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The simple AMG model accurately simulates organic carbon storage in soils after repeated application of exogenous organic matter 5 Florent Levavasseur^{1*}, Bruno Mary², Bent T Christensen³, Annie Duparque⁴, Fabien 6 Ferchaud², Thomas Kätterer⁵, Hélène Lagrange⁶, Denis Montenach⁷, Camille Resseguier¹, 7 Sabine Houot¹ UMR INRA AgroParisTech ECOSYS, Université Paris-Saclay, 78850 Thiverval-Grignon, 9 France ² BioEcoAgro Joint Research Unit, INRAE, Université de Liège, Université de Lille, Université de Picardie Jules Verne, 02000 Barenton-Bugny, France 12 3Department of Agroecology, AU Foulum, Aarhus University, Tjele, Denmark ⁴ Agro-Transfert Ressources et Territoires, 80200 Estrées-Mons, France ⁵ Department of Ecology, Swedish University of Agricultural Sciences, Uppsala, Sweden ⁶ ARVALIS-Institut du Végétal, 31450 Baziège, France ⁷UE INRA 0871 Service d'expérimentation Agronomique et Viticole, Colmar, France ***** Corresponding author: florent.levavasseur@inra.fr Acknowledgments This work was performed in partnership with the SAS PIVERT (www.institut-pivert.com). It was supported by the French Government (ANR-001-01) and the Genesys WP1 P13 Solebiom project. The QualiAgro and Colmar field experiments are part of the SOERE-PRO (a network of long-

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Abstract

 Repeated application of exogenous organic matter (EOM) contributes to soil organic carbon (SOC) stocks in cropped soils. Simple and robust models such as the AMG model are useful tools for predicting the effects of various EOM practices on SOC. In AMG, EOM is characterized by a single parameter: the isohumic coefficient *K1,* which represents the proportion of carbon that is humified and incorporated into SOC. The AMG model has been 42 validated under various pedoclimatic conditions and cropping systems, but has not yet been tested with data from long-term field experiments where EOM is regularly applied. The calibration of the EOM parameter *K¹* also remains an issue. In this study, AMG was used to simulate SOC stocks in seven long-term field experiments with EOM application. AMG 46 simulated changes in SOC stocks with a mean *RMSE* of 3.0 t C ha⁻¹ (the difference in SOC stocks between treatments with and without EOM). The optimized *K¹* values were highly correlated (*R²*=0.62) with the indicator of remaining organic carbon (*IROC*), which is measured by laboratory analysis. The present study (i) demonstrated the ability of the AMG model to accurately simulate changes in SOC stocks in long-term field experiments with regular EOM 51 application and (ii) validated the parameterization of EOM in AMG using I_{ROC} , which is routinely measured by commercial laboratories. Twenty-six different EOM types representing a wide range of EOM sources were parameterized using more than 600 *IROC* values. The AMG model could thus be used to predict the SOC increase following EOM addition with a very simple calibration.

 Keywords: organic amendment, organic fertilizer, EOM, soil organic carbon stock, model, AMG.

1 Introduction

 Exogenous organic matter (EOM) is organic matter that is not directly derived from crops but applied to soil as an organic fertilizer or organic amendment to improve soil fertility. EOM

 includes animal manures and other organic matter from urban or industrial activities, often treated to produce compost and anaerobic digestates. Depending on its characteristics, the use of EOM may increase the short-term and long-term nutrient supply (Gómez-Muñoz et al., 2017), increase soil organic matter (SOM) stocks (Zavattaro et al., 2017) and improve soil quality (Eden et al., 2017; Obriot et al., 2016). An increase in soil organic carbon (SOC) 66 stocks also contributes to reduced atmospheric $CO₂$ concentration (Maillard and Angers, 2014) provided that EOM had not been applied to the soil before (Powlson et al., 2011). Repeated EOM application in field experiments increase SOM stocks, depending on EOM type and characteristics, and the amount and frequency of EOM application (Bhogal et al., 2018; Maltas et al., 2018). The physico-chemical characteristics of applied EOM, such as nitrogen and lignin content, affect their initial rate of decomposition, whereas their influence in the longer term is uncertain (Dignac et al., 2017). Additionally, specific site conditions (e.g., soil texture and climate) influence the potential soil carbon storage.

 Models simulating soil carbon turnover are useful for predicting and comparing different EOM application practices and their efficiency in increasing SOM stocks. The simpler these models are, the more likely they are to be used by farming advisors. Several models have been developed over the past decades, including, the Hénin-Dupuis (Hénin and Dupuis, 1945), Roth-C (Jenkinson and Rayner, 1977), Century (Parton et al., 1987), ICBM (Andrén and Kätterer, 1997) and C-TOOL (Taghizadeh-Toosi et al., 2014) models. The AMG model (Andriulo et al., 1999) was developed in the late 1990s to improve the Hénin-Dupuis model. Compared to the Roth-C and Century models, AMG uses a simpler representation of organic 82 matter with only three pools: organic inputs (crop residues, roots or EOMs); an active SOM pool, which is supplied by humified organic inputs and undergoes mineralization; and a stable SOM pool, taken to be inert at the considered timescale (decades). AMG has been parameterized and validated for a wide range of cropping systems and pedoclimatic conditions (Bouthier et al., 2014; Clivot et al., 2019; Saffih-Hdadi and Mary, 2008). It is widely

 used in France by farming advisors, researchers, teachers and students as it has a 88 dedicated online version [\(http://www.simeos-amg.org/\)](http://www.simeos-amg.org/). However, its ability to simulate the effects of repeated EOM application on the long-term evolution of SOC stocks has not yet been verified. An update of the EOM parameters is also needed due to the recent development of the model (AMGv2). This version involves new formalisms and changes in some parameter values (Clivot et al., 2019). Moreover, because of the required increase in organic waste recycling in Europe (Directive 2018/851 of the European Parliament and of the Council) and in anaerobic digestion of both urban and agricultural wastes, the use of certain types of EOM is increasing together with the diversity of available EOM. Therefore, parameters for a wider range of EOM types are needed to facilitate AMG use by both researchers and stakeholders. The parameterization of EOM has previously been linked to the Roth-C model using field data from long-term experiments (Dechow et al., 2019; Peltre et al., 2012). However, long-term field trials are scarce and other methods based on laboratory 100 analysis are needed to parameterize AMG. The indicator of residual organic carbon (*I_{ROC}*) (Lashermes et al., 2009) is calculated from biochemical fractions of EOM (Van Soest and Wine, 1967) and the proportion of EOM carbon mineralized during 3 days of incubation with soil. *IROC* has been defined as a predictor of the EOM residual carbon after long-term 104 incubation of EOM with soil under controlled conditions. Peltre et al. (2012) used *I_{ROC}* to predict the partition of EOM into the different carbon pools of the Roth-C model. Preliminary results also indicate its potential use for parameterizing EOM in AMG (Bouthier et al., 2014).

 The objectives of this study were (i) to evaluate the ability of the AMG model to accurately simulate SOC stock evolution in long-term field experiments where EOM is regularly applied, (ii) to validate a method for EOM parameterization in AMG from laboratory analysis, and (iii) to use this method to parameterize a wide range of EOM types in AMG.

2 Materials and Methods

2.1 Field data

 Data from seven long-term field experiments were used for this study: Askov K2 in Denmark (Bruun et al., 2003; Christensen and Johnston, 1997), Colmar in eastern France (Obriot, 2016), QualiAgro in northern France (Obriot et al., 2016; Peltre et al., 2012), La Jaillière in western France (Bouthier and Trochard, 2015), SERAIL in southern France (Peltre et al., 2012), Broadbalk at Rothamsted Research in southern England (Jenkinson and Rayner, 1977; Perryman et al., 2018) and Ultuna in central Sweden (Gerzabek et al., 1997; Karhu et al., 2012; Kätterer et al., 2011). These seven experiments cover a wide range of durations, climates, soils, crop rotations and EOM types (Tables 1 and 2). More details about each experiment can be found in the references given above.

 Carbon inputs from plants were divided into aboveground and belowground inputs. Belowground carbon inputs (roots and exudates) were computed for each treatment and year according to the measured crop yields and allometric relationships (Clivot et al., 2019) adapted from Bolinder et al. (2007). For aboveground carbon, allometric relationships were also used for the Askov and SERAIL experiments, whereas carbon inputs were directly estimated from biomass measurements in the field for the QualiAgro, Colmar, La Jaillière, Ultuna and Rothamsted experiments.

 The EOM carbon inputs were computed from the known amount of applied EOM and the dry matter and carbon content of the EOM. When these variables were not determined every year, mean values of the available data were used to estimate carbon inputs from EOM.

 The SOC stocks were calculated by using SOC content and soil mass over the sampling depth, by considering an equivalent soil mass (Ellert and Bettany, 1995). When the soil bulk density changed over time (QualiAgro, Ultuna and Rothamsted experiments), SOC stocks were estimated by adding a variable amount of subsoil. In the Askov, Colmar, La Jaillière

 and SERAIL experiments, bulk density remained constant and SOC stocks were computed with a constant soil depth. If any rock fragments were present (as at La Jaillière), the bulk density of the fine soil was used, and the rock fragment fraction was subtracted from the soil volume to compute the SOC stocks, as suggested by Poeplau et al. (2017).

2.2 AMG model

 The AMG model originally developed by Andriulo et al. (1999) simulates the dynamics of SOC stocks at an annual time step in response to climate and cropping systems. A full description of the AMGv2 model version used in this study is given by Clivot et al. (2019) along with its validation under various pedoclimatic and cropping conditions. Only the main principles are described below.

 AMG considers three carbon pools [\(Figure](#page-32-0) **1**): (i) a pool including carbon inputs from crop residues, roots and EOM, (ii) an active carbon pool and (iii) a stable carbon pool. A fixed proportion, *K1,* of the carbon inputs is humified yearly and allocated to the active carbon pool. 149 The remaining 1- K_1 is mineralized as CO_2 . Aboveground crop residues, roots, EOM are each characterized by the specific *K¹* parameter (also termed isohumic coefficient). The active carbon pool decomposes according to first-order kinetics with a rate constant *k* affected by mean annual water balance and air temperature, soil clay and carbonate contents, soil pH and the C to N ratio of SOM (Clivot et al., 2017). The stable carbon pool is taken to be inert during the simulated period. The initial proportion of stable carbon is preset by default to 65% of the total SOC in long-term arable soils. The aboveground and belowground carbon inputs from plants are estimated by allometric relationships (Clivot et al., 2019) adapted from Bolinder et al. (2007) or directly as measured field data if available. The stocks are computed for a soil depth given as input data. AMG can be described by the following set of equations (Clivot et al., 2019):

$$
160 \t\t QC = QCS + QCA
$$
\t(1)

161
$$
\frac{dQC_A}{dt} = \sum_i m i K_{1i} - k QC_A
$$
 (2)

162 where *QC* is the total SOC stock (t ha^{−1}), *QC_A* and *QC_S* are the C stocks of the active and 163 stable C pools (t ha⁻¹), respectively, m_{*i*} is the annual C input from organic residue i (t ha⁻¹ 164 yr⁻¹), *K_{1i}* is its isohumic coefficient (the fraction of C inputs which is incorporated in SOM after 165 1 year) and *k* is the mineralization rate constant of the active C pool (yr⁻¹)

2.3 Determination of the coefficient *K¹* **in long-term field experiments**

2.3.1 EOM K¹ optimization

 The *K¹* value of each EOM type was estimated by fitting the model to the time series of differences in SOC stocks between the EOM treatments and the control treatments, i.e., the additional SOC storage in the treatments with EOM, designated ∆SOC stocks. The optimization was performed in R with the "optim" function and the "L-BFGS-B" method (Byrd et al., 1995) by minimizing the sum of squared differences between the observed and simulated ∆SOC stocks, with *K¹* varying between 0 and 100%.

 Using ∆SOC stocks instead of SOC stocks prevents the effects of a poor parameterization of the initial proportion of stable soil organic carbon, as suggested by Peltre et al. (2012). It also reduces the risk of bias, since any bias that occurred in simulating the SOC stocks of the control treatment is expected to also occur when simulating the EOM treatment. Beyond EOM inputs, we also considered the differences in crop residue input among treatments to explain ∆SOC stocks, since the carbon input from crop residues was input data specific to each treatment.

 All other parameters of the model were kept as the default values to preserve the consistency of the model and the general applicability of the optimized EOM *K¹* values. This included the *K¹* values of aboveground crop residues and roots (Table S3 in Clivot et al.,

184 2019). Finally, for the field experiments in which soil pH and soil C to N ratio changed over 185 time, these changes were considered in the model input data.

186 *2.3.2 Model evaluation*

 The model performance was first evaluated in control treatments without EOM (ASK_CT+N, COL_CT+N, QUA_CT-N and QUA_CT+N, LAJA_CT+N, ROTH_CT_S0 and ROTH_CT_S1, SER_CT_eqC and SER_CT_eqH, ULT_CT-N and ULT_CT+N). Indeed, because organic carbon from EOM is incorporated in SOC, it was important to verify the ability of the model to accurately simulate the evolution of SOC stocks under the various soil, climate and cropping system conditions. The mean error (*ME*), root mean square error (*RMSE*), relative *RMSE* (*RRMSE*), coefficient of determination (*R²*) and model efficiency (*EF*) were thus computed using the observed and simulated SOC stocks in each control treatment (Equations 3 to 7). The model performance for each treatment with EOM was then assessed by using the same statistical criteria.

197
$$
ME = \frac{1}{n} \cdot \sum_{i=1}^{n} (O_i - S_i)
$$
 (3)

198
$$
RMSE = \sqrt{\frac{1}{n} \cdot \sum_{i=1}^{n} (O_i - S_i)^2}
$$
 (4)

$$
199 \quad RRMSE = RMSE/\overline{O} \tag{5}
$$

200
$$
R^{2} = \left(\frac{\sum_{i=1}^{n} ((o_{i} - \bar{o}) \cdot (s_{i} - \bar{s}))}{\sqrt{\sum_{i=1}^{n} (o_{i} - \bar{o})^{2}} \cdot \sqrt{\sum_{i=1}^{n} (s_{i} - \bar{s})^{2}}}\right)^{2}
$$
(6)

201
$$
EF = 1 - \frac{\sum_{i=1}^{n} (o_i - s_i)^2}{\sum_{i=1}^{n} (o_i - \bar{o})^2}
$$
 (7)

202 where *O* and *S* are the observed and simulated values, respectively, and *n* is the number of 203 observations. \bar{O} and \bar{S} are the means of the observations and simulations, respectively.

2.4 Using *IROC* **to model SOC stocks**

205 The indicator of residual organic carbon (*I_{ROC}*) is computed from the biochemical fractions of EOM (Van Soest and Wine, 1967) and from the proportion of organic carbon mineralized 207 during 3 days of soil incubation (Lashermes et al., 2009). *I_{ROC}* was defined as a predictor of the EOM residual carbon after long-term incubation of EOM with soil under controlled conditions.

 The *IROC* values of the EOM types applied in the field experiments were available only for 4 211 out of the 7 experiments: Colmar, QualiAgro (one *I_{ROC}* value per EOM type and per application year), Ultuna (three *IROC* values per EOM type corresponding to three pooled 213 samples of different years of application) and SERAIL (one or two *I_{ROC}* values per EOM type 214 corresponding to one or two application years). When several *I_{ROC}* values were available for 215 a given EOM type in a field experiment, the mean of the I_{ROC} values was used as the K_1 value in AMG. For the 4 experiments with measured *IROC* values, the relationship between 217 the optimized K_1 values based on field data (section [2.3\)](#page-8-0) and the I_{ROC} values of the EOM was investigated by calculating the coefficient of determination (*R²*), mean error (*ME*) and root 219 mean square error (*RMSE*). The I_{ROC} values of the EOM were also directly used as $K₁$ input values in AMG (without any rescaling). The model performance was assessed under the same conditions as previously described (section [2.3.2\)](#page-9-0) and compared to the model performance with optimized *K1*.

2.5 The *IROC* **database**

 The original database developed by Lashermes et al. (2009) provided *K¹* values for EOM types for which we did not have any field estimates. This database was expanded with 226 additional *I_{ROC}* values from various research projects. Thus, the database contained more than 600 EOM types from 5 main groups: livestock liquid and solid manures (n=106), anaerobic digestates (n=54), sewage sludge (n=71), composts (n=352) and others (n=69). Most frequently used EOM types were present in the database; however, there were limited

 number of values for cattle slurry (n=5) considering its relative abundance and for different types of digestates (from 2 to 9 depending on the type of digestates).

3 Results

3.1 Evaluation of the AMG model for the control treatments

 The performance of the AMG model for the control treatments without EOM application was good for the Colmar, QualiAgro, La Jaillière, SERAIL and Ultuna experiments. In these 236 treatments, the *RMSE* was equal to or lower than 3 t C ha⁻¹ and the *RRMSE* was lower than 10% (electronic supplementary material, appendices A and B). The model thus succeeded in 238 representing the decrease in SOC stocks at La Jaillière (CT_N), Ultuna (CT_N and CT), SERAIL (CT_EQC and CT_EQH), and QualiAgro (CT_N-), and the nearly constant SOC 240 stocks at Colmar (CT_N) and QualiAgro (CT_N+). For Askov and Rothamsted, the model 241 performance was worse, with an *RMSE* between 4 and 5 t C ha⁻¹ in all control treatments (*RRMSE* from 19% to 30%). The model failed to simulate the increase in SOC stocks at Askov (CT_N) and overestimated the decrease in SOC stocks at Rothamsted (CT_S0 and CT_S1). If we consider all treatments and experiments, the simulation bias (*ME*) remained 245 low, smaller than 1 t C ha⁻¹, except for the Askov, Rothamsted, Ultuna (T0) and La Jaillière experiments.

3.2 Simulation of ∆SOC stocks with the optimized EOM *K¹* **value**

 Using the optimized values of *K1*, AMG succeeded in simulating the dynamics of ∆SOC 249 stocks in all treatments and experiments, except in two treatments at QualiAgro (BIO_N+, BIO_N-) and one at SERAIL (GW_EQH) which were underestimated [\(Figure](#page-32-1) **2**). This was confirmed by the low *ME* values [\(Table](#page-30-0) **3**), except those for the underestimated treatments (*ME* higher than 3.5 t C/ha). The *RMSE* remained low, with minimum, mean and maximum 253 values of 0.5, 3.0 and 8.2 t C ha⁻¹ respectively. Except for two treatments at La Jaillière

 (POM, CPOM), which had very low variability in ∆SOC stocks, and the SERAIL experiment, for which the dynamics of ∆SOC stocks were very noisy, the R² values were high (>0.6) to very high, indicating a good simulation of the variability of ∆SOC stocks. For example, the initial increase followed by stability in ∆SOC stocks at Rothamsted was well simulated, as was the initial increase at La Jaillère followed by a decrease in ∆SOC stocks after the cessation of EOM application. The model efficiency (*EF)* exhibited the same trends as *R²*.

3.3 Determination of isohumic coefficient (*K1)* **values from long-term field experiments**

 The optimized *K¹* values varied greatly among EOM types. For example, the lowest value was 8% for the poultry manure at La Jaillière, while values of 100% were found for several EOM types including green waste and sludge compost at QualiAgro [\(Table](#page-30-0) **3**). The mean optimized *K¹* for all treatments was 63%. For a given type of EOM, a large variation also 266 existed. For example, the optimized $K₁$ value of cattle manure varied from 44% (FYM treatment at Ultuna) to 99% (FYM_N+ at QualiAgro), with a mean value of 66%.

268 **3.4 Prediction of** K_1 **with** I_{ROC}

 Based on the four experiments for which *IROC* values of EOM were available, the *IROC* indicator appeared to be an acceptable predictor of *K¹* [\(Figure 3\)](#page-32-2). *ME* and *RMSE* between *K¹* and *IROC* were low for SERAIL and Ultuna (e.g., *RMSE* equal to 11% for Ultuna) but higher for 272 Colmar and QualiAgro [\(Table 4\)](#page-31-0). The *I_{ROC}* value systematically underestimated the optimized *K¹* (*ME*=-30%) for QualiAgro. Within each experiment, the variation in *K¹* was well explained 274 by the variability in I_{ROC} ($R^2 = 0.73 - 0.95$). Considering all experiments together, the I_{ROC} 275 explained slightly less of the K_1 variability (R^2 = 0.62).

3.5 Simulations of C storage from EOM with *K¹* **values equal to** *IROC*

 The *IROC* values were used as input *K¹* values in AMG to simulate the ∆SOC stocks for the 278 four experiments for which *I_{ROC}* values were available. *RMSE* logically increased in all

 treatments when replacing optimized *K¹* with *IROC* [\(Figure](#page-32-3) **4**). The mean *RMSE* increased 280 from 3.0 to 4.9 t C ha⁻¹, with a minimal increase of 0.01 t C ha⁻¹ for the FYM_EQH treatment 281 at SERAIL and a maximal increase of 5.4 t C ha⁻¹ for the treatment PEA+N at Ultuna. The largest increases in *RMSE* were associated with experiments with very high ∆SOC stocks 283 (more than 50 t C ha⁻¹ for PEA+N at Ultuna, for example), and thus the *RMSE* increases were relatively low.

3.6 Definition of a database of reference *K¹* **values for AMG**

 Using the *IROC* database (section [2.5\)](#page-10-0), some marked differences among EOM types appeared, [\(Figure 5,](#page-32-4) appendix D in electronic supplementary material). For example, the median *IROC* values of chicken droppings, cattle manure and green waste compost were 16%, 67% and 82%, respectively. Variability was also found within a given type of EOM: for example, the *IROC* values for cattle manure ranged from 21% to 79%. Some EOM types were grouped when an insufficient amount of data existed and if the differences among these EOMs were low. For example, all anaerobic digestates were grouped into one type. Another group was sewage sludge from urban wastewater treatment plants. Finally, we proposed a default *K¹* value for 26 EOM types: 9 types of livestock manure, 1 type of digestate, 2 types of sewage sludge, 13 types of compost and 1 type of other EOM.

4 Discussion

4.1 The accurate simulation by AMG of SOC stock evolution after repeated applications of EOM

 Many authors have observed an increase in SOC stocks following EOM addition, the increase being dependent on the amount and type of EOM applied (Bhogal et al., 2018; Maltas et al., 2018). The importance of the type of EOM in explaining differences in SOC storage was confirmed in experiments where various EOM types applied similar carbon

 inputs (Gerzabek et al., 1997). AMG succeeded in simulating these increases in SOC stocks with various EOM types and under various pedoclimatic conditions and cropping systems [\(Figure](#page-32-1) **2**), for example, in the Ultuna and QualiAgro experiments. The ability of the AMG model to simulate the additional SOC stocks from EOM addition (∆SOC stocks) was 307 satisfactory. The mean *RMSE* value (3.0 t C ha⁻¹) was comparable to the mean standard 308 deviation of the SOC measurements in the different experiments $(2.1 \text{ t C} \text{ ha}^{-1})$ in QualiAgro, 309 for 2.5 t C ha⁻¹ for Colmar). This mean *RMSE* for ∆SOC stocks was also comparable to that obtained by Peltre et al. (2012) with the Roth-C model. These authors found a mean *RMSE* 311 of 3.2 t C ha⁻¹ in the Askov, QualiAgro (1998-2009 only), Ultuna and SERAIL experiments. Even if EOM *K¹* values were optimized to minimize the *RMSE* of ∆SOC stocks, the *RMSE* of the simulated SOC stocks compared to the actual SOC stocks (electronic supplementary material, appendix C) was also similar to that found in other modeling studies. Karhu et al. 315 (2012) reported *RMSE* values of SOC stocks between 2.5 and 10.3 t C ha⁻¹ for different 316 treatments in the Ultuna experiment (from 2.1 to 6.6 t C ha⁻¹ for the same treatments in our study). Begum et al. (2017) obtained an *RRMSE* of 8.9% for the FYM treatment of the Rothamsted experiment (FYM_S1), compared with 13.0% obtained in our study with AMG. These simulations could have been further improved if we had optimized the size of the stable SOC pool. However, this would not have affected the simulation of the ∆SOC stocks.

321 Finally, the mean *RMSE* that we obtained in simulating ∆SOC stocks for all EOM types 322 (3.0 t C ha⁻¹) was close to the *RMSE* obtained for SOC stocks in the control treatments 323 without EOM in our study (2.7 t C ha⁻¹). The mean RMSE from simulating ∆SOC stocks was 324 also close to the mean *RMSE* obtained by Clivot et al. (2019) with the AMG model on a 325 dataset of 60 treatments located at 20 sites in France (2.6 t C ha⁻¹). However, the mean 326 RMSE for SOC stocks with EOM application (4.3 t C ha⁻¹, electronic supplementary material, 327 appendix C) was higher. It is recalled that EOM *K¹* values were not optimized to minimize the 328 *RMSE* for SOC stocks but rather to minimize the RMSE for ∆SOC stocks. Indeed, the poor

 modeling performance for the control treatments at Rothamsted and Askov also impacted the *RMSE* of SOC stocks for treatments with EOM (but not that of ∆SOC stocks). The poor simulation of these control treatments could be related to inappropriate allometric coefficients used for old crop varieties at Rothamsted and to the very low initial C content in the subsoil used at the Askov experiment, for which the SOC mineralization function was not adapted.

4.2 Variability and prediction of the *K¹* **isohumic coefficient for EOM**

 The optimized *K¹* values showed high variability among types of EOM, from 8% to 100%, with a mean value of 63%. Most of the EOM *K¹* values were higher than the *K¹* values of aboveground residues (from 22% to 32%) and roots (40%) (Clivot et al., 2019). This finding confirms that certain EOM types make a greater contribution to SOC than crop residues, as suggested by several authors (Kätterer et al., 2011; Kong et al., 2005), even if the continuation of this greater contribution in the long term is debated (Maillard and Angers, 2014).

 The variability in EOM *K¹* values well reflected the difference in stability of different EOM types, as approximated by their *IROC* value [\(Figure 3\)](#page-32-2). Using the *IROC* indicator as a predictor of *K¹* only slightly decreased model performance [\(Figure](#page-32-3) **4**). The largest increases in *RMSE* were associated with experiments with very high ∆SOC stocks. The small *RMSE* values confirmed the usefulness of *IROC* for parameterizing soil carbon models, as shown by Peltre et al. (2012) with Roth-C. In comparison to the parameters used for this latter study, the use of *IROC* was easier: *IROC* was directly used to derive *K¹* values in AMG without any rescaling, whereas several regressions were needed to predict the Roth-C EOM parameters. Some studies used other laboratory measurements to calibrate EOM in soil carbon models. Mondini et al. (2017) used EOM incubations to predict the partition coefficients and 352 mineralization rates of EOM in Roth-C. Pansu et al. (2017) used $13C$ nuclear magnetic resonance spectra to estimate the parameters of the TAO model. In comparison to these

 studies, the advantages of *IROC* are that it is a standardized (XP U44-162 standard), inexpensive method, and is already being conducted by several commercial laboratories.

4.3 The need for further improvement of the AMG formalism for EOM?

 Even if the performance of AMG in simulating the effects of EOM on SOC stocks was generally satisfying, AMG failed to simulate the substantial increase in SOC stocks observed 359 in some treatments of the QualiAgro experiment $(+24 \text{ t C} \text{ ha}^{-1})$ in 20 years for the BIO N+ 360 treatment), even with the maximal K_1 value (K_1 = 100%). In AMG, organic carbon from EOM is incorporated into the active SOC pool the year following EOM application and is then mineralized at the same rate as the native active SOC pool. The result for QualiAgro suggests that SOC derived from EOM could have a slower mineralization rate than the native SOC. This hypothesis is supported by other modeling studies: tests done with the STICS soil-crop model showed that a slower decomposition constant was needed to accurately simulate the SOC stocks in QualiAgro (Levavasseur et al., in prep). Noirot-Cosson et al. (2016) also calibrated a slower decomposition rate constant for EOM than for SOC to simulate the QualiAgro experiment with the CERES-EGC model. Peltre et al. (2017) observed a higher thermal stability of organic matter in EOM-amended soils. Yu et al. (2012) indicated a lower specific mineralization rate in EOM-amended soils, but Liu et al. (2018) found the opposite result.

372 Although the isohumic coefficient K_1 was generally well predicted by the I_{ROC} indicator, significant residual variability remained. Some very different *K¹* values were optimized for similar EOM types in different experiments. For example, the optimized *K¹* for cattle manure varied greatly among the different experiments. Part of this variability is certainly due to the true variability in OM stability, which could be related to differences in bedding materials and storage duration (Helgason et al., 2005). More surprisingly, we observed an unexplained 378 variability in the $K₁$ of the green waste and sludge compost at QualiAgro and Colmar. Although the compost often came from the same producer, the optimized *K¹* values were

 very different, at 100% and 41%, respectively. This unexplained variability in *K¹* indicates that some factors not considered in the AMG model affected the storage of carbon from EOM. Zhang et al. (2018) showed that accounting for the effects of litter stoichiometry and soil N availability may improve the simulation of SOM formation with the CENTURY model. Peltre et al. (2012) suggested that changes in soil pH may explain some of the larger simulation errors obtained with Roth-C. In our case, the evolution of pH was explicitly considered in the AMG model (annual soil pH was a model input). Brilli et al. (2017) also indicated that poor consideration of soil water conditions may lead to erroneous SOC modeling. The simple annual water balance used in the AMG model to account for the effect of climate on SOC dynamics might be a cause of some poor simulations. For example, in Colmar, the effect of common thunderstorms in summer is not well represented by an annual water balance.

 Finally, even if the AMG formalisms could be improved, AMG performed as well as other models in the literature despite its simpler formalisms. The simple formalisms of the AMG model make it possible for it to be a calibration-free model that is already easy to use outside of academic research, especially thanks to the online tool [\(http://www.simeos-amg.org/\)](http://www.simeos-amg.org/). Therefore, the need to modify the formalisms of AMG to more accurately simulate the effects of EOM should be carefully investigated in other field experiments from both the modeling and the soil chemistry points of view.

5 Conclusion

 Using the AMG model, we simulated the evolution of SOC stocks in seven long-term field experiments with repeated applications of different EOM types. The good model performance confirmed its ability to adequately simulate long-term effects of EOM application on SOC stocks. The only EOM parameter in AMG, the isohumic coefficient *K1,* was optimized for each EOM to minimize the difference in SOC stocks between treatments with and without EOM. 404 These optimized $K₁$ values were well correlated with the organic matter stability of EOM,

 approximated by the indicator of residual organic carbon in soil (*IROC*). This indicator is 406 routinely analyzed by commercial laboratories. The *I_{ROC}* value could be used directly as the EOM *K¹* value in AMG while maintaining satisfactory model performance. A database of *IROC* values was used to parameterize 26 types of EOM in AMG, which can thus be used by stakeholders to model the effects of different EOM practices.

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574 **7 Tables**

576

575 Table 1 Main general characteristics of the experiments used in this study

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578 Table 2 Description of the treatments in the different experiments used in this study

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579

- 581 Table 3 Model performance in simulating ∆SOC stocks in the treatments with EOM in the seven long-
- 582 term field experiments. The optimized value of the isohumic coefficient *K¹* is given for each EOM type

583

- 584 Table 4 Mean error (*ME*), root mean square error (*RMSE*) and coefficient of determination (*R²*)
- 585 between the *IROC* indicator and the optimized *K¹* for all treatments of experiments for which *IROC* values
- 586 were available

587

8 Figure captions

 Figure 1 Conceptual diagram of the AMG model. The organic carbon inputs (*m*) from crop residues, roots or EOM are either mineralized (1-*K¹* fraction) or incorporated (*K¹* fraction) into the active soil organic carbon pool (Active C, QCA). The active organic carbon pool mineralizes according to first-593 order kinetics with the *k* decay constant. The stable organic carbon pool (Stable C, QC_S) is considered completely inert (created with LibreOffice Draw)

 Figure 2 Observed (dots) and simulated (lines) differences in SOC stocks (∆SOC stocks) between the treatments with EOM and without EOM (controls) for the seven long-term field experiments (created with R)

 Figure 3 Relationship between the optimized value of the isohumic coefficient *K¹* and the *IROC* indicator in the long-term field experiments of Colmar (COL), QualiAgro (QUA), SERAIL (SER) and Ultuna 600 (ULT). Error bars around the mean *I_{ROC}* values represent the standard deviations (when available) (created with R)

 Figure 4 Root mean square error (*RMSE*) between simulated and observed differences in SOC stocks in treatments with and without EOM (∆SOC stocks) in four long-term experiments. The isohumic *K¹* 604 coefficient was either optimized (black bars) or equal to I_{ROC} (gray bars) (created with R)

 Figure 5 Distribution of *IROC* values for certain EOM types from the database and median values of *IROC* proposed as reference *K¹* values for AMG (in red). *n* represents the number of *IROC* values for a given EOM. DIG=anaerobic digestate, SS=urban sewage sludge, CM=cattle manure, HM=horse manure, PIS=pig slurry, CD=chicken droppings, GWC=green waste compost, CCM=composted cattle manure (created with R)

h (%)

optimized h \Box h= I_{ROC}

