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MEASURING AND MODELING SOIL INTRA-DAY VARIABILITY OF THE ¹³CO₂ & ¹²CO₂ PRODUCTION AND TRANSPORT IN A SCOTS PINE FOREST

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CONTEXT

- Future dynamics of soil CO₂ efflux (*F_s*) = key question in climate change research.
- Two processes lead to *F_s*: CO₂ production *P* (heterotrophic + autotrophic) + transport *F* to surface
- Factors affecting *P* & *F* (t°, moisture, substrate...) vary vertically and temporally.
- Better understanding of the mechanisms controlling *F_s* needs - soil **multilayer approach**
 - fine temporal resolution (intra-day) study
- Use of ¹³C & ¹²C improve *EF_s* origin (partitioning) and pathway (spatiotemporal tracer) knowledge

OBJECTIVE 1: Determine soil horizons *P_s* and $\delta^{13}P_s$

Based on ¹²CO₂ & ¹³CO₂ balances for layer i
&
Diffusive Flux-Gradient approach

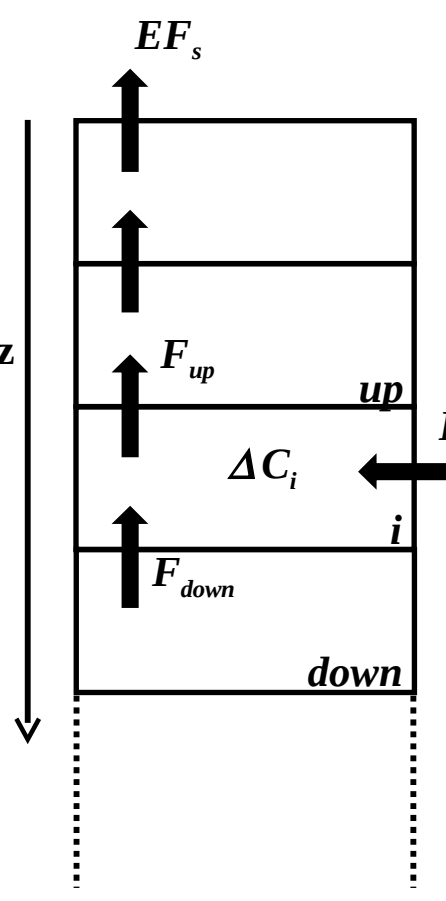
$$\varepsilon_i * \frac{\Delta C_i}{\Delta t} = \frac{F_{down} - F_{up}}{thick_i} + P_i \quad (Eq. 1)$$

$$F_x = D_{x-i} * \frac{C_x - C_i}{z_x - z_i} = \left(D * \frac{\Delta C}{\Delta z} \right)_{x-i} \quad (Eq. 2)$$

where *C_i* = [¹²CO₂] or [¹³CO₂]; ε_i & *thick_i* = layer i soil porosity & thickness;

$$D_{x-i} = \text{soil diffusivity at the x-i boundary}$$

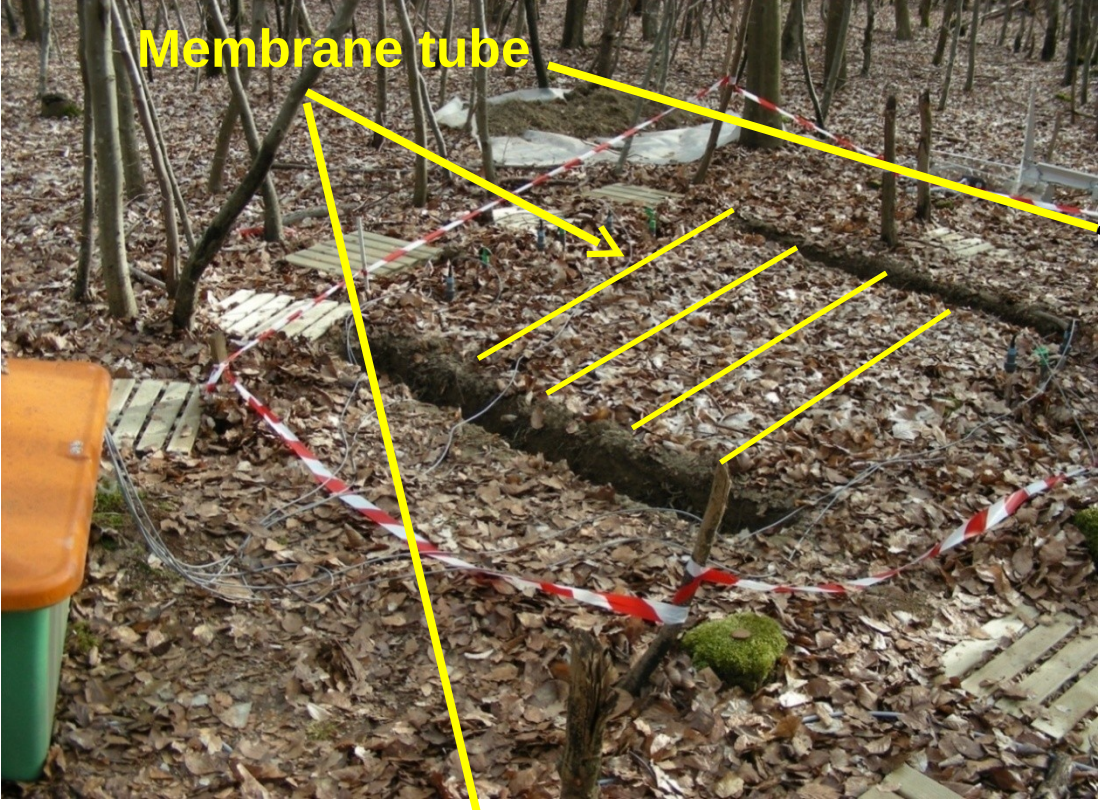
$$P_i = \varepsilon_i * \frac{\Delta C_i}{\Delta t} + \frac{\left(D * \frac{\Delta C}{\Delta z} \right)_{down-i} - \left(D * \frac{\Delta C}{\Delta z} \right)_{up-i}}{thick_i} \quad (Eq. 3)$$

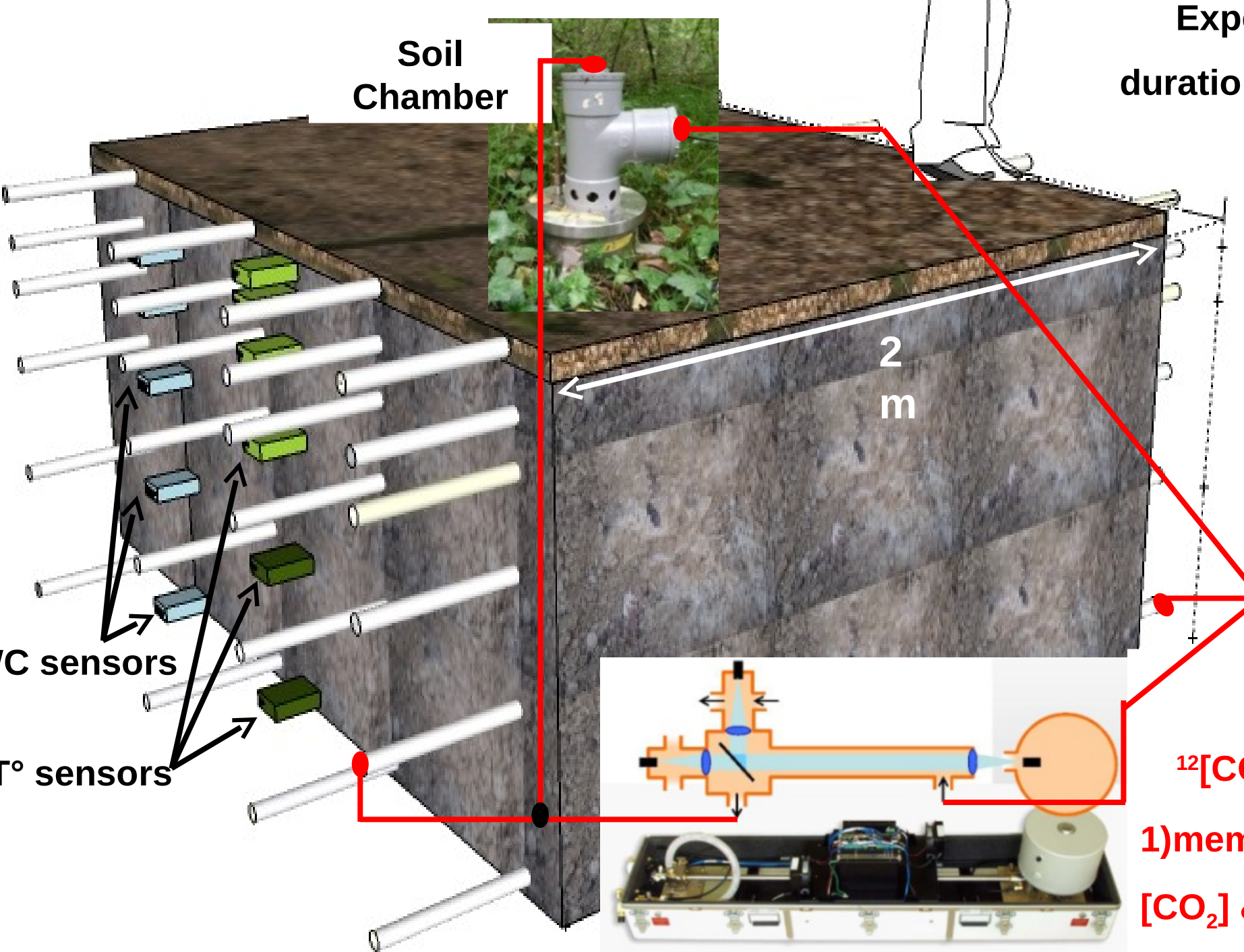


P_i and $\delta^{13}P_i$ can be obtain by measuring the temporal evolution of the [¹²CO₂] & [¹³CO₂] vertical profiles + Lab. measurement (on soil samples) of ε and dependence of *D* on soil water content (vertical profiles)

Boundary conditions : *F_{bottom}* = 0; *F_{top}* = *F_s*

Experimental device (Parent *et al.* 2013)





Experiment duration : 20 days

2 m

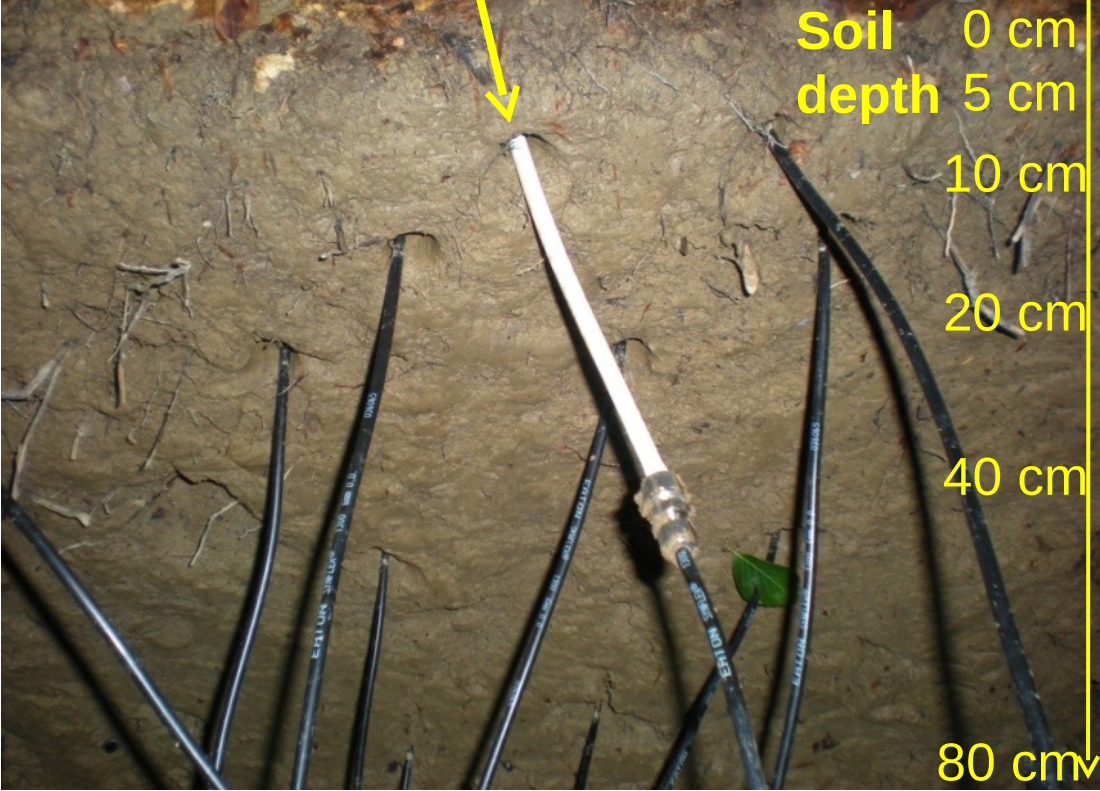
SWC sensors

T° sensors

TDLs: ¹²[CO₂] & ¹³[CO₂]

1)membrane tube ≡ [CO₂] & δ^{13} CO₂ in soil layers

2)from chamber ≡ *EF_s* & $\delta^{13}EF_s$



Soil depth

0 cm

5 cm

10 cm

25 cm

40 cm

80 cm

Results

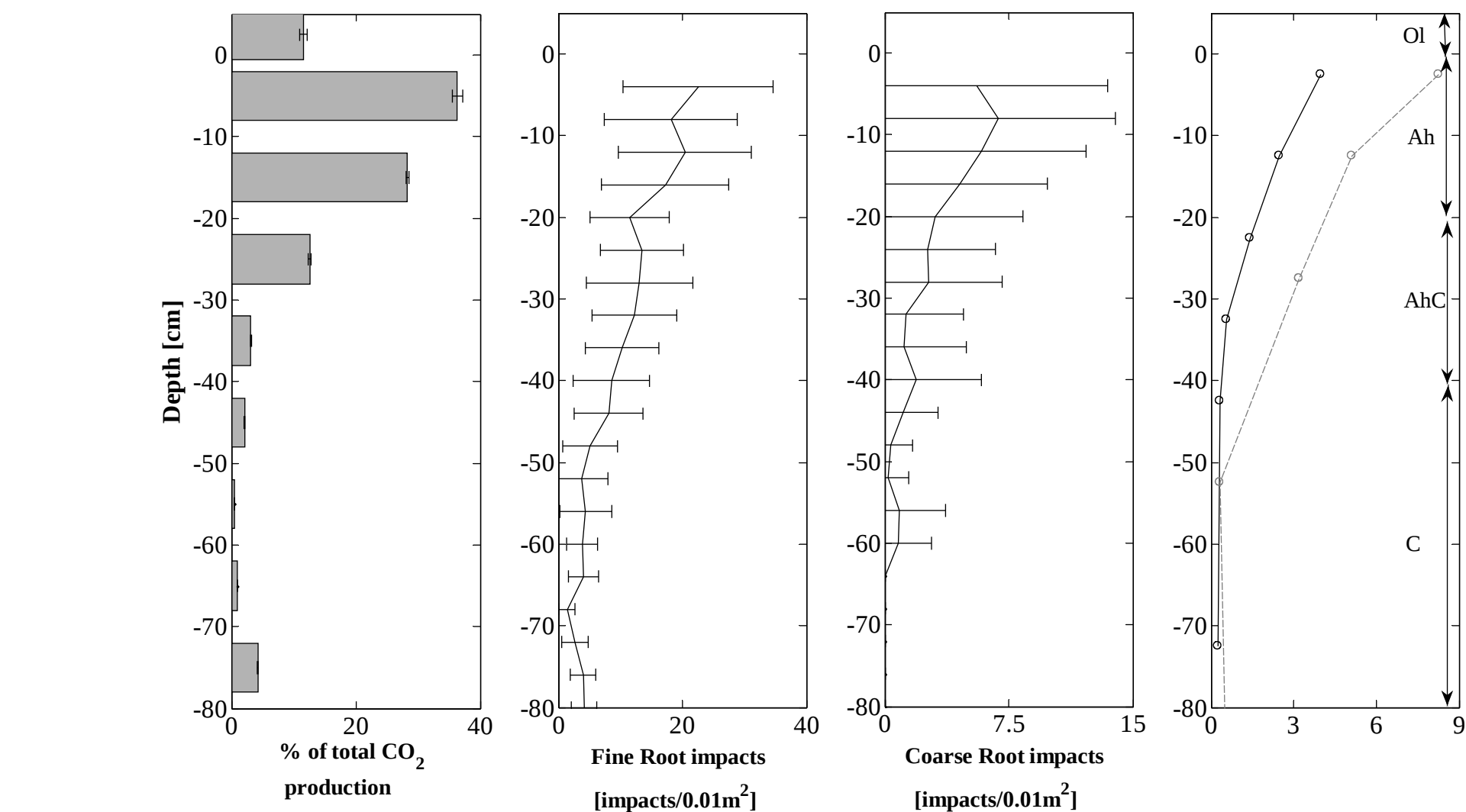


Fig. 1: Mean *P_s* (over the campaign) of soil horizons (Ol, Ah, AhC, C) in parallel with fine roots and carbon content vertical profiles

- **CO₂ production decrease with depth as fine roots & carbon contents**
- **Clear intraday variability in Ol Ah & AhC horizons**
- **Not plausible negative CO₂ production (consumption) in Ol ! (see explanation in Objective 2 section)**

Fig. 3: Temporal evolution of $\delta^{13}P_s$ for Ah (a & b) and Ahc (c & d). Hourly variability (a & c) and daily means (b & d) are presented

- **Significant day to day variation of $\delta^{13}P_s$ in Ah**

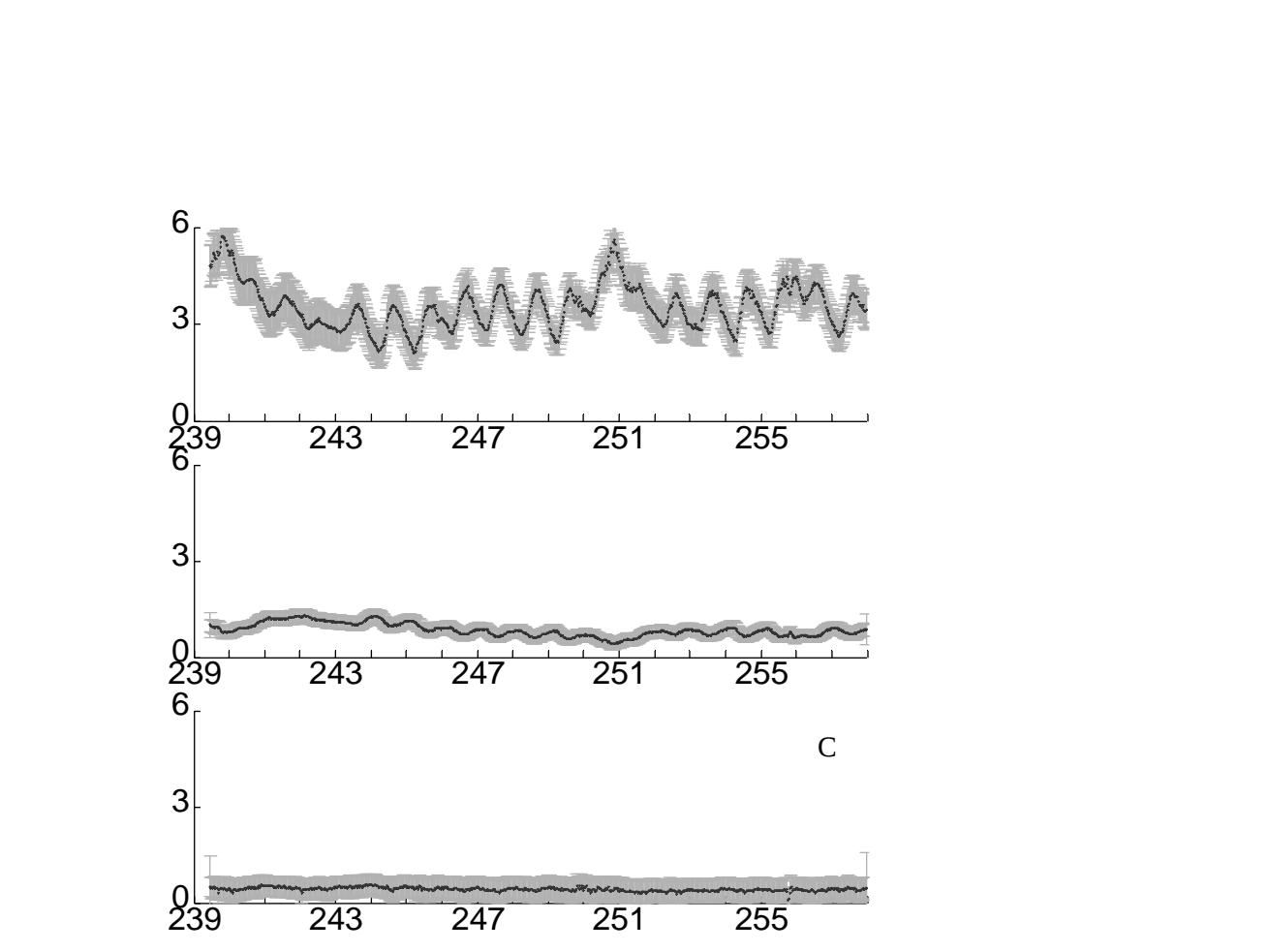
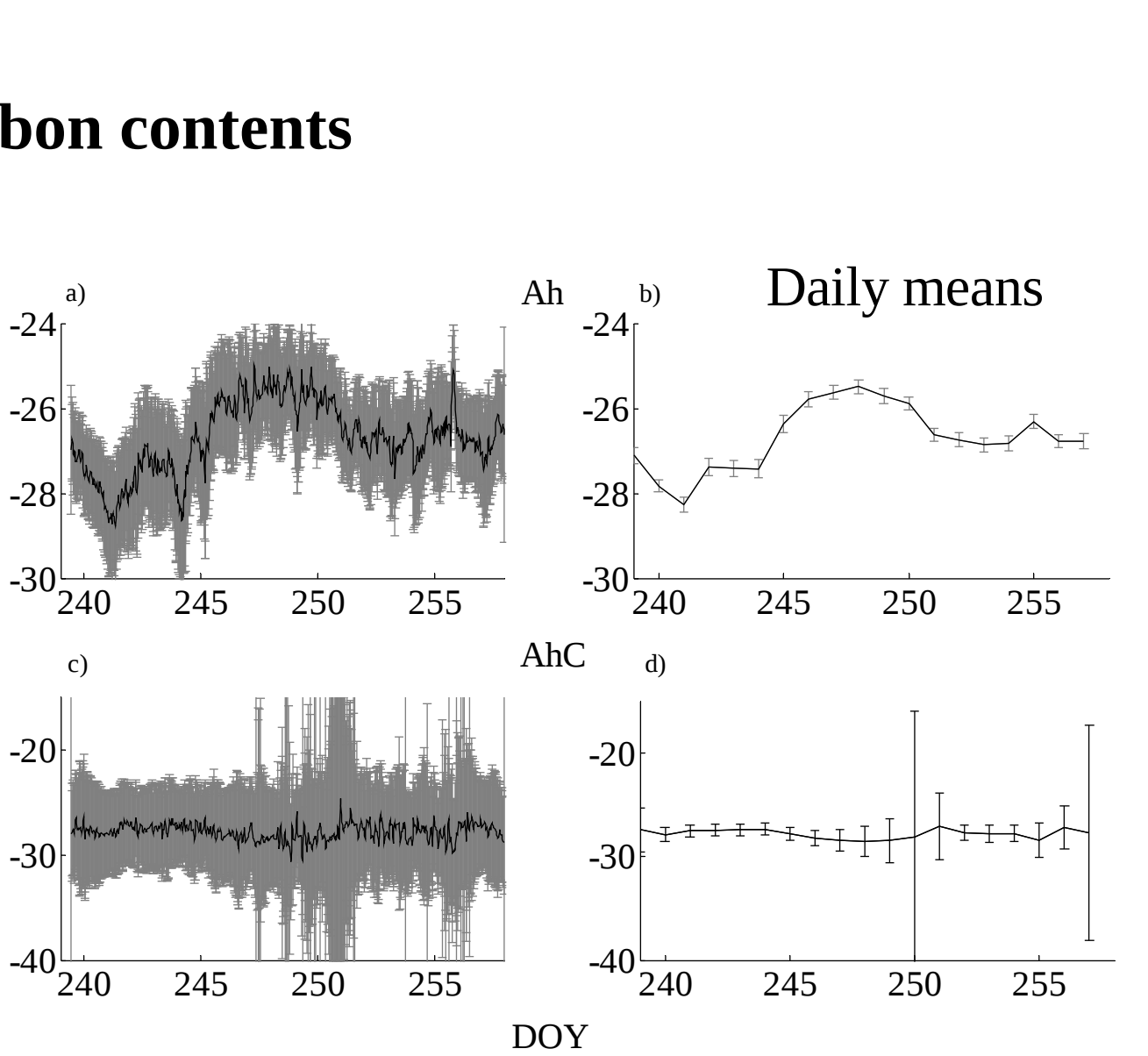


Fig. 2: Temporal evolution of the CO₂ production for the different horizons

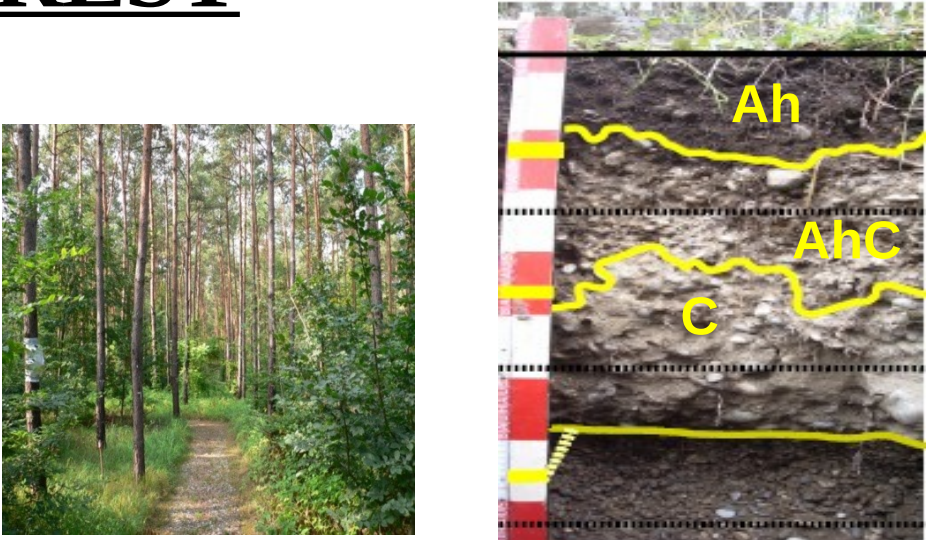


OBJECTIVES

- 1)Determine *P_s* and its isotopic signature $\delta^{13}P_s$ for the different soil horizons.
- 2)Find factors affecting *P_s* & $\delta^{13}P_s$ intra & inter day fluctuations
- 3) Evaluate by modeling which processes (those linked to production or transport) drive *F_s* temporal variability

EXPERIMENTAL SITE: HARTHEIM FOREST

- 46 year old Scots Pine Forest (*Pinus sylvestris* L.)
- Mean annual Air Temp/Prec: 10.3°C/642 mm
- Haplic Regosol (calcaric, humic)
- Humus type is mull (1-3 cm thick)



OBJECTIVE2: Factors affecting *P_s* and $\delta^{13}P_s$ (Goffin *et al.* 2014)

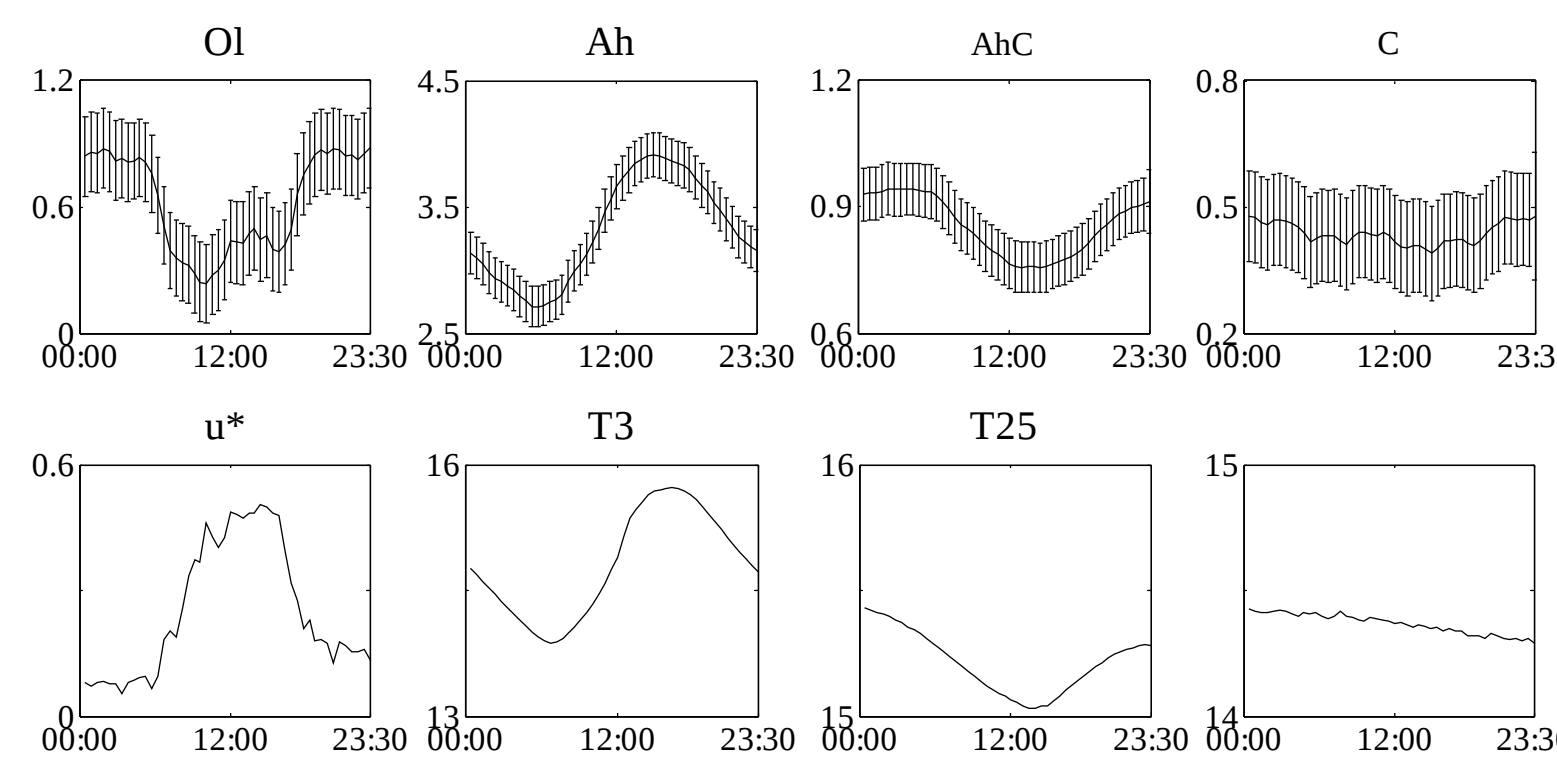


Fig. 4: Mean diurnal evolution of the *P_s*, soil temperature (3, 25, 70 cm depth) and friction velocity (*u**)

- *P_s* in Ah and AhC are correlated to their local temperature
- No diurnal evolution of *P_s* in C (t° ≈ constant over the campaign)
- *P_s* in Ol (litter) is correlated to *u**. The CO₂ transport processes due to turbulence are not represented in Eq. 1 & 3 when they are actually present in the litter layer. Consequently their influences are reflected back in the *P_s* term (≡ Eq. 3 output). This point explains also negative value of *P_s* in Ol met in Fig. 2

- Significant dependence of $\delta^{13}P_s$ in Ah on SWC (enrichment with drought) could be imputed to change in the substrate / microbial communities involved in C mineralization when SWC decreases.

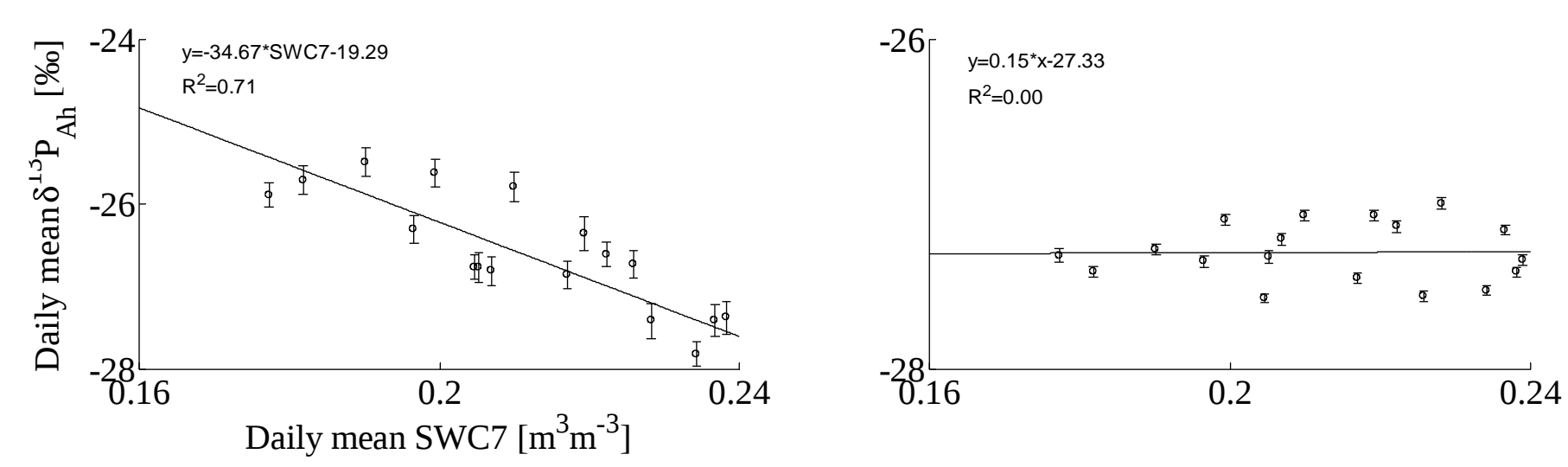


Fig. 5: Dependence of daily mean (a) $\delta^{13}P_s$ in Ah and (b) $\delta^{13}F_s$ on soil water content (SWC) at 7 cm depth

- This dependence is not observed for efflux ($\delta^{13}F_s$) because *F_s* is a mixed of ≠ layer contributions with ≠ dependences and the presence of soil chamber can have induced difference between SWC below them and the SWC measuring points.

OBJECTIVE 3: Model applications to evaluate the part of production and transport processes in the [CO₂]_{soil} & *F_s* temporal variability (Goffin *et al.* under review).

- 1) Model run with a reference version simulating production in the ≠ layers (rate depending on local t°) and diffusion as the only CO₂ transport process
- 2) Model runs with testing versions where (a) advection, (b) dispersion and (c) production regulated by phloem pressure concentration waves (PPCW) are successively added
- 3) Comparison of the reference and testing versions outputs with [CO₂]_{soil} & *F_s*

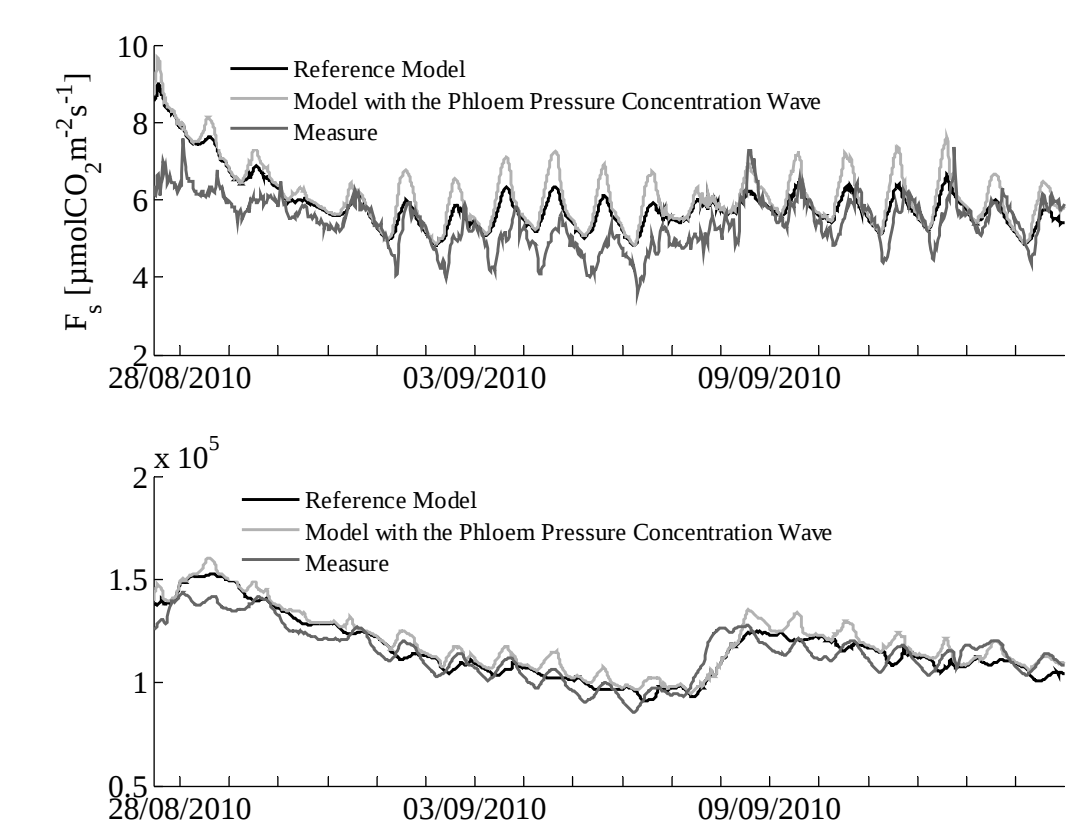


Fig. 6 a & c: Time evolutions of *F_s* and [CO₂]_{soil} obtained with reference model, model including the phloem pressure concentration wave (PPCW) and measurements

Fig. 6 b & d: Corresponding averaged intra-day variability

- **Adding PPCW**
 - Improve significantly the simulation of *F_s* & [CO₂]_{soil} amplitude and phase
 - Best way to reproduce the intra-day *F_s* & [CO₂]_{soil} variations