1 Pathways to persistence: plant root traits alter carbon accumulation in different soil carbon pools

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 respiration, microbial biomass

25 ABSTRACT

- Aims: Mineral-associated organic matter, mainly derived from microbial by-products, persists longer in soil compared to particulate organic matter (POM). POM is highly recalcitrant and originates largely from decomposing root and shoot litter. Theory suggests that root traits and growth dynamics should affect carbon (C) accumulation into these different pools, but the specific traits driving this accumulation are not clearly identified.
- 31 *Methods:* Twelve herbaceous species were grown for 37 weeks in monocultures. Root 32 elongation rate (RER) was measured throughout the experiment. At the end of the experiment, we 33 determined morphological and chemical root traits, as well as substrate induced respiration (SIR) as a

proxy for microbial activity. Carbon was measured in four different soil fractions, following particle size and density fractionation.

36 *Results:* In N₂-fixing Fabaceae species, root biomass, RER, root diameter, hemicellulose content 37 and SIR, were all positively correlated with increased C in the coarse silt fraction. Root diameter and 38 hemicellulose content were also negatively correlated with C in the POM fraction, that was greater 39 under non N₂-fixing Poaceae species, characterized by lignin-rich roots with a high carbon:nitrogen 40 ratio that grew slowly. The accumulation of C in different soil pools was mediated by microbial activity.

Conclusions: Our results show that root traits determine C input into different soil pools, mediated
primarily by microbial activity, thus determining the fate of soil organic C. We also highlight that C in
different soil pools, and not only total soil organic C, should be reported in future studies to better

44 understand its origin, fate and dynamics.

45 Abbreviations

Abbreviation	Meaning						
С	Carbon						
POM	Particulate organic matter						
C:N	Carbon – nitrogen ratio in plant tissue and/or soil						
N_2 -fixing	Dinitrogen fixing						
t0	Time zero, beginning of the experiment						
t37	Time 37 weeks, end of the experiment						
٨٢	Delta carbon, as difference between carbon at time 0 and carbon at time 37, in						
	different fractions (mg C g ⁻¹ soil)						
Сром	Carbon in the coarse POM 200-2000 μm fraction (mg C g^-1 soil)						
C _{finePOM}	Carbon in the fine POM 50-200 μm fraction (mg C g $^{\text{-1}}$ soil)						
C _{SILT}	Carbon in the 20-50 μm coarse silt fraction (mg C g^-1 soil)						
C _{SILT+CLAY}	Carbon in the fine silt+clay <20 μm fraction (mg C g $^{-1}$ soil)						
	Sum of delta carbon in different fractions, $\Delta C_{SUM} = \Delta C_{POM} + \Delta C_{finePOM} + \Delta C_{SILT} +$						
	$\Delta C_{SILT+CLAY}$ (mg C g ⁻¹ soil)						
RER	Root elongation rate (mm d ⁻¹)						
RLP	Root length production (m)						
	RER and RLP of 'new' roots initiated during the 2 weeks interval between						
RER _{NEW} , RLP _{NEW}	measurements						
RER _{OLD} , RLP _{OLD}	RER and RLP of 'old' roots, initiated more than 2 weeks before the measurement						

SIR	Substrate induced respiration (μ g C-CO ₂ g ⁻¹ soil h ⁻¹)
PCA	Principal component analysis

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47 INTRODUCTION

48 Given the current climate change emergency, several international initiatives have been launched to 49 unlock the potential of soils to sequester atmospheric carbon (C) (e.g. 4 per Thousand Initiative, 50 Minasny et al. 2017). Better understanding the interactions between vegetation and soil has become 51 central for sequestering C into anthropogenically disturbed soil, such as agricultural fields, mining 52 waste soil, road embankments and technosols (Paustian et al. 2016; Griscom et al. 2017). Plants act as a major conduit for transferring C into soils via litterfall, root mortality and exudation (Six et al. 2004; 53 Derrien et al. 2016; Sokol et al. 2019). Some C is transformed by soil microbes and released back into 54 55 the atmosphere by respiration (Jones et al. 2009; Kuzyakov and Larionova, 2005), but C can also be 56 stabilized in soil, increasing its residence time (Besnard et al. 1996; Lal, 2004; Rasse et al. 2005; 57 Bardgett et al. 2014; Vidal et al. 2018; Sokol et al. 2019). Carbon persists in soil at different time scales 58 based on recalcitrance (short-term preservation), spatial inaccessibility to decomposers due to 59 occlusion in soil aggregates, and adsorption to mineral and metal surfaces (Kleber et al. 2011, Schmidt 60 et al. 2011, Poirier et al. 2018). These mechanisms are influenced by abiotic and biotic factors and 61 especially by plant roots, since their C is preferentially stabilized compared to aerial parts (Balesdent 62 and Balabane, 1996; Rasse et al. 2005, Kätterer et al. 2011). In a recent review, Poirier et al. (2018) 63 argued that the root traits most influencing C stabilization are those related to chemical composition, 64 root exudation and the presence of symbionts (e.g. mycorrhizas and *Rhizobium* bacteria), whereas the 65 role of morphological traits is not yet clear. More specifically, root traits increasing chemical 66 recalcitrance promote short-term C stabilization by slowing root decomposition rates, whereas root 67 traits increasing exudation rate promote long-term C stabilization faster. Several studies have analysed 68 the link between plant functional traits, microbial activity and C accumulation (Chapin 2003; Lavorel 69 et al. 2007; De Deyn et al. 2008; Poirier et al. 2018). However, as yet, it is not understood how root 70 traits can alter the accumulation and potential persistence of C.

71 Through differences in chemical and physiological traits, roots should affect C accumulation into 72 different C pools depending on soil texture. These pools are defined as: i) coarse particulate organic 73 matter (coarse POM, > 200µm fraction), that is free in the soil at different levels of degradation, ii) fine 74 POM (50-200µm fraction), that comprises free organic C and organic C occluded in soil 75 macroaggregates. These two pools are mostly derived from the decomposition of roots and shoots 76 (Kögel-Knabner, 2002), and their short-term C protection from microbial consumption relies mainly on 77 the recalcitrance of their lignocellulose C structures and the physical protection given by 78 macroaggregate structure (Six et al. 2002). Finally, iii) C protected in the coarse silt and fine silt+clay 79 pools (20-50µm and <20µm fractions, respectively); that is highly processed and protected from 80 microbial consumption via occlusion in microaggregates and through organo-mineral adsorption to 81 clay particles and metals. This C is mostly derived from dissolved organic C originating from the degradation of above and belowground plant C input (Bird et al. 2008; Rubino et al. 2010; Sanderman 82 83 et al. 2014), from root exudation of labile rhizospheric compounds and from microbial compounds 84 (Simpson et al., 2007; Mambelli et al., 2011; Cotrufo et al. 2013; 2014; Vidal et al. 2018; Rossi 2019). It 85 is now generally accepted that labile low molecular weight compounds persist in soil longer than 86 chemically recalcitrant C structures, when protected by organo-mineral adsorption (Mikutta et al. 87 2006; Kleber et al. 2015; King et al. 2019; Robertson et al. 2019; Sokol et al. 2019). The stability of 88 sequestered C in soil is therefore linked to the fraction of soil to which it is associated, with a greater 89 stability of C pools associated with finer fractions (Torn et al., 2009).

90 As C accumulation into the coarse POM pool is related to the amount of recalcitrant matter present, it 91 should therefore be greater in soils containing roots with high cellulose, lignin and carbon:nitrogen 92 ratio (C:N) (Poirier et al. 2018). However, it is C-rich exudates produced during fine root elongation 93 that promote long-term C stabilization in the coarse silt and fine silt+clay fractions (Mikutta et al. 2006; 94 King et al. 2019; Robertson et al. 2019; Sokol et al. 2019), and together with mucilage and border cells 95 (shed during growth), are important substrates for microbial communities (Dennis et al., 2010). These 96 C substrates that are assimilated by microorganisms close to the root apex are utilized rapidly for 97 respiration and growth, or lost as microbial exudates or exopolysaccharides that are then used as a 98 substrate for subsequent microbial communities. Since microbial byproducts (from activity in any soil 99 C pool) are believed to be the main precursor of protected C due to organo-mineral interactions 100 (Simpson et al., 2007; Mambelli et al., 2011; Cotrufo et al., 2013; Rossi 2019), root exudation should 101 influence the C storage in this fraction. It is however important to mention that exudation and the 102 resulting microbial activity can also negatively influence soil C storage, increasing the consumption of 103 preexistent soil C (i.e. priming effect; Hamer and Marschner, 2005; Shahzad et al. 2017). Root traits 104 related to exudation are however poorly understood; the few studies available showed contrasted 105 results and relate to root morphological traits measured at the whole root system level (Roumet et al. 106 2006; Guyonnet et al. 2018). Because root exudates are mainly released at the elongating root tip, 107 where rhizospheric microbial activity is high (Jones et al. 2009; Canarini et al. 2019), it can be expected 108 that root elongation rate (RER) is a powerful predictor of C deposits in coarse silt and fine silt+clay C 109 pools (Holz et al. 2018). Root elongation rate is affected principally by local abiotic soil conditions such 110 as soil temperature, moisture, and compaction, but also differs among species, although most known 111 data is related to woody species (Steinaker et al. 2011; Mohamed et al. 2016; Wang et al. 2018). Fast 112 growing species with small diameter fine roots, high specific root length and N uptake rate usually have 113 high RER (Larson and Funk 2016).

114 Determining plant traits that increase C accumulation in different soil C pools should therefore enable 115 the selection of species that promote C persistence in soil. Recent debate has focused on the ability of 116 dinitrogen fixing (N_2 -fixing) species to sequester large amounts of C in soil (Plaza-Bonilla et al. 2016; 117 King et al. 2018). Bacteria such as *Rhizobium*, present in nodules of N_2 -fixing species, produce large 118 amounts of exopolysaccharides (Downie, 2010; Sasse et al. 2018), that are adsorbed onto fine silt and 119 clay particles (Fehrmann and Weaver, 1978). Also, N_2 -fixing species have roots that are easily 120 degradable with a high content of hemicelluloses (Hernández et al. 2017) and low C:N ratio 121 (Warembourg et al. 2003; Roumet et al. 2005), therefore enhancing microbial activity (Poirer et al. 122 2018). However, it is not known whether N₂-fixing species promote greater C accumulation in the fine 123 silt and clay soil fractions, thereby enhancing C persistence in soil.

124 We explored the effect of root traits on C accumulation into different soil C pools beneath 12 125 herbaceous species grown in monocultures for 37 weeks. These species had diverse root traits in terms 126 of morphology, chemical composition, and elongation rate and belonged to different plant families: 127 five N₂-fixing Fabaceae, five Poaceae, one Rosaceae and one Plantaginaceae. Our main hypothesis is 128 that C accumulation into different soil C pools is driven by root traits and their effects on microbial 129 activity and biomass. More specifically, we hypothesize that (i) traits related to high RER promote C 130 accumulation in the coarse silt and fine silt + clay C pools, since these traits are expected to favour 131 exudation and subsequent microbial activity, whereas (ii) root traits related to chemical recalcitrance 132 (high lignin and cellulose content and high C:N ratio), promote C accumulation in the unprotected POM 133 pool, and iii) N₂-fixing species favour C accumulation in the coarse silt and fine silt + clay pools. Results 134 should enable us to disentangle the relationships between root growth, traits and the accumulation 135 and stabilization of C in different soil C pools, between different families (Poaceae and Fabaceae) and 136 N₂ and non N₂-fixing species.

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139 MATERIALS AND METHODS

140 Experimental setup

The experiment was set up in the experimental garden of CEFE-CNRS Montpellier, France (43.6389° N°, 3.864125° E and lasted 37 weeks (from t0: Sept-2016 to t36: July-2017). Twelve herbaceous species were grown as monocultures in steel boxes (0.7 m length x 0.7 m width x 0.3 m depth): five N₂-fixing species from the Fabaceae family and seven non N₂-fixing species, including five Poaceae, one Plantaginaceae and one Rosaceae species (Table 1). A weather station was set up permanently in the experimental garden, and air humidity, air temperature (minimum, maximum and mean daily) andsolar irradiation (measured daily) were monitored throughout the experiment.

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149 Seventy-eight boxes were prepared: six replicate boxes per species and six additional boxes of bare 150 soil used as controls. Boxes were organized in three blocks with two rows of 13 adjacent boxes in each 151 block, and with a distance of 50 cm between each box. Each row comprised 12 monocultures (one per 152 species) and a bare soil, randomly arranged in each row (Fig S1). Boxes of the first row were used for 153 destructive plant and soil sampling, while the boxes of the second row were equipped with rhizotrons 154 for the study of root elongation (Section 1.1, Fig. S1). In the second block, each box was equipped with 155 soil temperature and humidity sensors placed at a depth of 0.1 m. Soil temperature was recorded 156 every 4 hours with an i-button sensor (iButtonLink, Wisconsin, USA); soil relative humidity was 157 recorded every hour with moisture sensors (Waterscout SM100, Spectrum Technologies Inc.) and a 158 datalogger (WatchDog weather station 200 series, Spectrum Technologies Inc.). These boxes were 159 undisturbed for the duration of the experiment. Rhizotrons comprised a 0.2 m width x 0.3 m depth x 160 0.05 m thick pane of transparent plexiglass set into the lower walls of the boxes, through which roots were observed and root elongation rate (RER, in mm root⁻¹ day⁻¹) and root length production (RLP, in 161 162 mm mm⁻² day⁻¹), were calculated (Fig. S2). For RER and RLP, only one replicate box per species was 163 analyzed, because the analysis of root images was extremely time consuming.

Boxes were inclined at 20° relative to the horizon to encourage the positive geotropism of roots when they came into contact with rhizotron windows (Huck and Taylor, 1982). Boxes were filled with soil sieved to 8 mm. Five layers of soil were successively added and manually compacted to attain a volume of 0.113 m³, i.e., a total of 190 kg of soil per box (bulk density = 1.70 ± 0.02 g cm⁻³). The soil, excavated in Villefort (France; 44°26′25″ N, 3°55′58″ E), was sandy-loam (62.6 % sand, 26.1 % silt, 11.3 % clay); with 1.36 g kg⁻¹ of total N, 16.9 g kg⁻¹ of total C, 0.069 g kg⁻¹ of phosphorus (P Olsen), pH in water was 7.06, and cation-exchange capacity (CEC) was 7.98 cmol_c kg⁻¹.

171 On 17-19 October 2016, 72 boxes were sown as monocultures (12 species x 6 replicates with one 172 replicate species per row). Seeds of each species were sown in lines, the distance between lines was 173 75 mm and the distance between plants within a line was 75 mm, leading to a final plant density of 174 155 plants m⁻². Once seeds had germinated, each box was inoculated with a purified solution of local 175 *Rhizobium* bacteria strains (Incolum Valorhiz[™], France) and was netted to avoid birds disturbing seeds. 176 During the experiment, mean air temperature was 13 C° (Figure S3) with a maximum of 30 °C and a 177 minimum of -0.4 °C (Figure S3a), and the cumulative precipitation was 349 mm (Fig S3 for additional 178 information on climatic conditions over the 37 weeks of the experiment). Soil temperature in the boxes followed closely the air temperature over the 37 weeks period, with a mean of 13.5 °C, a maximum of 25.9 °C and a minimum of 3 °C (Figure S3a). Air humidity ranged from 53 to 87%, with a mean value of 74% (Figure S3a) and solar irradiation ranged from 320 to 897 W m⁻² with a mean value of 568 W m⁻² (Figure S3b). During the experiment, boxes were carefully weeded by hand and plants were cut to ground level every 4 months to maintain a regular aerial cover. In addition, each box was watered with sprinklers when required.

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186 Analysis of carbon content in different soil fractions

187 Soil C content was measured before filling the boxes, as a reference for time 0 (t0), on three samples 188 from the initial homogenized soil batch, and at the end of the experiment, i.e. at 37 weeks (t37) after 189 sowing. At t37, soil samples were taken at 0-200 mm depth using a soil corer (75 mm in diameter) in 190 each box dedicated to soil and plant sampling. All soil samples were separated into two depths (0-100 191 mm and 100-200 mm), air dried and separately sieved to 2 mm. A subsample of 40 g of soil was 192 collected at a depth of 0-100 mm for subsequent fractioning into POM fractions (coarse POM: 200-193 2000 μ m and fine POM: 50-200 μ m), coarse silt (20-50 μ m), fine silt + clay (<20 μ m) fractions. Soil 194 fractionation was carried out using the method from Gavinelli et al. (1995). Soil samples were 195 presoaked overnight in 300 ml of deionized water at 4°C with 0.5 g of hexametaphosphate to enhance 196 disaggregation. Soil was then shaken at 300 rpm (digital orbital shaker, Intertek) with five agate 197 marbles for 2 h (i.e., the time suggested for sandy soils, to avoid the transfer of C into finer fractions, 198 Gavinelli et al. 1995). The soil was wet sieved with a 200 µm sieve, and the resulting 200-2000 µm 199 fraction was then transferred into a separate container and soaked in deionized water. The floating 200 coarse particulate organic matter (POM) was then carefully collected. The remaining 200-2000 μm 201 fraction represented the coarse sand fraction in soil and was carefully collected by washing the content 202 of the sieve in a beaker using deionized water. Then, the remaining fraction was sieved with a 50 μ m 203 sieve, to separate and collect the fine sand fraction and the fine POM fraction (50-200 μ m). The 204 remaining fraction <50 μ m was sonicated with a 1510E-MT Bransonic sonicator for 10 minutes to break 205 microaggregates before sieving at 20 µm. The 20-50 µm fraction (coarse silt) was collected and the 206 resulting solution of deionized water and <20 µm fraction collected in a beaker and filled up to 1.0 L. 207 This solution was tumbled 30 times to homogenize it and an aliquot of 100 mL was collected with the 208 aid of a syringe, representing the fine silt + clay fraction. All the fractions were oven dried at 40°C until 209 all the water evaporated. The dried fractions were weighed to check that the sum of the fraction's 210 weight did not differ from +/-5% the initial 40 g total weight. The quality of the soil particle dispersion was checked and did not differ +/-5% compared to the soil texture analysis, being 62.6% in the sand
fraction and 37.4% in the fine silt + clay fraction.

Carbon content in each of the soil fractions (C_{POM} - carbon in the coarse POM fraction; $C_{finePOM}$ – carbon in the fine POM fraction; C_{SILT} - carbon in the coarse silt fraction and $C_{SILT+CLAY}$ - carbon in the fine silt+clay fraction), was analyzed using an elemental analyzer (CHN model EA 1108; Carlo Erba Instruments, Milan, Italy) to assess the amount of C present in each pool. A subsample of 0.1 g was taken from each 40 g sample and analyzed without fractioning to determine the total C in the bulk sample. The difference between total C in bulk soil and the sum of C in the different soil fractions was used to assess the correctness of the fractionation (+/-5%) and was equal to 97.3% (SD=22%; n=34).

220 We calculated soil C changes (Δ C) in each soil fraction as the difference between C content in mg C g⁻¹

soil at 0 and 37 weeks ($\Delta C = C_{t37} - C_{t0}$). The sum of ΔC in each soil fraction ($\Delta C_{SUM} = \Delta C_{POM} + \Delta C_{finePOM} + \Delta C_{SILT} + \Delta C_{SILT+CLAY}$) was also calculated to investigate the variation in the totality of the soil fractions.

223 Note that ΔC can be either positive (accumulation) or negative (depletion due to the positive priming

effect, that is the increase of pre-existing soil C consumption and losses due to vegetation, Kuzyakov,

225 2002).

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227 Measurement of root elongation rate (RER) and root length production (RLP)

228 As soon as the first root of a given species was visible in a rhizotron, roots of that species were scanned 229 every 2 weeks for the experiment experiment (i.e. n = 19 measurement dates) using a smartphone 230 scanner application CamScanner (INTSIG Information Co., Ltd, Shanghai, China; version 3.9.5). A 231 smartphone (Samsung Xcover3, Samsung Electronics, Korea) was kept at a fixed distance from the 232 rhizotron (0.3 m) and a ruler was included in the picture to set the scale (Mohamed et al. 2017). Images 233 were then analyzed with the SmartRoot software (Lobet et al. 2011), a freeware plugin of ImageJ 234 software (Schneider et al. 2012). The images acquired were converted into 8-bit grey scale and, when 235 necessary, color-inverted, so that roots were dark on a lighter background. SmartRoot allows the semi-236 automatic tracing of roots by clicking on the basal point of each root (Fig. S4a). Data extracted include 237 the length and diameter of the roots. The resulting traced image of roots could then be imported and 238 superimposed onto a new image, allowing analysis of subsequent images and creating a time-239 dependent dataset acquiring root length at different time steps.

Root elongation rate (RER; mm d⁻¹) is indicative of primary root growth and is defined as the difference
in individual root length measured between two dates. RER is a frequent but punctual observation of
root dynamics over time. As soil C storage is a cumulative process, root length production (RLP; m)

243 after 37 weeks was also calculated for up to 60 roots (randomly chosen) per rhizotron. RLP is the total 244 length of all roots produced in a specific period of time (Mommer et al. 2015). Of these 60 roots 245 measured, 30 were selected from the 'new roots', i.e., the roots that were not present in the previous 246 scan, and so had formed in the previous two weeks. Of the 30 'new roots', 20 were primary and first 247 order roots and ten were second order laterals (Fig. S4a, according to the developmental centrifugal protocol of root topology; Berntson, 1997). Then, 30 'old roots' were selected at each subsequent 248 249 sampling date. The 'old roots' were the roots already present in the previous scan (again, 20 primary 250 axes and ten second order lateral roots). Fig. S5 shows an example of rhizotron analysis for new and 251 old roots. To have a more representative sample of the 'old roots,' ten primary roots were selected 252 from the 20 primary 'old roots' of the previous scan, ten were selected from the 20 newly emerged 253 roots of the previous scan, five were chosen from the ten second order lateral 'old roots' of the 254 previous scan and five were selected from the newly emerged second order laterals of the previous 255 scan. This method was used to select roots at each subsequent sampling date. If one or more roots 256 had: 1) reached the boundaries of the rhizotron, or 2) were in a bundle and not distinguishable (Fig. 257 S4b), or 3) could not be analyzed for any other reason (e.g. soil masking the root), they were discarded 258 and different roots were then selected.

The mean daily RER was calculated by subtracting from the length of a root (L_{t2}) the length of the root acquired at the previous sampling date (L_{t1}). This result was then standardized dividing by the number of days between the two sampling dates (t) to have the mean elongation rate of a single root:

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$$RER(t) = (L_{t2} - L_{t1})/[t2 - t1]$$

Root length production (RLP) of roots over the 37 weeks was chosen as a cumulative indicator for root
dynamics, adapted from Mommer et al. (2015):

265
$$RLP = \sum_{t=1}^{T} \sum_{r=1}^{R} (RER(t) * \frac{R_{30}}{R})$$
 Equation 2

266 Where *t* represents the sampling date ; $_{(t)}$; RER(t) is the daily RER; R the real number of roots analyzed 267 in that interval. Since the number of analyzed roots varied depending on dates and species, we decided 268 to standardize the analysis of RLP for R_{30} = 30 roots.

To refine the understanding of root dynamics, the RER and RLP were calculated separately for the new roots (RER_{NEW} and RLP_{NEW}, i.e. roots initiated during the 2 weeks interval between measurements), old roots (RER_{OLD} and RLP_{OLD}, i.e. roots older than 2 weeks), and also the total root system, regardless of root age. For all species, RER was high during the first two samplings after their initiation and then decreased rapidly or stopped. Therefore, mean RER could be biased by the development of new roots, justifying our decision to separate roots based on age and order for the statistical analysis.

Equation 1

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276 Analysis of root traits

After 37 weeks, a soil core (75 mm diameter, 200 mm depth), centered on one individual plant per species and per box was collected. In each core, roots were separated from the aboveground part and washed. Roots were sorted into absorptive roots, typically the first, second and third root orders (defined as the most distal root orders), and transport roots, that were higher order roots (all orders above third order roots, following McCormack et al. 2015). A subsample of absorptive roots (0.1 g dry mass on average) was selected, stained in a solution of methyl violet (0.5 g L⁻¹), spread into a transparent water filled tray and scanned at 800 dpi (Epson Expression 1680, Canada).

The software Winrhizo Pro (Regent Instruments, Quebec, Canada) was used to determine the root diameter (from 0 to 2 mm, with a 0.1 mm diameter interval) of absorptive roots. Roots were then oven dried at 40°C for 3 days and weighed to determine the total root dry mass for each core.

For each species, determination of root chemical composition was conducted on three subsamples of absorptive roots reserved for chemical analyses. C and N concentrations were determined on ground material using an elemental analyser (CHN model EA 1108; Carlo Erba Instruments, Milan, Italy). Concentrations of water-soluble compounds + hemicelluloses, cellulose and lignin were obtained following the Van Soest method (Van Soest, 1963) and using a fiber analyser (Fibersac 24; Ankom, Macedon, NJ, USA). This method consists of measuring the various plant tissue constituents by sequential extraction with neutral detergent, acid detergent and sulfuric acid (76%).

294 Substrate-induced respiration (SIR) was used as a proxy for potential soil microbial respiration and 295 activity, according to Beare et al. (1990). Briefly, 20 g air-dried 2 mm sieved soil samples were incubated in 150 mL sealed serum flasks with 1.5 mg C-glucose g⁻¹ soil, at 80 % field capacity and at 296 297 25°C. A 200 μ L aliquot of the flask headspace was analyzed for CO₂ concentration after 2 and 6 hours 298 using a microcatharometer (MicroGC Serie S, SRA Industries, Marcy l'Etoile, France), equipped with a 299 PoraPlot column (Agilent, Santa Clara, United States). Substrate induced respiration rates were 300 calculated as the mass of C-glucose converted to C-CO₂ per g of soil dry weight and per hour (in μ g C-301 $CO_2 g^{-1}$ soil h^{-1}).

302 Statistical analysis

First, a one-way analysis of variance (ANOVA) and post-hoc Tukey honestly significance difference (HSD) tests were performed to test the effects of species on mean RER, mean RLP, root traits and ΔC sequestration in soil C pools. Secondly, one way ANOVAs were conducted on the five Poaceae species and the five Fabaceae species, i.e excluding *P. lanceolata* and *S. minor* (hypothesis 3). In order to select 307 the environmental parameters to be included in the constrained ordination, an initial db-RDA including 308 all parameters was performed followed by a stepwise model selection using Generalized Akaike 309 Information Criterion (AIC, ordistep function with a backward direction). The normal distribution of 310 residuals was verified using a Shapiro-Wilk test (p = 0.05). If the data were not normally distributed, 311 one way ANOVA was substituted with a Kruskal-Walls test. Finally, the same procedure using one way 312 ANOVA was performed to compare the mean effect of N_2 -fixing and non N_2 -fixing species (for the 313 latter, grouping together Poaceae, P. lanceolata and S. minor) on root traits and C storage (hypothesis 314 3).

315 A principal component analysis (PCA) was performed on 12 variables (six root traits, four ΔC of each 316 soil C pool plus their sum, and SIR) using the mean for three replicate boxes (n = 12) to investigate the 317 effect of root traits at the species level. RER and RLP were not included in the PCA since they were 318 measured on one replicate box per species. Then, Pearson's correlation coefficients were calculated 319 to study the relationships between root traits and ΔC in each soil C pool (hypotheses 1 and 2) and 320 linear models of the significant correlations were analyzed to study the data dispersion. To deepen the 321 undesrtanding of these correlations, Pearson's correlation analysis and a study of the linear models 322 were performed on raw data (n = 34) to study relationships at the individual level.

To investigate the effect of abiotic factors on root growth dynamics, Pearson's correlation coefficients were calculated between mean daily RER, mean RLP, mean daily soil and air temperatures, mean daily soil humidity and mean daily solar irradiation (n = 12 for each variable). Means of daily climate data were calculated for the 2 weeks preceding the measurement of RER.

All the statistical analyses were performed in the open-source statistical environment R, version 3.4.3
(R Development Core Team, 2017) using the packages *Hmisc* (Harrel 2007) and *vegan* (Oksanen et al.
2019).

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331 RESULTS

332 Effect of plant species on soil carbon accumulation (ΔC) in different C pools associated with soil fractions

Plant species did not significantly influence the accumulation of C in different pools, nor in the sum of C pools (Fig. 1). The mean ΔC_{SUM} increase was $1.72 \pm 1.45 \text{ mg C g}^{-1}$ soil, and was highest in soil beneath *L. corniculatus* ($3.60 \pm 0.70 \text{ mg C g}^{-1}$ soil) compared to the bare soil control ($0.21 \pm 3.87 \text{ mg C g}^{-1}$ soil, Fig. 1a). The mean increase in the coarse pool ΔC_{POM} was $0.58 \pm 0.34 \text{ mg C g}^{-1}$ soil (Fig. 1b) and in the $\Delta C_{finePOM}$ was $1.21 \pm 0.74 \text{ mg C g}^{-1}$ soil (Fig. 1c). In the protected C_{SILT} pool, the ΔC mean increase was

338 0.57 ± 0.34 mg C g⁻¹ soil (Fig. 1d), while the $\Delta C_{SILT+CLAY}$ decreased by -0.50 ± 0.77 mg C g⁻¹ soil (Fig. 1e).

However, no significant differences were found between any species and bare soil with regard to any
C pool (Fig. S6, C data in different soil C pools for each species at t37).

341 Significant differences in mean ΔC between N₂-fixing Fabaceae and non N₂-fixing Poaceae were found 342 with regard to CPOM and CSILT. Mean CPOM was significantly higher in soil beneath Poaceae species 343 (ANOVA, p = 0.024, Tukey HSD test, Fig. 2a), whilst C_{SILT} was significantly higher in Fabaceae species 344 (ANOVA, p = 0.060, Tukey HSD test, Fig. 2b), and no significant differences were found in C_{SILT} between 345 Poaceae and bare soil. When grouping the data for all the non N_2 -fixing species (i.e., Poaceae, P. 346 *lanceolata* and *S. minor*), mean C_{POM} was higher compared to N₂-fixing Fabaceae (ANOVA, p = 0.06, F = 347 3.61) but C_{silt} was lower (ANOVA, p = 0.01, F = 7.01) (Fig. 2), although a Tukey HSD test did not find 348 significant differences between N₂-fixing and non N₂-fixing species.

349

350 Root elongation rate (RER) and root length production (RLP)

351 More than a threefold variation in mean daily RER_{TOT} occurred among species, ranging from 0.23 mm 352 d^{-1} (*F. rubra*) to 0.75 mm d^{-1} (*T. repens*) (Table 1). Mean daily RER_{TOT} did not differ between N₂-fixing 353 Fabaceae (0.57 \pm 0.08 mm d⁻¹ on average) and non N₂-fixing Poaceae (0.42 \pm 0.13 mm d⁻¹, ANOVA, p =354 0.221, Table 1), even when grouped with non N₂-fixing species (0.46±0.14 mm d⁻¹, ANOVA, p = 0.075). Mean daily RER_{TOT} peaked at 0.75 mm d⁻¹ in mid-February for Poaceae and then decreased, attaining a 355 356 value of 0.4 mm d⁻¹ from April to June 2017 (Figs. S7). For Fabaceae species, mean daily RER_{TOT} peaked 357 at 1.1 mm d⁻¹ in May 2017, before decreasing sharply in June 2017 (Fig. S7, mean RER_{TOT} for Fabaceae 358 and Poaceae species).

- 359 The mean daily RER for new roots (RER_{NEW}, 0.83 ± 0.22 mm d⁻¹) was significantly higher than that of old roots (RER_{OLD}, 0.17 ± 0.09 mm d⁻¹, ANOVA, p < 0.001). Mean daily RER_{NEW} ranged from 0.32 mm d⁻¹ (*F*. 360 rubra) to 1.13 mm d⁻¹ (*D. glomerata*) whereas RER_{OLD} ranged from 0.05 mm d⁻¹ (*P. pratensis*) to 0.40 361 362 mm d⁻¹ (*T. pratense*). Mean daily RER_{NEW} did not differ in N₂-fixing Fabaceae compared to non N₂-fixing 363 Poaceae or all non N₂-fixing species grouped together. However, mean daily RER_{OLD} was greater in N₂fixing Fabaceae (0.25 ± 0.09 mm d⁻¹) than in non N₂-fixing Poaceae (0.13 ± 0.03 mm d⁻¹, ANOVA, p =364 0.020) or all non N₂-fixing species grouped together (0.12 \pm 0.04 mm d⁻¹, ANOVA, p = 0.005, Table 1). 365 366 After 37 weeks, the highest cumulative RLP_{TOT} was observed in in *O. viciifolia* (3.62 m) and the lowest 367 in F. rubra (1.19 m) (Table 1). N₂-fixing Fabaceae species possessed a greater RLP_{TOT} (3.37 ± 2.32 m)
- 368 compared to non N₂-fixing Poaceae (2.32 \pm 0.70 m, ANOVA, p = 0.032), as well as all the N₂-fixing
- 369 species grouped together (2.42 ± 0.63 m, ANOVA, p = 0.009). Root dynamics of only three species were
- 370 correlated with climate factors. In *L. corniculatus,* mean daily RER_{TOT}, RER_{OLD}, RER_{NEW}, RLP_{TOT}, RLP_{OLD} and

RLP_{NEW} were all positively correlated with soil and air temperature and solar irradiation (Tables S1, S2).
In *T. repens*, RER_{TOT} and RER_{NEW}, RLP_{NEW} were significantly and positively correlated with soil and air
temperature (Tables S1, S2). With regard to Poaceae species, mean RER_{NEW} of *D. glomerata* was
negatively correlated with soil and air temperature and solar irradiation (Table S1). In *O. viciifolia*,
RLP_{TOT} was slightly and positively correlated with solar irradiation (Table S2). In *D. glomerata*, RLP_{NEW}
only, was negatively correlated with soil and air temperature (Table S2).

377

378 Root biomass, diameter and chemical composition

379 At 37 weeks, *M. sativa* had significantly greater mean root biomass $(4.23 \pm 0.42 \text{ g})$ compared to all 380 other species (Tukey HSD test, Table 1). In general, N₂-fixing Fabaceae species had a significantly higher 381 mean root biomass (2.08 \pm 1.33 g) compared to non N₂-fixing Poaceae (0.62 \pm 0.11 g) and all the non 382 N₂-fixing species grouped together (0.65 \pm 0.17 g, ANOVA, p < 0.001). The mean diameter of absorptive 383 roots differed significantly between species, with O. viciifolia having the thickest absorptive roots and 384 D. glomerata the thinnest (0.21 \pm 0.14 mm, Table 1). Species from the N₂-fixing Fabaceae family had 385 significantly thicker absorptive roots (0.39 ± 0.11 mm) compared to non N₂-fixing Poaceae (0.23 ± 0.03 386 mm) or all the non N₂-fixing species grouped together (0.25 \pm 0.03 mm, ANOVA, p < 0.001).

387 The chemical composition of absorptive roots strongly varied among species and between N_2 -fixing Fabaceae and non N₂-fixing Poaceae or all the non N₂-fixing species grouped together (Table 1). 388 389 Absorptive roots of N₂-fixing Fabaceae possessed more hemicelluloses + water-soluble compounds $(705 \pm 74 \text{ mg g}^{-1})$ than non N₂-fixing Poaceae (543 ± 33 mg g⁻¹) or all non N₂-fixing species grouped 390 391 together (583 \pm 69 mg g⁻¹), a lower mean lignin content (N₂-fixing Fabaceae: 173 \pm 56 mg g⁻¹, non N₂-392 fixing Poaceae: 302 ± 59 mg g⁻¹, all non N₂-fixing species grouped together: 264.18 ± 79.06 mg g⁻¹), and 393 a lower mean C:N ratio (N_2 -fixing Fabaceae: 19.15 ± 3.07, non N_2 -fixing Poaceae: 58.67 ± 6.34, and all 394 non N₂-fixing species grouped together: 62.04 ± 7.41). Mean root cellulose content did not differ either 395 among species or between N₂-fixing Fabaceae and non N₂-fixing Poaceae. However, when all the non 396 N₂-fixing species were grouped together, absorptive roots had a significantly higher mean cellulose 397 content compared to N₂-fixing Fabaceae (Table 1).

398

399 Soil substrate induced respiration (SIR)

400 Mean SIR for soil microbial communities varied significantly among species and between N₂-fixing 401 Fabaceae (5.28 ±1 μ g C-CO₂ g⁻¹ soil h⁻¹) and non N₂-fixing Poaceae (3.12 ± 0.41 μ g C-CO₂ g⁻¹ soil h⁻¹, 402 ANOVA, *p* < 0.001, Table 1). Mean SIR ranged from 2.47 ± 0.34 μ g C-CO₂ g⁻¹ soil h⁻¹ (beneath *B. erectus*) 403 to $6.41 \pm 0.56 \ \mu g \ C-CO_2 \ g^{-1}$ soil h⁻¹ (beneath *M. sativa*). When grouping all the non N₂-fixing species 404 together, mean SIR was still significantly lower ($3.1 \pm 0.46 \ \mu g \ C-CO_2 \ g^{-1}$ soil h⁻¹) compared to N₂-fixing 405 Fabaceae ($5.28 \pm 1.02 \ \mu g \ C-CO_2 \ g^{-1}$ soil h⁻¹, ANOVA, *p* < 0.001, Table 1).

406

407 Relationships between soil carbon accumulation (ΔC), root growth dynamics, root traits, and substrate 408 induced respiration (SIR)

409 The PCA conducted on the ΔC in the different C pools, SIR and root traits explained 64.6% of the 410 variance of the variables analyzed (Fig. 3). The first PCA axis (horizontal), accounting for 44.4% of the 411 variation, opposed ΔC_{POM} (negative) and ΔC_{SILT} (positive), while the remaining C pools, as well as the 412 sum of C pools, covaried and were quite orthogonal to ΔC_{POM} and ΔC_{SILT} and related to the second PCA 413 axis, that accounted for 20.2% of the variation. SIR and root biomass, diameter, and hemicelluloses + water soluble compounds content of absorptive roots all went along the 1st axis (positive) together 414 415 with ΔC_{SILT}. Root traits linked with recalcitrance, lignin, cellulose and C:N ratio, went along the 1st axis 416 (negative) together with ΔC_{POM} . Convex hull polygons reflecting intraspecific variations generally had 417 small areas and were segregated over the biplot (Fig. 3). The PCA strongly discriminated Poaceae from 418 Fabaceae. Poaceae were all on the negative end of the first axis and were characterized by high lignin 419 and cellulose contents, high C:N and accumulation of C in the coarse POM fraction. Fabaceae species 420 were at far right of the first axis and were characterized by a higher biomass and thicker roots that 421 were rich in hemicelluloses, favoring accumulation of C in the coarse silt fraction. The two other non 422 N₂-fixing species were situated in intermediate positions on the axis.

423 When analyzing the species effect of root traits on C storage (n = 12) regression analyses showed that 424 mean ΔC_{POM} was not related to RER, but was slightly significantly and negatively related to two traits: 425 diameter and hemicelluloses + water-soluble compounds content of absorptive roots (Table S3a, 426 Figure 4a,b). Mean ΔC_{SILT} was significantly and positively correlated with mean daily RER_{OLD and} RLP_{OLD} 427 and with the mean diameter of absorptive roots, root biomass, hemicellulose + water-soluble 428 compounds of absorptive roots, and SIR (Table S3s, Figures 5a,b,c,d,f, h), whereas mean ΔC_{SILT} was 429 negatively correlated with mean lignin and C:N ratio (Table S3s, Figures 5e,g). Linear regressions of 430 mean ΔC_{SILT} and C:N ratio show two segregated clusters of points: one with low C:N related to N₂-fixing 431 species and one with non N2-fixing species having a high C:N ratio and low accumulation in ΔC_{SILT} 432 (Figure 5g). Variations in mean ΔC_{SUM} , $\Delta C_{finePOM}$ and $\Delta C_{SILT+CLAY}$ were not explained by any variables. 433 Mean SIR was significantly and positively correlated to mean REROLD and RLPOLD, root biomass and 434 hemicelluloses + water-soluble compounds (Table S3, Fig. 6a,b,c,e), but negatively correlated with

435 mean lignin and C:N ratio (Table S3a, Figures 6d,f). Mean hemicelluloses + water-soluble compounds
436 were significantly and negatively correlated with mean lignin content and C:N ratio (Table S3a).

When considering Pearson's correlations at the individual level (n = 36), significant correlations were found only between absorptive root diameter and ΔC_{POM} , SIR and ΔC_{SILT} (Table S3b). Correlations between root traits and mean SIR were similar compared to correlatins of raw data (Table S3b). The data dispersion in linear models showed that at the individual level, even if R² was low, the tendency remained the same as that when mean data were used for ΔC_{POM} (Fig. S9), ΔC_{SILT} (Fig. S10), and SIR (Fig. S11).

443

444 DISCUSSION

445 Total C accumulation in soil did not differ among plant species (Fig. 1), but as expected, C accumulation 446 was significantly greater in the C_{SILT} pool beneath N₂-fixing Fabaceae, whereas in soil beneath non N₂-447 fixing Poaceae species, C accumulation was greater in the CPOM pool (Fig. 2). In line with our hypotheses, 448 the accumulation of C into different soil C fractions, specifically CPOM and CSILT, was correlated with root 449 traits (Fig. 4, 5, 6). The more rapid RER and greater RLP of older roots promoted C accumulation into 450 the C_{SILT} pool, but smaller root diameter and low content of labile compounds (hemicelluloses and 451 water soluble compounds) enhanced C accumulation into the CPOM pool. Although measuring total soil 452 organic carbon can be an easy method to evaluate C storage, it is not as sensitive to short-term C 453 dynamics or effect of plant species and families, as the C changes in different soil fractions. Studies of 454 C sequestration should therefore focus on better estimating C input into different C pools associated 455 with soil textural fractions (Wiesmeier et al. 2019).

456

457 Hypothesis 1: Root elongation rate and root length production are expected to favour carbon
 458 accumulation in the C_{SILT} and C_{SILT+CLAY} fractions

459 We hypothesized that a fast RER would promote C accumulation in coarse silt and fine silt+clay soil 460 fractions, through an increase in exudation and microbial activity along newly initiated roots. 461 Interestingly, REROLD and RLPOLD were significantly and positively correlated with soil microbial SIR and 462 ΔC_{SILT} (Fig. 5a,b, Table S3a), but not with the RER and RLP of newly initiated roots, that had very high 463 rates of growth. Dennis et al. (2010) hypothesized that rapidly elongating root tips grow quickly out of 464 the main zone of microbial activity, that is established once root exudates have been consumed. These 465 microbial communities then consume rhizodeposits from mucilage and cell senescence as well as 466 exudates from roots growing in proximity. Therefore, slow growing older roots would be maintained 467 in this zone of high microbial activity, and C accumulation in the coarse silt fraction would be higher, 468 especially in N₂-fixing species with populations of bacteria distributed in nodules all along roots. N₂-469 fixing Rhizobium bacteria also increase root elongation (Garrido-Oter et al. 2018), likely inducing a 470 feedback mechanism whereby a stimulated RER results in a higher exudation rate (Garcia et al. 2001), 471 acting as a substrate for newly colonizing Rhizobium communities. Although the role of microbial 472 communities is of utmost importance for C input into the soil, differences in the use of C within plants 473 could also explain the lack of a relationship between RER_{NEW}, RLP_{NEW}, SIR and C_{SILT}. In fast-growing, 474 newly initiated roots, we suggest that C in the form of non-structural carbohydrates (NSC, produced 475 during photosynthesis), will be used preferentially for cell production and expansion, as found in a 476 recent seasonal study of root elongation and NSC fluxes (Wang et al. 2018). In older roots with lower 477 RER, less NSC is required for growth, and excess NSC would be freely exuded, reflected in the high SIR 478 that we observed.

479 Contrary to our hypothesis, RER and RLP did not promote accumulation of C in the C_{SILT+CLAY} pool. 480 Surprisingly, the C_{SILT+CLAY} pool was the only pool where C was actually lost over the 37 weeks, in both 481 bare soil and beneath all plant species, and this mineralisation of C could not be explained by microbial 482 activity or by any root traits. When soil was prepared in our study, its excavation, crushing and sieving 483 would have disrupted soil aggregates (Franzluebbers, 1999). This increase in C mineralisation is higher 484 in clays, as organic matter that was highly protected within the clay fraction will be released during 485 disruption, providing a new pool of C available to microorganisms (Hassink, 1992). The presence of 486 plant roots can also lead to an increased mineralization and loss of preexistent soil C due to an 487 increased microbial activity (positive priming effect; Kuzyakov et al., 2000). In our study, the origin of 488 C was not assessed, so it was not possible to quantify any priming effect. The decrease in C was mainly 489 observed in the silt+clay fraction, challenging the assumption that the C pool associated to this soil 490 fraction has greater C stability (Torn et al., 2009). However, these results are in line with the findings 491 from Keiluweit et al. (2015) who observed a major priming effect in the organomineral associated C 492 after the input of oxalic acid, a common component of root exudates. The soil disturbance, together 493 with the input of highly degradable C, may well increase the release and priming of C in the 494 organomineral associated fraction, thereby decreasing the C content in the silt+clay fraction, but 495 further studies are required to better understand this phenomenon.

496

497 Hypothesis 2: more recalcitrant root traits are expected to favour the unprotected coarse POM fraction

Root traits linked to recalcitrance (high cellulose and lignin content and high C:N ratio) did not correlate
 to C accumulation in the coarse POM fraction (C_{POM}) (Fig. 4), but the PCA showed that this suite of traits

500 was sharing similar coordinates with C_{POM} on axis 1 (Fig. 3). This result is mainly due to the Poaceae 501 species that all have higher C accumulation in the C_{POM} pool, as well as recalcitrant root traits, 502 compared to species from other families. Recalcitrant compounds have all been reported to decrease 503 root decomposition rates (Silver and Miya, 2001; Aulen et al. 2012; Poirier et al. 2018). Lignin-504 carbohydrate complexes prevent polymer-hydrolyzing enzymes access to substrates, thus reducing the 505 degradability of plant organic matter (Cornu et al. 1994, Malherbe and Cloete, 2002). SIR, as a proxy 506 for microbial activity, was also strongly and negatively correlated to lignin content (Fig. 6f), probably 507 because lignin reduces the accessibility of polysaccharides to microorganisms through the formation 508 of links between lignin and polysaccharides (Bertrand et al. 2006). Products of lignin degradation can 509 also react with ammonia or amino acids to form further recalcitrant complexes that are less available 510 to microorganisms (Nömmik and Vahtras, 1982). The trend observed in the PCA (Fig. 3), that species 511 with recalcitrant tissues were linked to higher CPOM accumulation, is in contradiction with the lack of 512 significant correlations between ΔC_{POM} and lignin or C:N ratio. One reason for this lack of correlation 513 might be that the experiment was shorter than the root life span of some or all species, and full 514 accumulation of C in the CPOM pool had not yet occurred (Van der Krift et Berendse, 2002). Another 515 reason may be because CPOM was derived from the input of fresh C from plants, as well as losses of 516 older C that already existed in soil. While the accumulation of new C in this fraction is influenced by 517 the chemical composition of the root system, the losses of older, pre-existing C are not. A C labeling 518 approach would be helpful to assess the different fluxes of new and old C and to better explain the 519 relationships between root traits and C storage.

520

521 Interestingly, C accumulation in the C_{SILT} pool was negatively correlated with recalcitrant root traits (lignin and C:N ratio, Fig. 5e,g), but positively with hemicelluloses content and root diameter (Fig. 5d,f). 522 523 Hemicelluloses comprise polysaccharides soluble in alkali and are easily degradable by microorganisms 524 (Dekker, 1985). Hemicelluloses are usually produced to the detriment of lignin and enhance tissue 525 degradability through higher accessibility to amorphous phases in the lignocellulose structure (Malherbe and Cloete, 2002). Microorganisms will use this easily degradable C for growth and 526 527 respiration and then produce exudates and exopolysaccharides, that are used as a substrate for 528 subsequent microbial communities (Dennis et al. 2010). These exopolysaccharides and low molecular 529 weight compounds are belived to be the main precursors of C in the coarse silt pool (Simpson et al., 530 2007; Mambelli et al., 2011; Cotrufo et al. 2013, Vidal et al. 2018), probably explaining the high C_{SILT} 531 we found beneath N₂-fixing species (Fig. 2) with high hemicelluloses + water soluble compounds and 532 low lignin contents. Absorptive roots were negatively correlated with C accumulation in the CPOM pool 533 (Fig. 4a) and positively correlated with C accumulation in the C_{SILT} pool (Fig. 5d). Absorptive roots 534 generally have a higher turnover rate and undergo rapid transformation through microbial degradation 535 (McCormack et al., 2015), explaining the low accumulation in the C_{POM} pool and the positive correlation 536 with the C_{SILT} pool. However, this relationship may also be an artefact because Poaceae roots are 537 inherently very fine compared to Fabaceae roots (Roumet et al. 2006; 2016), highlighting that the 538 understanding of relationships between C accumulation and morphological traits is challenging 539 because of their inherent nature and plasticity (Poirier et al. 2018).

540

541 Hypothesis 3: Fabaceae and Poaceae strongly differ in their influence on accumulation of C into 542 different soil fractions

543 Contrary to that observed in previous studies (Binkley 2005; Fornara and Tilman, 2008; Plaza-Bonilla 544 et al., 2016; King et al., 2018), we did not find evidence of a greater accumulation of total C (ΔC_{SUM}) in 545 soil beneath N₂-fixing species, because variability was high within Fabaceae. However, we showed that 546 N₂-fixing species and non N₂-fixing species (especially Poaceae and Fabaceae) strongly differed in their 547 effect on the accumulation of C into different soil C pools. Roots of Poaceae, as compared to Fabaceae, 548 had a lower RER and RLP. Poaceae produced thinner roots, rich in lignin and cellulose with a high C:N 549 ratio. These more recalcitrant tissues slow down microbial activity and hence root decomposition rate 550 (Roumet et al. 2016; Freschet et al. 2017). Due to their particular chemical composition, non N₂-fixing 551 species, especially Poaceae species, promote C accumulation in the unprotected C_{POM} pool, and have 552 a lower C accumulation in the more stable C_{SILT} pool (Fig. 4). On the other hand, roots of N₂-fixing 553 Fabaceae grow faster and produce thick roots that are easily degradable, since they are rich in N (low 554 C:N ratio) and in hemicelluloses and water soluble compounds. These traits favour the development 555 of microbial biomass and enhance their activity, as observed from the SIR that was 40% higher beneath 556 Fabaceae than Poaceae species (see Fig. 7 for a conceptual model describing C accumulation in 557 different pools).

558 Because of the lack of correlations between root traits and C accumulation in C_{SILT+CLAY} pool, we cannot 559 establish that the C_{SILT+CLAY} pool increases when there is a higher input of labile C from N₂-fixing species. 560 C accumulation is the result of the input of new C and losses of pre-existing C that can be influenced 561 by the input of fresh C from plants (priming effect). The difference in behavior between these two 562 pools could result in poor correlations between root traits and C accumulation in the C_{SILT+CLAY} pool. An 563 isotopic approach differentiating between the changes in new C and old C in different C pools, allowing 564 us to assess the priming effect, would be fundamental to understanding the mechanisms behind soil 565 C storage (Rossi 2019).

567 CONCLUSION

568 Our findings show that specific plant root traits influence the accumulation of C into different pools, 569 largely through the mediation of microbial activity, shaping the C pathway in soil and, finally, its 570 persistence. Our results showed that non N₂-fixing Poaceae species, characterized by high contents of 571 lignin and cellulose and a high C:N, promoted accumulation of C in the unprotected coarse POM 572 fraction, while root traits associated with high labile C input (high hemicelluloses + water soluble 573 compound contents, high REROLD and RLPOLD) and microbial activity, typical of N2-fixing Fabaceae 574 species, stimulated C accumulation in the protected coarse silt fraction. Root elongation rate and 575 length production promoted microbial activity in older roots only, potentially suggesting either a 576 spatial influence of root exudate accessibility on microbial communities, or a relationship between 577 non-structural carbohydrate use in roots and available exudates for microbial consumption. The 578 planting of vegetation in bare soil also led to a loss of C in the fine silt+clay fraction, commonly belived 579 to be the most stable. Differentiating the source of C loss (pre-existing C in soil or fresh C from live 580 plants), is a fundamental step to assess the priming effect and understand the mechanisms behind C 581 loss in the finer soil fractions, and could be achieved through an isotope labelling approach. Longer 582 term studies on C dynamics are needed to understand these species and root trait effects over time 583 and the consequent C accumulation in different pools. Moreover, the influence of different soils and 584 associated microbial communities need to be taken into consideration for a broader understanding of 585 C pool dynamics. Our results will not only be useful for identifying plant species capable of enhancing 586 long-term C storage in soil, but will also contribute significantly to the understanding of mechanistic 587 processes within the C cycle.

588

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600 **REFERENCES**

- 601Aulen M, Shipley B, Bradley R (2012) Prediction of in situ root decomposition rates in an602interspecific context from chemical and morphological traits. Annals of Botany 109: 287-297
- 603Balesdent J and Balabane M (1996) Major contribution of roots to soil carbon storage inferred from604maize cultivated soils. Soil Biol Biochem 9: 1261–1263.
- 605Bardgett RD, Mommer L, Vries FT De (2014) Going underground : root traits as drivers of ecosystem606processes. Tends Ecol Evol 29: 692–699.
- Beare MH, Neely CL, Coleman DC, Hargrove WL (1990) A substrate-induced respiration (SIR)
 method for measurement of fungal and bacterial biomass on plant residues. Soil Biol Biochem
 22(5): 585-594
- 610 Berntson GM (1997) Topological scaling and plant root system architecture: developmental and 611 functional hierarchies. New Phytol 135: 621–634
- 612 Bertrand I, Chabbert B, Kurek B, Recous S (2006) Can the biochemical features and histology 613 ofwheat residues explain their decomposition in soil? Plant Soil 281:291–307
- 614 Besnard E, Chenu C, Balesdent J, Puget P, Arrouays D (1996) Fate of particulate organic matter in 615 soil aggregates during cultivation. Eur J Soil Sci 47: 495–503.
- Binkley D (2005) How Nitrogen-Fixing Trees Change Soil Carbon. In: Binkley D., Menyailo O. (eds)
 Tree Species Effects on Soils: Implications for Global Change. NATO Science Series IV: Earth Env Sci,
 vol 55. Springer, Dordrecht
- 619Bird JA, Kleber M, Torn MS (2008) C-13 and N-15 stabilization dynamics in soil organic matter620fractions during needle and fine root decomposition. Organic Geo- chemistry, 39, 465–477
- 621 Canarini A, Kaiser C, Merchant A, Richter A, Wanek W (2019) Root exudation of primary 622 metabolites: mechanisms and their roles in plant responses to environmental stimuli. Front Plant 623 Sci 10: 157
- 624 Chapin FS III (2003) Effects of plant traits on ecosystem and regional processes: a conceptual 625 framework for predicting the consequences of global change. Ann Bot 91: 455–463.
- 626 Cornu A, Besle JM, Mosoni P, Grent E (1994) Lignin- carbohydrate complexes in forages: Structure
 627 and consequences in the ruminal degradation of cell-wall carbohydrates. Reprod Nutr Dev 24:
 628 385–398
- 629 Cotrufo MF, Wallenstein MD, Boot CM, Denef K, Paul E (2013) The Microbial Efficiency-Matrix
 630 Stabilization (MEMS) framework integrates plant litter decomposition with soil organic matter
 631 stabilization: do labile plant inputs form stable soil organic matter? Global Change Biol 19: 988–
 632 995.
- 633 De Deyn GB, Cornelissen JHC, Bardgett RD (2008) Plant functional traits and soil carbon 634 sequestration in contrasting biomes. Ecol Lett 11: 516–531.
- 635Dekker M (1985) Biosynthesis and Biodegradation of Wood Components: Biodegradation of the636Hemicelluloses. Ed. T. Higuchi pp. 505 533: Academic Press
- Dennis PG, Miller AJ, Hirsch PR (2010) Are root exudates more important than other sources of
 rhizodeposits in structuring rhizosphere bacterial communities? FEMS Microbiol Ecol 72: 313–327.

- Derrien D, Barot S, Chenu C, Chevallier T, Freschet GT, Garnier P, Guenet B, Hedde M, Klumpp K,
 Lashermes G, et al. (2016) Stocker du C dans les sols : Quels mécanismes, quelles pratiques
 agricoles, quels indicateurs ? Étude et Gestion des Sols 23: 193–224.
- Downie JA (2010) The roles of extracellular proteins, polysaccharides and signals in the interactions
 of rhizobia with legume roots. Fed of Eur Microbiol Soc 34: 150–170.
- Dungait JA, Hopkins DW, Gregory AS, Whitmore AP (2012) Soil organic matter turnover is governed
 by accessibility not recalcitrance. Global Change Biol 18: 1781–1796
- Fehrmann RC and Weaver RW (1978) Scanning electron microscopy of *Rhizobium* sp. adhering to
 fine silt particles. Soil Sci Soc Am J 42: 279-281
- Fornara D, Tilman D (2008) Plant functional composition influences rates of soil carbon and
 nitrogen accumulation. J Ecol 96: 314–322.
- Franzluebbers AJ (1999) Potential C and N mineralization and microbial biomass from intact and
 increasingly disturbed soils of varying texture. Soil Biol Biochem 31: 1083-1090,
- Fujisaki K, Chapuis-Lardy L, Albrecht A, Razafimbelo T, Chotte JL, Chevallier T (2018) Data synthesis
 of carbon distribution in particle size fractions of tropical soils: implications for soil carbon storage
 potential in croplands. Geoderma 313, 41-51.
- Garcia JAL, Barbas C, Probanza A, Barrientos ML, Manero FJG (2001) Low molecular weight organic
 acids and fatty acids in root exudates of two *Lupinus* cultivars at flowering and fruiting stages.
 Phytochem Analysis 12: 305–311.
- Garrido-Oter R, Nakano RT, Dombrowski N, Ma KW, AgBiome Team, McHardy AC, Schulze-Lefert
 P (2018) Modular traits of the Rhizobiales root microbiota and their evolutionary relationship with
 symbiotic rhizobia. Cell Host Microbe 24(1): 155-167.
- Gavinelli E, Feller C, Larré-Larrouy MC, Bacye B, Djegui Z,Nzila JD (1995) Routine method to study
 soil organic matter by particle-size fractionation: examples for tropical soils. Comm Soil Sci
 Plan26(11&12): 1749-1760
- 664 Gleixner G, Poirier N, Bol R, Balesdent J (2002) Molecular dynamics of organic matter in a cultivated 665 soil. Org Geochem 33: 357-366.
- Griscom BW, Adams, J, Ellis PW, Houghton RA, Lomax G, Miteva DA, Schlesinger WH, Shoch D,
 Siikamäki J V, Smith P, Woodbury P, Zganjar C, Blackman A, Campari J, Conant RT, Delgado C, Elias
 P, Gopalakrishna T, Hamsik MR, Herrero M, Kiesecker J, Landis E, Laestadius L, Leavitt SM,
 Minnemeyer S, Polasky S, Potapov P, Putz FE, Sanderman J, Silvius M, Wollenberg E, Fargione J
 (2017) Natural climate solutions. Proceedings of the National Academy of Science of the United
 States of America. 114, 11645–11650
- Hamer U, Marschner B, (2005) Priming effects in soils after combined and repeated substrateadditions. Geoderma 128(1):38-51
- 674 Harrell FE (2007) Package 'Hmisc'. Harrell Miscellaneous
- Hassink J (1992) Effects of soil texture and structure on carbon and nitrogen mineralization in
 grassland soils. Biol Fertil Soils 14 : 126–134.
- Heinze, Gensch S, Weber E, Joshi J (2017) Soil temperature modifies effects of soil biota on plant
 growth. J Plant Ecol 10 (5): 808 821
- 679 Hernández MA, Romero J, Jaime C, León-pulido J (2017) Lignocellulosic biomass from fast-

- growing species in Colombia and their use as bioresources for biofuel production. Chem Eng
 Trans 58: 541–546.
- Holz M, Zarebanadkouki M, Kaestner A, Kuzyakov Y, Carminati A (2018) Rhizodeposition under
 drought is controlled by root growth rate and rhizosphere water content. Plant Soil 423: 429–442.
- Huck MG, Taylor HM (1982) The Rhizotron as a Tool for Root Research In Advances in Agronomy,
 ed. NC Brady, pp. 1-35: Academic Press
- Huo C, Luo Y, Cheng W, (2017) Rhizosphere priming effect: A meta-analysis. Soil Biology and
 Biochemistry 111: 78–84
- 588 Jones DL, Nguyen C, Finlay RD (2009) Carbon flow in the rhizosphere: carbon trading at the soil-589 root interface. Plant Soil 321: 5–33.
- Kätterer T, Bolinder MA, Andrén O, Kirchmann H, Menichetti L (2011) Roots contribute more to
 refractory soil organic matter than above-ground crop residues, as revealed by a long-term field
 experiment. Agric. Ecosyst. Environ. 141, 184–192
- Keiluweit M, Bougoure JJ, Nico PS, Pett-Ridge J, Weber PK, Kleber M (2015) Mineral protection of
 soil carbon counteracted by root exudates. Nature Climate Change 5: 588-595
- King AE, Blesh J, (2018) Crop rotations for increased soil carbon: Perenniality as a guiding principle.
 Ecol. Appl. 28, 249–261
- 697 King AE, Congreves KA, Deen B, Dun KE, Voroney RP, Wagner-riddle C (2019) Quantifying the 698 relationships between soil fraction mass , fraction carbon, and total soil carbon to assess 699 mechanisms of physical protection. Soil Biol Biochem 135: 95–107.
- Kleber M, Nico PS, Plante A, Filley T, Kramer M, Swanston C, Sollins P (2011) Old and stable soil
 organic matter is not necessarily chemically recalcitrant: Implications for modeling concepts and
 temperature sensitivity. Glob. Chang. Biol. 17, 1097–1107
- Kleber M, Eusterhues K, Keiluweit M, Mikutta C, Mikutta R, Nico PS (2015) Mineral-Organic
 Associations: Formation, Properties, and Relevance in Soil Environments, Advances in Agronomy.
 Elsevier Ltd.
- Kogel-Knabner I (2002) The macromolecular organic composition of plant and microbial residues
 as inputs to soil organic matter. Soil Biol Biochem 34: 139-162
- Kuzyakov Y, Friedel JK, Stahr K, (2000) Review of mechanisms and quantification of priming effects.
 Soil Biology and Biochemistry 32: 1485-1498
- Kuzyakov Y, Larinova AA (2005) Root and rhizomicrobial respiration: a review of approaches to
 estimate respiration by autotrophic and heterotrophic organisms in soil. J Plant Nutr Soil Sci 168:
 503 520
- Lal R (2004) Soil carbon sequestration impacts on global change and food security. Science 304:1623-1627.
- Larson JE, Funk JL (2016) Seedling root responses to soil moisture and the identification of a
 belowground trait spectrum across three growth forms. New Phytol 210: 827–838.
- Lavorel S, Diaz S, Cornelissen JHC, Garnier E, Harrison SP, McIntyre S, Juli G, Pérez-Harguinde PS,
 Roumet C, Urcelay C (2007) Plant functional types: are we getting any closer to the Holy Grail? In:
 Terrestrial Ecosystems in a Changing World, eds Canadell, J., Pitelka, L.F. & Pataki, D. Springer,
 Berlin, pp. 171–186.

- Lobet G, Pagès L, Draye X (2011) A novel image-analysis toolbox enabling quantitative analysis of
 root system architecture. Plant Physiol 157: 29-39
- von Lutzow M, Kogel-Knabner I, Ekschmitt K, Matzner E, Guggenberger G, Marschner B, Flessa H
 (2006) Stabilization of organic matter in temperate soils: mechanisms and their relevance under
 different soil conditions a review. Eur J Soil Sci 57: 426–445.
- Malherbe S, Cloete TE (2002) Lignocellulose biodegradation: fundamentals and applications. Rev
 Environ Sci Bio 1: 105–114,: 105–114.
- Mambelli S, Bird JA, Gleixner G, Dawson TE, Torn MS (2011) Relative contribution of foliar and fine
 root pine litter to the molecular composition of soil organic matter after in situ degradation.
 Organic Geochemistry, 42: 1099–1108
- McCormack LM, Dickie IA, Eissenstat DM, Fahey TJ, Fernandez CW, Guo D, Erik A, Iversen CM,
 Jackson RB.(2015) Redefining fine roots improves understanding ofbelow-ground contributions to
 terrestrial biosphere processes. New Phytol 207: 505–518.
- 734Mikutta R, Kleber M, Torn MS, Jahn R (2006) Stabilization of soil organic matter: association with735minerals or chemical recalcitrance?. Biogeochem 77: 25-56
- Minasny B, Malone BP, McBratney AB, Angers DA, Arrouays D, Chambers A, Chaplot V, Chen ZS,
 Cheng K, Das BS, Field DJ, Gimona A, Hedley CB, Hong SY, Mandal B, Marchant BP, Martin M,
 McConkey BG, Mulder VL, O'Rourke S, Richer-de-Forges AC, Odeh I, Padarian J, Paustian K, Pan G,
 Poggio L, Savin I (2017) Soil carbon 4 per mille. Geoderma 292, 59–86.
- Mohamed A, Monnier Y, Mao Z, Lobet G, Maeght JL, Ramel M, Stokes A (2017) An evaluation of
 inexpensive methods for root image acquisition when using rhizotrons. Plant Methods 13: 1–13.
- Mohamed A, Stokes A, Mao Z, Jourdan C, Sabatier S, Pailler F, Fourtier S, Dufour L, Monnier Y
 (2018) Linking above- and belowground phenology of hybrid walnut growing along a climatic
 gradient in temperate agroforestry systems. Plant and Soil 424: 103
- Mommer L, Padilla FM, van Ruijven J, de Caluwe H, Smit-Tiekstra A, Berendse F, de Kroon H (2015)
 Diversity effects on root length production and loss in an experimental grassland community. Funct
 Ecol 29: 1560–1568.
- Nguyen C (2003) Rhizodeposition of organic C by plants: mechanisms and controls. Agronomie 23,
 375–396.
- Nömmik H, Vahtras Z (1982) Retention and fixation of ammonium and ammonia in soils. In:
 Stevenson, F.J. (Ed.),. Nitrogen in agricultural soils. Agronomy monographs, No. 22. Agronomy
 Society of America, Madison, WI, USA.
- 753 Oksanen J, Blanchet FG, Friendly M, Kindt R, Legendre P, McGlinn D, Minchin PR, O'Hara RB, 754 Simpson GL, Solymos P, Henry M, Stevens H, Szoecs E, Wagner H (2019) Package 'vegan'. 755 1.17-Community Ecology Package, R package version 2. http://CRAN.R-756 project.org/package=vegan
- Paustian K, Lehmann J, Ogle S, Reay D, Robertso GP, Smith P (2016) Climate-smart soils. Nature
 532: 49–57.
- Plaza-Bonilla D, Nolot JM, Passot S, Raffaillac D, Justes E, (2016) Grain legume-based rotations
 managed under conventional tillage need cover crops to mitigate soil organic matter losses. Soil
 Tillage Res. 156, 33–43

- Poirier V, Roumet C, Munson AD (2018) The root of the matter: linking root traits and soil organic
 matter stabilization processes. Soil Biol Biochem 120: 246–259.
- R Development Core Team (2013) R: A Language and Environment for Statistical Computing.
 Vienna, Austria
- Rasse DP, Rumpel C, Dignac M (2005) Is soil carbon mostly root carbon ? Mechanisms for specific
 stabilization. Plant Soil 269: 341–356
- Robertson AD, Paustian K, Ogle S, Wallenstein MD, Lugato E, Cotrufo FM (2019) Unifying soil
 organic matter formation and persistence frameworks: the MEMS model. Biogeosci 16: 1225-1248
- Rossi LMW (2019) Embankment as a carbon sink: a study on carbon sequestration pathways and
 mechanisms in topsoil and exposed subsoil. Ph. D. Thesis, University of Montpellier, France
- Roumet C, Urcelay C, Díaz S, Roumet C (2006) Suites of root traits differ between annual and
 perennial species growing in the field. New Phytol 170 : 357–368
- Roumet C, Birouste M, Picon-Cochard C, Ghestem M, Osman N, Vrignon-Brenas S, Cao K-F, Stokes
 A (2016) Root structure-function relationships in 74 species: Evidence of a root economics
 spectrum related to carbon economy. New Phytol 210: 815–826.
- Rubino M, Dungait JAJ, Evershed RP et al. (2010) Carbon input belowground is the major C flux
 contributing to leaf litter mass loss: evidences from a (13)C labelled- leaf litter experiment. Soil
 Biology & Biochemistry, 42: 1009–1016
- Sanderman J, Maddern T, Baldock J, (2014) Similar composition but differential stability of mineral
 retained organic matter across four classes of clay minerals. Biogeochemistry 121: 409-424
- Sasse J, Martinoia E, Northen T (2018) Feed Your Friends: Do Plant Exudates Shape the Root
 Microbiome? Trends in Plants Science 23(1): 25-41
- Schmidt MWI, Torn MS, Abiven S, Dittmar T, Guggenberger G, Janssens IA, Kleber M, KögelKnabner I, Lehmann J, Manning DAC, Nannipieri P, Rasse DP, Weiner S, Trumbore SE (2011)
 Persistence of soil organic matter as an ecosystem property. Nature 478, 49–56.
- 787 Schneider CA, Rasband WS, Eliceiri KW (2012) NIH Image to ImageJ: 25 years of image analysis.
 788 Nature Methods 9: 671–675
- Shahzad T, Chenu C, Genet P, Barot S, Perveen N, Mougin C, and Fontaine S (2015) Contribution of
 exudates, arbuscular mycorrhizal fungi and litter depositions to the rhizosphere priming effect
 induced by grassland species. Soil Biology and Biochemistry, 80: 146–155
- Silver WL, Miya RK (2001) Global patterns in root decomposition: comparisons of climate and litter
 quality effects. Oecologia 129:407-419
- Simpson AJ, Simpson MJ, Smith E, Kelleher BP (2007) Microbially derived inputs to soil organic
 matter: are current estimates too low? Environmental Science & Technology, 41: 8070–8076
- 796Six J, Bossuyt H, Degryze S, Denef K (2002) A history of research on the link between797(micro)aggregates, soil biota, and soil organic matter dynamics. Soil Till Res 79: 7-31
- 798Sokol NW, Kuebbing SE, Karlsen-ayala E, Bradford MA (2019) Evidence for the primacy of living799root inputs, not root or shoot litter, in forming soil organic carbon. New Phytol 221: 233–246.
- Sollins P, Hofmann P and Caldwell BA (1996) Stabilization and destabilization of soil organic matter:
 mechanisms and controls. Geoderma 74: 65–105.

- 802Steinaker DF, Wilson SD, Peltzer DA (2010) Asynchronicity in root and shoot phenology in grasses803and woody plants. Global Change Biol 16: 2241–2251
- 804Torn MS, Swanston CW, Castanha C, Trumbore SE (2009) Storage and Turnover of Organic Matter805in Soil. In: Biophysico-Chemical Processes Involving Natural Non-living Organic Matter in806Environmental Systems, Publisher: John Wiley & Sons
- Van Der Krift TAJ, Berends F (2002) Root life spans of four grass species from habitats differing in
 nutrient availability. Functional Ecology 16(2): 198 203
- Van Soest PJ (1963) Use of detergents in the analysis of fibrous feeds. II. A rapid method for the
 determination of fiber and lignin. J AOAC Int 46: 829–835.
- Vidal A, Hirte J, Bender SF, Mayer J, Gattinger A, Höschen C, Schädler S, Iqbal TM, Mueller CW
 (2018) Linking 3D Soil Structure and Plant-Microbe-Soil Carbon Transfer in the Rhizosphere. Front
 Env Sci 6: 1–14.
- Wang Y, Mao Z, Bakker MR, Kim JH, Brancheriau L, Buatois B, Leclerc R, Selli L, Rey H, Jourdan C,
 Stokes A (2018) Linking conifer root growth and production to soil temperature and carbon supply
 in temperate forests. Plant Soil 426: 33–50
- Warembourg FR, Roumet C, Lafont F (2003) Differences in rhizosphere carbon-partitioning
 among plant species of different families. Plant Soil 256: 347–357.
- 819 Wiesmeier M, Urbanski L, Hobley E, Lang B, von Lützow A, Marin-Spiotta E, van Wesemael B, Rabot
 820 E, Ließ M, Garcia-Franco N, Wollschläger U, Vogel HJ, Kögel-Knabner I (2019) Soil organic carbon
 821 storage as a key function of soils A review of drivers and indicators at various scales. Geoderma
 822 333: 149-162,
- Žifc^{*}áková L, Ve^{*}trovský T, Howe A, Baldrian P (2016) Microbial activity in forest soil reflects the
 changes in ecosystem properties between summer and winter. Env Microbiol 18: 288-301
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TABLES AND FIGURES 827

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Table 1: Plant root traits and SIR for the 12 herbaceous species. Mean data ± standard deviation are given for species from N₂-fixing Fabaceae , N₂-fixing Poaceae families, and N₂-fixing species aggregated (Poaceae species, *P. lanceolate* 828 and S. minor). For each species and for N₂-fixing Fabaceae, N₂-fixing Poaceae families, and N₂-fixing species aggregated, mean and standard deviation are given. Abbreviations: RER_{TOT} – root elongation rate of the entire root system; 829 830 RER_{OLD} – of old roots older than 2 weeks; RER_{NEW} – of new roots younger than 2 weeks; RLP_{TOT} root length production of the entire root system; RLP_{OLD} – of old roots; RLP_{NEW} – of new roots; Root biomass – total root biomass of a core 831 sampled after 37 weeks; Diameter absorptive – mean diameter of absorptive roots after 37 weeks; Hemicell. + H₂O – concentration of hemicelluloses and water soluble compounds in absorptive roots; Cellulose, Lignin – concentrations 832 of cellulose and lignin in absorptive roots; C:N – ratio of carbon and nitrogen in absorptive roots; SIR – microbial subsrate induced respiration. Different letters next to the average value indicate statistically significant differences (p < 833 0.05) between species or families according to Tukey HSD tests. DF – degree of freedom (number of species - 1, number of observations). Statistically significant values (*p* < 0.05) are indicated in bold text.

Family Species	Lotus corniculatus	Trifolium repens	Fabaceae Trifolium pratense	Onobrychis viciifolia	Medicago sativa	Bromus erectus	Festuca rubra	Poaceae Dactylis glomerata	Poa pratensis	Lolium perenne	Rosaceae Sanguisorba minor	Plantaginaceae Plantago lanceolata	Effect of species (ANOVA)			N2.fixing (Fabaceae)	Non N2- fixing (Only Poaceae)	Effect of families (ANOVA) Poaceae VS Fabceae			Non N ₂ -fixing (All: Poaceae, S. minor, P. lanceolata)		Effect of N aquisition strategy (ANOVA) N2-fixing VS non N2fixing		
Acronym code	Lc	Tr	Тр	Ov	Ms	Be	Fr	Dg	Рр	Lp	Sm	Pl	Df	F	р		-	Df	F	р	-	Df	F	р	
RER _{TOT} (mm d ⁻¹)	0.55±1.1 (a)	0.75±1.5 (a)	0.59±1 (a)	0.53±10 (a)	0.57±0.9 (a)	0.42±0.6 (a)	0.23±0.50 (a)	0.58±0.9 (a)	0.66±2.90 (a)	0.42±0.8 (a)	0.39±0.5 (a)	0.5±0.7 (a)	11,175	18.9	0.061	0.57±0.08	0.42±0.13	2,9	1.8	0.22	0.46±0.14	1,10	3.94	0.075	
RER _{OLD} (mm d ⁻¹)	0.17±0.07 (a)	0.2±0.08 (a)	0.4±0.11 (a)	0.21±0.05 (a)	0.26±0.08 (a)	0.12±0.03 (a)	0.09±0.04 (a)	0.14±0.06 (a)	0.05±0.02 (a)	0.11±0.04 (a)	0.13±0.04 (a)	0.16±0.04 (a)	11,178	17.9	0.081	0.25±0.09	0.13±0.03	2,9	6.3	0.02	0.12±0.04	1,11	12.46	0.005	
RER _{NEW} (mm d ⁻¹)	0.99±0.27 (a)	0.91±0.21 (a)	0.73±0.15 (a)	0.81±0.16 (a)	0.97±0.12 (a)	1.06±0.29 (a)	0.32±0.08 (a)	1.13±0.25 (a)	0.96±0.31 (a)	0.67±0.11 (a)	0.66±0.08 (a)	0.76±0.11 (a)	11,178	17.4	0.11	0.88±0.11	0.79±0.33	2,9	0.37	0.7	0.79±0.28	1,12	0.425	0.529	
RLP _{TOT} (m)	3.03	3.26	3.36	3.62	3.61	2.55	1.19	2.76	2.16	2.95	2.31	3.04	-	-	-	3.37±2.32	2.32±0.7	2,9	5.17	0.03	2.42±0.63	1,13	10.05	0.009	
RLP _{OLD} (m)	4.51	4.91	4.26	4.96	5.23	3.47	1.05	3.3	2.76	4.29	3.32	4.5	-	-	-	1.33±0.61	0.61±0.2	2,9	7.5	0.01	0.65±0.23	1,14	16.05	0.003	
RLP _{NEW} (m)	0.89	1.05	1.78	1.35	1.6	0.78	0.53	0.81	0.32	0.62	0.47	0.99	-	-	-	4.78±2.97	2.97±1.21	2,9	5.12	0.03	3.24±1.14	1,15	8.181	0.02	
Root biomass (g)	1.53±2.6 (bc)	0.55±1.13 (c)	2.01±0.62 (b)	2.06±1.1 (b)	4.23±0.42 (a)	0.65±0.08 (c)	0.70±1.6 (c)	0.58±0.51 (c)	0.60±0.83 (c)	0.57±0.57 (c)	0.91±2.13 (c)	0.49±0.60 (c)	11,24	27.3	<0.001	2.08±1.33	0.62±0.11	1,28	10.9	<0.0011	0.65±0.17	1,31	19.3	<0.0011	
Diameter absorptive roots (mm)	0.47±0.05 (b)	0.28±0.05 (de)	0.32±0.04 (cd)	0.55±0.18 (a)	0.35±0.18 (c)	0.27±0.02 (def)	0.22±0.05 (fg)	0.21±0.14 (g)	0.22±0.19 (g)	0.22±0.05 (fg)	0.26±0.09 (efg)	0.28±0.14 (de)	11,23	98.7	<0.001	0.39±0.11	0.23±0.03	1,27	31.7	<0.001	0.25±0.03	1,31	33.5	<0.001	
Hemicell. +H ₂ O soluble compounds (mg g ⁻¹)	779.5±58.1 (a)	612±8.3 (bcde)	674.5±44.6 (abc)	704.1±78.8 (ab)	755.1±31 (a)	533.9±12.1 (de)	572.3±27.3 (cde)	530.4±57.6 (de)	562.4±31.2 (cde)	520.5±31.7 (e)	703.2±NA (ab)	639.4±12.7 (bcd)	11,21	17.4	<0.001	705.11±74.39	543.51±33.56	1,25	51.5	<0.001	583.86±69.51	1,31	23.11	<0.001	
Cellulose (mg g-1)	101.6±6.3 (a)	163.2±8.7 (a)	102.5±17.7 (a)	120±69.3 (a)	123.9±27.7 (a)	177.9±76.5 (a)	160.7±44.8 (a)	137.5±7.7 (a)	89.1±6 (a)	181.6±32.3 (a)	140±NA (a)	151.2±54 (a)	11,21	1.7	0.13	122.39±33.61	154.89±50.8	1,25	3.9	0.06	151.96±45.55	1,31	4.2	0.049	
Lignin (mg g-1)	118.9±57.8 (c)	224.8±0.4 (abc)	223±33.7 (abc)	175.9±9.4 (abc)	120.9±3.5 (c)	288.3±88.5 (ab)	267±39.5 (ab)	332.1±65.3 (a)	348.4±25.1 (a)	297.9±62.1 (ab)	156.8±NA (bc)	209.3±66.7 (abc)	11,21	6.7	<0.001	172.5±56.53	301.6±59.02	1,25	33.7	<0.001	264.18±79.06	1,31	13.63	<0.001	
C:N ratio	15.4±0.8 (c)	21.2±1.1 (c)	17.6±3.20 (c)	20.5±3.2 (c)	21.4±2 (c)	50.8±7.6 (b)	61.3±2.6 (ab)	62.5±6.4 (ab)	59.1±4.1 (ab)	61±4.4 (ab)	69.8±3.2 (a)	68.9±3.1 (a)	11,21	96.4	<0.001	19.15±3.07	58.67±6.34	1,25	436	<0.001	62.04±7.41	1,31	413.4	<0.001	
SIR (µg C-CO ₂ g ⁻¹ soil h ⁻¹)	5.37±0.46 (ab)	4.43±0.48 (bc)	6.11±0.35 (a)	4.07±0.18 (bcd)	6.41±0.56 (a)	2.47±0.34 (d)	3.22±0.11 (cd)	3.58±0.28 (cd)	3.15±0.15 (cd)	3.17±0.1 (cd)	3.16±0.23 (cd)	2.99±0.11 (cd)	9,20	16	<0.001	5.28±1	3.12±0.4	1,28	45.9	<0.001	3.1±0.46	1,31	60.05	<0.001	

- **Fig.1:** Comparison of the difference ($\Delta C = C_{t37} C_{t0}$) in carbon (C) after 37 weeks between different soil fractions for each species. a) total C (ΔC_{SUM}), b) C in the coarse POM fraction (ΔC_{POM}), c) C in the fine POM fraction ($\Delta C_{finePOM}$), d) C in
- the coarse silt fraction (ΔC_{SILT}) and e) C in the fine silt+clay fraction (ΔC_{SILT+CLAY}). In each boxplot, the lower edge of the box corresponds to the 25th percentile data point, while the top edge of the box corresponds to the 75th percentile

data point. The line within the box represents the median.



Fig.2: Comparison of the difference ($\Delta C = C_{t37} - C_{t0}$) in carbon (C) after 37 weeks among N₂-fixing Fabace, non N₂-fixing Poaceae species only and non N₂-fixing species aggregated (Poaceae, P. lanceolate, S. Minor), and control.a) C in the second seco



Fig.3: Principal Component Analysis of six soil variables (five carbon pool changes and SIR) and six root variables measured on 12 species. Black dots are Fabaceae, white dots are Poaceae, red dots are *Sanguisorba minor*, and green dots

849 are *Plantago lanceolata*. The Hull polygons unify the different replicates for the same species. Abbreviations: SIR – microbial subsrate induced respiration; Hemicelluloses + H₂0 – concentration of hemicelluloses and water soluble

850 compounds in absorptive roots, cellulose, lignin – concentrations of cellulose and lignin in absorptive roots; C:N – ratio of carbon and nitrogen in absorptive roots; ΔC_{POM} – difference ($\Delta C = C_{t37} - C_{t0}$) in carbon (C) after 37 weeks for the

851 coarse POM C pool; $\Delta C_{\text{finePOM}}$ – for the fine POM C pool; ΔC_{SILT} – for the silt C pool; $\Delta C_{\text{SILT+CLAY}}$ – for the silt + clay f C pool; ΔC_{SUM} – sum of different fractions, ΔC as the total change in C concentration in soil.



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e, red dots are *Sanguisorba minor*, and green dots oncentration of hemicelluloses and water soluble $C = C_{t37} - C_{t0}$ in carbon (C) after 37 weeks for the centration in soil.

- Fig. 4: Linear regression at species level (n=12) between ΔC_{POM} as the difference (ΔC = C_{t37} C_{t0}) in carbon (C) after 37 weeks for the coarse POM C pool and a) diameter of absorptive roots and b) hemicelluloses + water soluble compounds. 853
- The black symbols are the N₂-fixing Fabaceae species, the white symbols the non N₂-fixing Poaceae species, the red dots are *S. minor* and the green dots are *P. lanceolata*. The red line is the linear model function of the variables and R², 854
- F and p of the linear model are shown. 855



Fabaceae

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Fig. 5: Linear regression at species level (n=12) between ΔC_{sit} as difference ($\Delta C = C_{37} - C_0$) in carbon (C) after 37 weeks for the coarse silt C pool and a) root elongation rate of old roots (RER_{OD}), b) root length production of old roots (RLP_{OD}), c) root biomass, d) diameter of absorptive roots, e) lignin content, f) 858

hemicelluloses + water soluble compounds, g) CiN ratio and h) substrate induced respiration rate (SIR). The black symbols are the N₂ fixing Fabaceae 859

species, the white symbols the non Nz-fixing Poaceae species, the red dots are S. minor and the green dots are P. lanceolata. The red line is the linear model 860

function of the variables and \mathbb{R}^2 , F and p of the linear model are shown. 861







Fig. 6: Linear regression at species level (n=12) between substrate induced respiration rate (SIR) and a) root elongation rate of old roots (RER_{OD}), b) root length production of old roots (RLP_{OD}), c) root biomass, d) lignin content, e) hemicelluloses + water soluble compounds, f) C:N ratio. The black symbols are the N_{2} -fixing Fabaceae species, the white symbols the non N_{2} -fixing Poaceae species, the red dots *S. minor* and the green dots *P. lanceolata*. The red line is the linear model function of the variables and R^{2} , F and *p* of the linear model are shown.



Fig.7: Conceptual scheme of carbon (C) sequestration mechanisms into different soil pools under N₂-fixing (Pabaceae) and non N₂-fixing (Poaceae) species. Square boxes refer to the major factors affecting C sequestration. Ellipses show 870

- the destination of C into unprotected pools (POM and finePOM) and protected pools (coarse silt and fine silt+clay). Text in the central column describes the mechanisms favoring C sequestration into soil C pools. The arrows symbolize 871
- the processes of transformation or transport of C into different pools. Arrow colors represent the nature of the C: red C derived from root turnover, green C from rhizodeposition, vellow C from microbial exudates and exopolysaccharides 872
- 873 and blue the C respired back into the atmosphere as CO₂. The thickness of the arrows is qualitative, with wider arrows reflecting higher C fluxes. The signs: "+" (in green) means an increase and "-" (in red) means a decrease.
 - N₂-Fixing **Non N2-Fixing** (Fabaceae) (Poaceae) **CO**2 Mechanisms favouring soil C **CO**2 sequestration in different C pools Recalcitrant C input (high C:N ratio lignin and cellulose content) decreases microbial **Root traits Root traits** activity and degradation Root biomass Root biomass Hemicellulose Hemicellulose 🕂 Lignin Lignin + Cellulose Cellulose **Root dynamics** Root elongation rate 🕂 C:N C:N High labile C input increases **Root dynamics** Root length production microbial activity and the Root elongation rate consequent exudates and Root turnove Root length production exopolysaccharides production Soil microbial Soil microbial communities commuties - SIR - SIR Lower microbe induced decomposition increases the residence time of particulate organic matter and Microbiological Microbiological the overall C content in the unprotected effect) exudates and exudates and exopolysaccharides coarse and fine POM pools exopolysaccharides Protected C Unprotected C Unprotected C Protected C Higher deposition of labile compounds increases the 题 题 A A A C storage in the protected finePOM POM Fine silt+clay silt pool and decreases the finePOM Fine silt+clay Coarse silt losses in the silt+clay pool



