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






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# The Retreat of Mountain Glaciers since the Little Ice Age: A Spatially Explicit Database

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**Abstract:** Most of the world's mountain glaciers have been retreating for more than a century in response to climate change. Glacier retreat is evident on all continents, and the rate of retreat has accelerated during recent decades. Accurate, spatially explicit information on the position of glacier margins over time is useful for analyzing patterns of glacier retreat and measuring reductions in glacier surface area. This information is also essential for evaluating how mountain ecosystems are evolving due to climate warming and the attendant glacier retreat. Here, we present a non-comprehensive spatially explicit dataset showing multiple positions of glacier fronts since the Little Ice Age (LIA) maxima, including many data from the pre-satellite era. The dataset is based on multiple historical archival records including topographical maps; repeated photographs, paintings, and aerial or satellite images with a supplement of geochronology; and own field data. We provide ESRI shapefiles showing 728 past positions of 94 glacier fronts from all continents, except Antarctica, covering the period between the Little Ice Age maxima and the present. On average, the time series span the past 190 years. From 2 to 46 past positions per glacier are depicted (on average: 7.8).

**Dataset:** 10.6084/m9.figshare.13700215

**Dataset License:** CC-BY-4.0

**Keywords:** glacier retreat; climate change; little ice age; pre-satellite era; global scale

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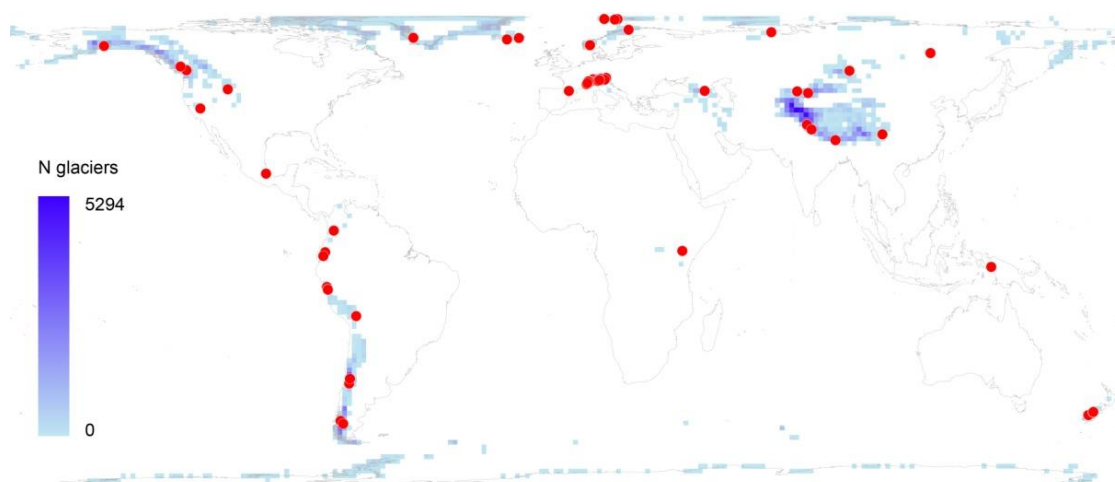
## 1. Summary

Most of the world's mountain glaciers have been losing mass since the second half of 19th century due to the rise of global temperature [1]. Glacier retreat is evident on all continents, and the rate of retreat has accelerated during recent decades [2–4]. In the European Alps, for example, glaciers have lost 25–30% of their surface area over the past 60 years, and the rate of ice loss is accelerating rapidly—it has been 200–300% faster in the past two decades than 40 years ago [5–7], and similar rates of retreat have been measured in other areas of the world [8]. The biotic and abiotic consequences of glacier retreat have received increasing attention in recent years, with research focusing on the

biotic colonization, the formation and evolution of soils along glacier forelands, and the geomorphological hazards related to deglaciation [9–14], as well as on the impacts of glacier retreat on meltwater availability and human wellbeing [15,16]. In this context, broad-scale, spatially explicit information on the dynamics of glacier retreat is essential to assess the ecological dynamics of biotic colonization across multiple regions and to develop adequate adaptation and mitigation strategies to reduce geomorphological risks and cope with meltwater scarcity in arid regions.

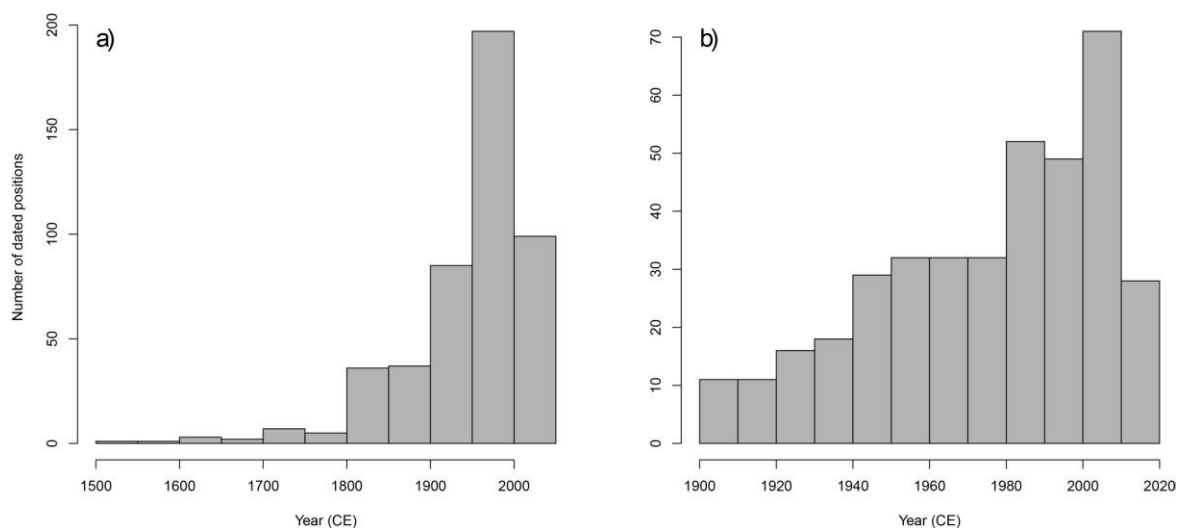
Several databases summarizing information on glacier retreat are currently available (e.g., World Glacier Monitoring Service [17] and Global Land Ice Measurements from Space [18,19]). In most cases, they provide recent outlines obtained through remote sensing. For some glaciers, the GLIMS initiative also provides past outlines, such as glacier extent at the end of the Little Ice Age (LIA). However, these databases generally do not provide information on glacier extent at multiple time points, covering the retreat occurring during the last century. For many glaciers, high-quality data on margins are available since the end of the LIA (from the late 19th century in part of Northern Europe, to as early as the 17th–18th century in other mountain ranges such as the tropical Andes; see e.g., [20]). These data have been obtained through geomorphologic analyses mainly based on morpho-stratigraphic positions, morphology, and relationships of moraines which is further dated by in situ relative and absolute dating methods (e.g., radiocarbon, lichenometry, dendrochronology, optically stimulated luminescence, and terrestrial cosmogenic nuclide dating), analysis of old/repeated photographs and paintings, historical archives and maps including topographical maps, and remotely sensed data [21–24]. The data are typically analyzed using multi-data integrative methods and summarized in long multi-temporal retreat maps. However, because they are derived from disparate sources, the data require manual processing for analysis and presentation. As a consequence, such datasets are mainly available fragmentally for some specific glaciers and for some restricted areas. Thus, there is a need to synthesize such long multi-temporal glacier fluctuation datasets from all over the world to develop spatially explicit datasets showing positions of glacier fronts since the Little Ice Age (LIA) maxima at one place.

We focused on time-series of glacier margins from the LIA maximum extent to the present, with representative examples from the major mountain ranges of the world, except Antarctica. We performed a literature search of glaciers with well-documented retreat series worldwide (i.e., long and spatially explicit time series of glacier margins); the dataset was further complemented with data from several alternative sources (i.e., topographic maps, historical images, and drawings), field work, and remotely sensed data. We focused on mountain glaciers (see [25] for definitions), even though our dataset also included a few glaciers that are linked to icecaps in Iceland and Greenland (Figure 1).



**Figure 1.** Distribution of glaciers included in the dataset (red dots). Due to proximity, some dots are superimposed. The blue shaded areas show the number of extant glaciers for  $1.5^\circ \times 1.5^\circ$  cells (source: [18,19]).

The dataset includes dated margins for 94 glaciers from all the continents except Antarctica (Figure 1). From 2 to 46 past positions are included (average: 7.8 lines per glacier); at least four past positions are shown for 97% of the glaciers. In total, we provide 728 glacier outlines and/or frontal positions for the period from the 16th century to the present. The average length of the time series is 188 years; the length is  $\geq 85$  years for 94% of the glaciers. About 97% of the glacier margins date to the period from the 19th century to today, with a marked increase of data over the second half of the 20th century (Figure 2). The oldest outlines are largely restricted to areas where researchers have dated the LIA maximum back to the 16th–18th centuries (e.g., South America [26,27]).



**Figure 2.** Temporal distribution of glacier margins in the dataset. (a) All margins and (b) 20th and 21st century margins. CE: Common Era.

Although the dataset includes glaciers from all continents (Figure 1), there are differences in coverage among areas, as observed for other environmental datasets [28,29]. Specifically, 35% of the data are from Europe (including Svalbard); 29% are from Asia (including Papua New Guinea); 15% from South America; 11% from Northern and Central America, 8% from Oceania (New Zealand); and 2% are from Africa. Our primary objective was not to obtain a complete, global scale dataset with equal coverage from all the continents, but instead to collate high-quality data with multiple positions from several glaciers around the world. We encourage users to add to our dataset information from additional glaciers.

The present work is part of the European Community's Horizon 2020 project IceCommunities (Grant Agreement no. 772284). IceCommunities combines innovative methods and a global approach to boosting our understanding of the evolution of ecosystems in recently deglaciated areas. IceCommunities investigates chronosequences ranging from recently deglaciated terrains to late successional stages of soil pedogenesis. Through environmental DNA metabarcoding IceCommunities identifies taxa from multiple taxonomic groups (bacteria, fungi, protists, soil invertebrates, and plants), to obtain a complete reconstruction of biotic communities along glacier forelands over multiple mountain areas across the globe and to measure the rate of colonization at an unprecedented level of detail. Information on assemblages is then combined with analyses of soil, landscape, and climate to identify the drivers of community change. IceCommunities also assesses the impact of ecogeographical factors (climate and the regional pool of potential colonizers) on colonization. Analyses of functional traits are also used to reconstruct how functional diversity emerges during community formation, and how it scales to the functioning of food webs. IceCommunities will help to predict the future development of these increasingly

important ecosystems, providing a supported rationale for the appropriate management of these areas.

## 2. Data Description

The dataset is downloadable from <https://doi.org/10.6084/m9.figshare.13700215> (accessed on 4 October 2021). It contains the following data files (Tables 1–4):

**Table 1.** Description of the datasets.

Filename	Description
IC_glac_lines (* .shp, * .shx, * .prj, * .dbf)	ESRI shapefile (EPSG:4326) containing the 728 reconstructed positions of glacier margins for the 94 glacier analyzed. For each line, glacier name, dating, source, GLIMS id, and maximum extent are reported in the associated table.
IC_glac_sites (* .csv)	Table reporting the description of the study sites (glacier name, GLIMS id, country, coordinates of the centroid, mean elevation, mean annual temperature, annual precipitation, lithology, area, and number of reconstructed lines)
IC_glac_references (* .csv)	Table reporting the references cited in IC_glac_lines

### IC\_glac\_lines

**Table 2.** Variable identities, definitions, and attributes for the dataset IC\_glac\_lines.

Identity	Definition	Unit	Storage	Range
glacier	Glacier name	-	Character	-
dating	Dating	year CE	Integer/character	1500 to 2019
source	Reference	-	Character	-
GLIMS_id	Id	-	Character	-
max_extent	Maximum glacier extent <sup>1</sup>	-	Character (y/n)	-

<sup>1</sup> For each glacier, “y” identifies the margin(s) at LIA maximum. The information is missing (i.e., all “n”) when the reconstructed series is incomplete (e.g., Brewster glacier: 1986–2011).

### IC\_glac\_sites

**Table 3.** Variable identity, definition, and attributes for the dataset IC\_glac\_sites.

Identity	Definition	Unit	Storage	Range
glacier	Glacier name	-	Character	-
country	Country name	-	Character	-
GLIMS_id	Id	-	Character	-
lon_wgs84	Longitude (centroid)	DD	Numeric	−149.631 to 170.173
lat_wgs84	Latitude (centroid)	DD	Numeric	−47.530 to 78.899
elev_m	Mean elevation	m a.s.l.	Integer	51 to 5120
mat_°C	Mean annual temperature <sup>1</sup>	°C	Numeric	−13.28 to 9.49
pcp_mm	Annual precipitation <sup>1</sup>	mm	Integer	181 to 4515
litho	Lithology <sup>2</sup>	-	Character	-
area_km2	Glacier area <sup>3</sup>	km <sup>2</sup>	Numeric	0.014 to 8091.670
N_positions	Number of reconstructed lines	-	Numeric	1 to 27

<sup>1</sup> Retrieved from CHELSA [30]. <sup>2</sup> Codes identifying lithology refer to the lithological classes used in Hartmann and Moosdorf [31]. Specifically: mt, metamorphics; ss, siliciclastic sedimentary rocks; pa, acid plutonic rocks; su, unconsolidated sediments; pb, basic plutonic rocks; va, acid volcanic rocks; sc, carbonate sedimentary rocks; vb, basic volcanic rocks; sm, mixed sedimentary rocks; and vi, intermediate volcanic rocks. <sup>3</sup> Most-recent available estimate; retrieved from GLIMS database v20200630 [18,19].

IC\_glac\_references

**Table 4.** Variable identity, definition, and attributes for the dataset IC\_glac\_references.

Identity	Definition	Unit	Storage	Range
References	Sources used to reconstruct or date glacier margins	-	Character	-

### 3. Methods

We focused on time-series of glacier margins from the LIA maximum extent to the present, with representative examples from the major mountain ranges of the world, except Antarctica. We first performed a literature search of glaciers for which there are long and spatially explicit time series of glacier margins. Data from the literature were complemented with new data, obtained mostly from topographic maps; historical, aerial, and satellite images; and field surveys. Some of the glacier margins and dates are based on our measurements made in the field. Older positions are based mainly on moraines that are clearly visible on images and in the field, and have been dated using lichenometry, dendrochronology, radiocarbon, and cosmogenic nuclides. The reconstruction of LIA maxima and subsequent glacier extent have been carried out differently by different studies, in most of cases using a multi-data layer integration approach (MDIA, [32]). This approach incorporates individual layers of information extracted from geomorphological mapping, analysis of photo sequences, historical archives, maps inferences, and hillshade DEM analysis into a GIS environment. For many glaciers, glacial geomorphological evidence and landforms (e.g., lateral, recessional, and hummocky moraines, supraglacial morainic ridges, trim lines, and palaeo-channels) resultant due to LIA glaciation and latterly molded by deglaciation are initially mapped using high-resolution remote sensing images and DEMs and further validated in the field. These data are integrated with sequences of pictures taken in the field in different times, or obtained from satellite/aerial images. Moreover, additional information on the historical terminus, surface characteristics, and the extents of individual glaciers was extracted from historical descriptions, documents, and maps preserved since LIA maxima, and existing marks in the field. All the spatial data were integrated into a spatial database, and the output was further validated against known LIA positions from available regional chronologies (e.g., [26]).

We used four approaches to validate the dated margins for each glacier: (i) we performed a double-check against the original publication; (ii) each shapefile was checked by more than two co-authors, to confirm the consistence across areas of the world; (iii) the database was reviewed by regional experts, i.e., by researchers experienced in the geomorphology and mapping of glaciated areas of a study region; and (iv) we then performed a final check based on available high resolution satellite images in Google Earth.

Images were georeferenced and lines were digitized using QGIS 3.4.12; additional analyses were performed using R 4.0.5.

### 4. User Notes

The final dataset is provided in ESRI shapefile format (WGS 84, decimal degrees—EPSG:4326). Missing/anomalous data are present in both IC\_glac\_lines and IC\_glac\_sites. They refer to some GLIMS IDs lacking (glacier not in the references database or extinct). Additionally, it was not always possible to obtain precise datings for the glacier margins, particularly those older than the first half of the 20th century (marked as “NA”, “LIA”, “M2”, or “M3” and “(estimated- . . .)”). Sources of uncertainty included the following:

(1) For a number of glaciers, dating of old margins were based on published geomorphological chronologies of the region, rather than on the glacier itself. For example, LIA moraines in the Peruvian Andes, although clearly visible in the field, have not been directly dated for all the glaciers, therefore we assume ages similar to those of nearby glaciers [25]. Similarly, in the absence of direct dating, we assume that LIA moraines of glacier margins

in the European Alps date to the last half of the 19th century, even though some variation might exist among glaciers due to their different response time [33]. Cases with large age uncertainties are explicitly acknowledged in the dataset.

(2) Even if a moraine has been directly dated (e.g., using lichenometry or radiocarbon dating), the user must be aware that every technique has inherent uncertainties. The user should refer to the reference(s) cited in the dataset for further information on this uncertainty.

(3) Finally, some level of spatial uncertainty exists, for instance when data are based on old maps or images, mostly because of their limited quality and/or spatial resolution.

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