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Didier Bert, Jean-Luc Dupouey, S. Leavitt. Long-Term increasing WUE revealed by  $\delta 13$ C in tree-rings.. International Seminar on "Causes and consequences of accelerating tree growth in Europe", May 1998, Nancy, France. 1 p., 1998. hal-02836720

### HAL Id: hal-02836720 https://hal.inrae.fr/hal-02836720

Submitted on 7 Jun 2020

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## LONG-TERM INCREASING WUE REVEALED BY 813C IN TREE RINGS



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INTERNATIONAL SEMINAR ON " CAUSES AND CONSEQUENCES OF ACCELERATING TREE GROWTH IN EUROPE".
NANCY, FRANCE 18-19 MAI 1998.

The isotopic composition of tree rings carries an integrated annual record of environmental conditions, in particular temperature and precipitation. However, models of carbon discrimination during carbon fixation show that C3 plants are not passive recorders: the carbon isotopic variations are subjected to strong physiological control through leaf gas exchange regulation (Francey and Farquhar 1982). Therefore, records of carbon discrimination in tree-ring cellulose could be used to study past variations of the ecophysiology of silver fir (Abies alba Mill.), a common conifer in European mountains, over the last century (Bert et al., 1997).

#### **PRINCIPLE**

The isotopic composition is measured by mass spectrometry and expressed in reference to the PDB standard:

$$\delta^{13}\text{C }\%_{o} = (\frac{13_{\text{Csample}}}{12_{\text{Csample}}} - 1) \times 1000 \qquad (1)$$

$$\frac{13_{\text{CPDB}}}{12_{\text{CPDB}}} - 10 \times 1000 \qquad (1)$$

A typical value for air in 1990 was \$13Cair = -8% therefore the A splical value of  $\alpha$  in Firsty was  $\alpha \sim C_{eff} \approx -0.8$ , therefore the discriminations in the C3 plants give typical values  $\delta^{13}C_p \approx -22\%$ , to -34%. This carbon discrimination is expressed as  $(\delta_a$  is  $\delta^{13}C_{air}$  and  $\delta_p$  is  $\delta^{13}C_{plant}$ ):

$$\Delta\% = \frac{\delta_a \cdot \delta_p}{1 + \delta_p} \tag{2}$$

dels of discrimination (Farquhar et al., 1982) show that  $\Delta$  related to the ratio between the internal gas phase ssure of CO<sub>2</sub>, P<sub>i</sub>, and the atmospheric partial pressure,

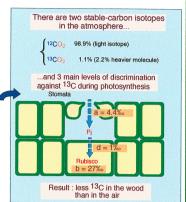
$$\Delta = a - d + (b - a) \frac{P_i}{P_a}$$
 (3)

In addition, net photosynthesis, A, measured as CO2 uptake and leaf conductance to CO2, gc, are linked by Fick's law

$$A = g_c (P_a - P_i)$$
 (4)

Given that, gw, the leaf conductance to water vapor, is 1.6 g and using equations (3) and (4), discrimination can be related to the intrinsic water-use efficiency "WUE" (Ehleringer et al. 1993) by:

$$\Delta = a - d + (b - a) \left( 1 - \frac{1.6}{P_a} \left( \frac{A}{g_w} \right) \right)$$
 (5)



Direct and indirect atmospheric estimates of  $\delta_{\boldsymbol{a}}$ are now available; therefore, equations (2, 5) can be used in order to study long-term changes of  $\Delta$  and thus of intrinsic water-use efficiency from

measurements of  $\delta_a$  and  $\delta_p$ . This *intrinsic* WUE is part of the WUE of the whole plant which depends also on leaf to air wa gradient, the environmental component.

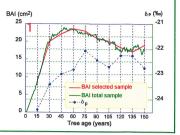
However, the wood formed when the tree was young is significantly more <sup>13</sup>C depleted than wood formed at later stages of tree life. Thus, an adequate sampling scheme has to be designed to distinguish such an 'age' effect and the effects of environmental changes on Δ.

#### TREE-AGE EFFECT ON δ<sub>n</sub>

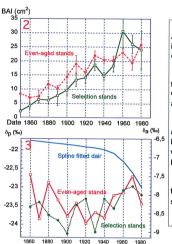
We confirm the existence of an 'age' trend, independent of any long-term environmental changes (Fig.1). This  $\delta_D$  pattern could be attributed to

 internal age-related physiological factors
 changes of environmental factors associated with tree height, canopy closure and stand level competition (concentration of respired CO<sub>2</sub>, soil moisture and light regime, etc.)

This aging effect has to be taken into account when comparing the stable-carbon isotope content of rings formed at different tree ages.



#### **LONG-TERM CHANGES**



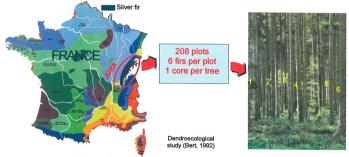
The basal area of tree rings formed at the age of 40 increased after circa 1860 (Fig.2). The mean BAI is significantly different for the two types of stands but  $\delta_{\rm D}$  values showed no significant difference (Fig.3).

 $\delta_{D}$  should have paralleled  $\delta_{B}$  if the plant fractionation was not affected by any long-term environmental changes. Actually, the variations of  $\delta_{D}$ and δ<sub>a</sub> were very different after 1930s (Fig.3).

Discrimination  $\Delta$  kept constant before the 1930s and significantly decreased afterwards (Fig.4). The P<sub>i</sub>/P<sub>a</sub> ratio followed the same changes (Eq.3, Fig.4), which allows to reconstruct the Pi variation because Pa is known (Fig.5).

During the period 1930-1980, Pi increased slower than Pa and the intrinsic water-use efficiency significantly increased 30% (Fig.6).

# SAMPLING SILVER FIRS IN THE JURA MOUNTAINS



#### Δ (%) 18 0.57 0,56 17.5 0.55 0.52 0.51 1860 1880 1900 1920 1940 1960 1980

Our observation of an in situ increase of intrinsic WUE of a forest tree species is a first upscaled confirmation of similar observations made in controlled experiments (CO<sub>2</sub> increase, temperature etc.).

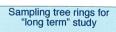
However, we are faced with the difficulty of 0,53 ascribing this observed ecophysiological change in the forest to the effect of improvement of silvicultural methods or macroclimatic changes, increasing atmospheric CO<sub>2</sub> concentration, atmospheric deposition of nitrogen, etc.

#### Sampling tree rings for 'age effect" study

150 years of stand life were studied with 15-yrs steps. Bias Du years of stand life were studied with 15-yrs steps. Bias of  $\delta_{p}$  measurements was limited: studied tree ring pooled with its 4 neighbors; studied age reached between 1940 and 1988; site effects averaged by choosing ten trees at different

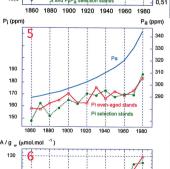
For each studied age, ten tree-ring pentads were ground

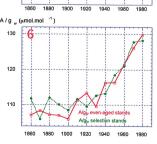
Example for 5-yr-old



Ten trees which reached 40-yrs-old during the time period were selected per decade (1860-69... 1980-88). The tree-ring pentad of 38 to 42-yr-old was extracted from each core. The 10 pentads were ground together and one  $\delta_p$  value measured for each decade.

Selection forests and even-aged forests revealed different increasing BAI long-term trends (Fig.2) and the sampling process was repeated for both.



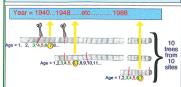


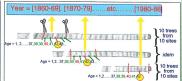
#### CONCLUSION

The studied variations have important implications for the use of carbon isotope ratios in tree rings as a palaeorecorder of past environments.

The long-term component is better suited for the reconstruction of year-to-year deviations of climatic parameters from their long-term average (Dupouey et al. 1993).

The relationship between the isotopic composition of tree rings and climatic parameters is often calibrated during the recent period of rapid environmental changes. It may perhaps not be directly applicable to previous time periods, at least for the reconstruction of absolute climatic values.





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