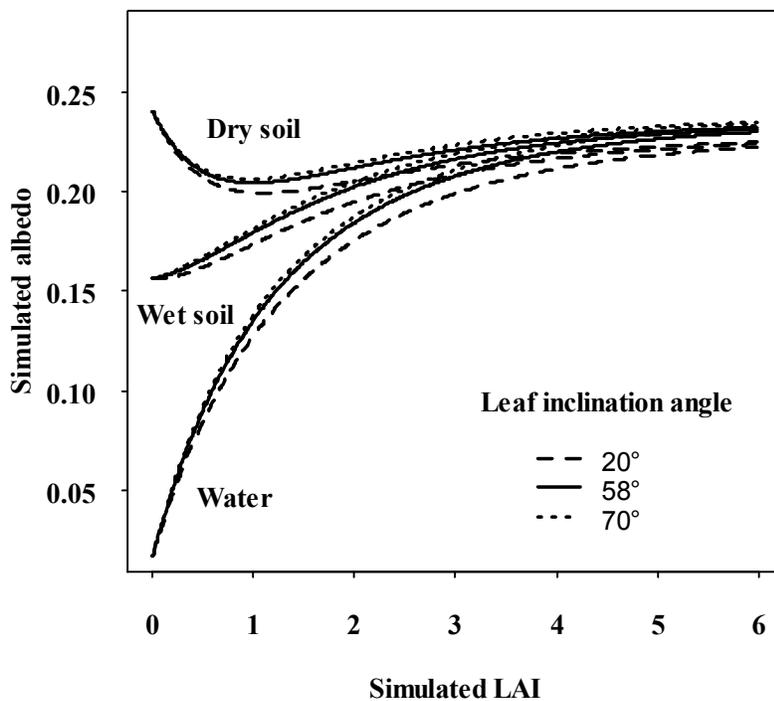
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964 **Figure 10: Boxplots of FORMOSAT-2 albedo retrievals as a function of FORMOSAT-2 estimated leaf area**
 965 **index LAI classes (0.2 LAI steps), for all available observations and for each vegetation type derived from the**
 966 **land use classification (section 2.1.1). (+) symbols represent the mean estimated albedo for each estimated LAI**
 967 **class, and (•) symbols represent albedo versus LAI measured in the field. The rectangles limit the lower and**
 968 **upper quartiles. The vertical bars limit the largest non-outlier observations.**

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974 **Figure 11: Theoretical relationship between albedo and LAI as simulated by the radiative transfer model**
975 **PROSAIL over three contrasted back-ground surfaces and three leaf inclination angles (Parameterization**
976 **details in Table 5).**

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981

982 **Table 1: Overview of vegetation characteristics observed for the five study fields. Coefficient of Variation (CV)**
983 **was derived from FORMOSAT-2 pixels by considering the whole field (borders excluded) for LAI (LAI CV) and**
984 **the albedometer footprint for albedo (Albedo CV). For meadow (Field #3), 1st and 2nd cycle corresponded to**
985 **growth cycles after cuts. Values of LAI and albedo CV are discussed in Section 2.1.5.**

986

Studied Fields	Field area (ha)	Period of data collection	Number of days with collection of ground based LAI data	Maximum LAI	LAI CV (%)	Albedo CV (%)
#1 - Wheat 1	30	March 10 th – May 30 th	6	1.55	1.0	1.5
#2 - Wheat 2	10	March 10 th – May 30 th	6	2.0	1.5	1.0
#3 - Meadow	30	March 10 th – October 10 th	14	6.3 (1 st cycle), 4.8 (2 nd cycle)	3.6	1.1
#4 - Maize	2	June 1 st – August 8 th	5	2.5	17.1	2.6
#5 - Rice	4	June 1 st – August 27 th	5	4.8	2.1	1.2

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990 **Table 2: Description of datasets for albedo and LAI ground based measurements, after spatial and temporal**
991 **matching with FORMOSAT-2 data.**

992

Variable	Ground measurements	Spatial matching	Temporal matching	Dataset size	Range of values
Albedo	Albedometer	Measurement footprint (4×4 pixels)	Measurements at 10:30 UT	130	[0.10 0.27]
LAI	Hemispherical images	Entire field (≈300 pixels)	Temporal interpolation using equation (1)	97	[0.00 6.8]

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996 **Table 3: Models and variables used for simulating of albedo and LAI. Values corresponded to typical average**
997 **soil and vegetation characteristics representative of our experimental conditions.**

998

MODELS & Variables	Values
PROSPECT (Leaf optical characteristics)	
Chlorophyll content	60 $\mu\text{g}/\text{cm}^2$
Dry matter	0.0075
Relative water content	80 %
Mesophyll structure parameter	1.2
Simulation spectrum	300 to 2400 nm
SAIL (radiative model)	
Zenith angle	30°
Diffuse fraction	Derived from 6S (See below)
Hot spot parameter	0.01
Leaf inclination angles	20°, 58° and 70° with ellipsoidal distribution
LAI values	0 to 6 (Step = 0.2)
Soil background	Water, wet soil, dry soil
6S (radiative transfer model)	
Aerosol model	Continental
Aerosol optical thickness	0.2347 at 550 nm
Environmental reflectance	Equal to the reflectance of the target

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1002 **Table 4: Coefficients sets used to compute albedo as a linear combination of waveband FORMOSAT-2**
 1003 **reflectances. Set 1 and NNT were calibrated over the Crau-Camargue dataset using FORMOSAT-2 off nadir**
 1004 **bidirectional reflectances. Sets 2 and 3 were designed for generic spectral configurations, and calibrated over**
 1005 **simulated data set by considering hemispherical reflectances. Set 4 was devoted to the PoIDER derived**
 1006 **hemispherical reflectances, and calibrated over the Alpilles-Reseda measured dataset. Set 5 was devoted to the**
 1007 **MISR sensor, and calibrated over a simulated dataset by considering nadir bidirectional reflectances. The**
 1008 **waveband limits (in nm) considered for each coefficient set are reported. Are also indicated the sum of**
 1009 **coefficients ($\Sigma\beta$), the performances of each coefficients set in terms of absolute bias (ME_A), absolute ($RMSE_A$)**
 1010 **and relative ($RMSE_R$) root mean square errors. The last column corresponds to the number of coefficients to be**
 1011 **adjusted in each empirical model.**

1012

	Green	Red	NIR	β_0	$\Sigma\beta$	ME_A	$RMSE_A$	$RMSE_R$	Freedom degree
FORMOSAT	520-600	630-690	760-900						
Set 1	0.000	0.619	0.402	0.000	1.021	0.000	0.015	7.3 %	3
NNT	-	-	-	-		0.000	0.007	3.5 %	11
Weiss et al. (1999)	560	665	855						
Set 2	0.000	0.570	0.460	0.000	1.030	-0.012	0.019	9.7 %	2
Set 3	0.680	0.080	0.350	0.000	1.110	0.021	0.029	14.2 %	3
Jacob et al. (2002a)	530-570	650-690	845-885						
Set 4	0.000	0.591	0.374	-0.001	0.965	0.014	0.020	9.7 %	3
Liang et al.(1999)	544-571	662-682	847-886						
Set 5	0.126	0.343	0.415	0.004	0.884	0.018	0.027	13.4 %	4

1013

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1015

1016 **Table 5: Obtained parameters when calibrating the various empirical methods we considered for the retrieval of**
 1017 **LAI from FORMOSAT-2 data. Are also indicated the performances of each method in terms of absolute bias**
 1018 **(ME_A), absolute (RMSE_A) and relative (RMSE_R) root mean square errors. The last column corresponds to the**
 1019 **number of coefficients to be adjusted in each empirical model. First is the multiple linear regression over**
 1020 **FORMOSAT-2 reflectances. Second is the NNT based method that uses as inputs FORMOSAT-2 reflectances.**
 1021 **Third is the heuristic formulation of LAI as a function of NDVI. Fourth and last is the NNT based method that**
 1022 **uses as input FORMOSAT-2 NDVI.**

1023

Coefficients	Red	NIR	β_0	K_{LAI}	$NDVI_{\infty}$	$NDVI_s$	ME _A	RMSE _A	RMSE _R (%)	Freedom degree
Multiple regression	-17.91	12.26	0	-	-	-	0.03	1.29	64.58	3
NNT_{Ref}	-	-	-	-	-	-	-0.03	0.69	34.63	11
NDVI based method	-	-	-	0.71	0.89	0.10	0.05	0.66	33.37	3
NNT_{NDVI}	-	-	-	-	-	-	0.02	0.55	27.54	7

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