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1 **Topsoil clay content mapping in croplands from Sentinel-2 data: influence** 2 **of atmospheric correction methods across a season time series**

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15

16 **Abstract**

17 Recent studies demonstrated the capability of Sentinel-2 (S2) data to estimate topsoil
18 properties and highlighted the sensitivity of these estimations to soil surface conditions
19 depending on the S2 acquisition date. These estimations are based on Bottom Of Atmosphere
20 (BOA) reflectance images, obtained from Top Of Atmosphere (TOA) reflectance values using
21 Atmospheric Correction (AC) methods. AC of optical satellite imagery is an important pre-
22 processing stage before estimating biophysical variables, and several AC methods are
23 currently operational to perform such conversion. This study aims at evaluating the sensitivity
24 of topsoil clay content estimation to atmospheric corrections along an S2 time series. Three
25 AC methods were tested (MAJA, Sen2Cor, and LaSRC) on a time series of eleven Sentinel-2

26 images acquired over a cultivated region in India (Karnataka State) from February 2017 to
27 June 2017. Multiple Linear Regression models were built using clay content analyzed from
28 topsoil samples collected over bare soil pixels and corresponding BOA reflectance data. The
29 influence of AC methods was also analysed depending on bare soil pixels selections based on
30 two spectral indices and several thresholds: the normalized difference vegetation index
31 (NDVI below 0.25, 0.3 and 0.35) and the combination of NDVI (below 0.3) and Normalized
32 Burned Ratio 2 index (NBR2 below 0.09, 0.12 and 0.15) for masking green vegetation, crop
33 residues and soil moisture.

34 First, this work highlighted that regression models were more sensitive to acquisition
35 date than to AC method, suggesting that soil surface conditions were more influent on clay
36 content estimation models than variability among atmospheric corrections. Secondly, no AC
37 method outperformed other methods for clay content estimation, and the performances of
38 regression models varied mostly depending on the bare soil pixels selection used to calibrate
39 the regression models. Finally, differences in BOA reflectance among AC methods for the
40 same acquisition date led to differences in NDVI and NBR2, and hence in bare soil coverage
41 identification and subsequent topsoil clay content mapping coverage. Thus, selecting S2
42 images with respect to the acquisition date appears to be a more critical step than selecting an
43 AC method, to ensure optimal retrieval accuracy when mapping topsoil properties assumed to
44 be relatively stable over time.

45

46 **Keywords:** clay content; Sentinel-2; atmospheric correction; multiple linear regression; soil
47 property mapping; India.

48

49 **1. Introduction**

50 Soils are key to meeting global environmental sustainability challenges for food security,
51 water security, energy sustainability, climate stability, biodiversity, and ecosystem service
52 delivery (McBratney et al., 2014). An accurate and spatially referenced characterization of
53 soil properties over cultivated areas, including soil organic matter, soil texture, or iron
54 content, is essential for meeting these global environmental sustainability challenges and
55 would also allow to help for planning agricultural engineering work such as land
56 consolidation, drainage management, and soil erosion prevention. Particle-size distribution,
57 also called soil texture, refers to relative amounts of sand, silt, and clay in grams per kilogram
58 (g kg^{-1}) in a soil, the sum of these particle size fractions equaling 1000 g kg^{-1} . Soil texture is a
59 major component of soil, as it has an important influence on water infiltration and soil
60 stabilization (Le Bissonnais et al., 2007, 2018). To ensure soil security, adequate decisions
61 both at global and local levels are required to favor the beneficial roles of soil (Rodrigo-
62 Comino et al., 2020). Such decisions require accurate spatially referenced soil information
63 systems that can be used in environmental modeling.

64 Spectroscopy covering the Visible, Near Infrared, and Short Wave Infrared domains
65 (VNIR/SWIR, 400 – 2500 nm) is a technology that proved its relevance for the estimation of
66 soil properties (Viscarra Rossel et al., 2006) as the soil reflectance spectrum results from the
67 position and shape of absorption features of chemical constituents (“peaks”) (e.g., water
68 molecules influence the absorption features at specific wavelengths that are the results of
69 overtone and combination modes from the IR region) and overall spectral shape of the
70 physical properties (e.g., texture) (Ben-Dor and Banin, 1995a, 1995b). Most of the
71 absorptions in the VNIR/SWIR region are characteristic of clay and organic matter, and are
72 dominated by C–H, O–H, N–H and metal–OH bonds (e.g., Clark et al., 1990). Soil physical
73 properties, including particle size, scatter the light in a way that the spectrum shape and base
74 line are changing (Wetzel, 1983; Chabrilat et al., 2019). According to Ben-Dor et al. (2002),

75 an accurate estimate of a soil property can be expected from VNIR/SWIR data if the targeted
76 soil property i) is related to a chemical species that impacts soil surface reflectance values
77 through absorption bands (e.g., OH⁻ ion for clay) or ii) is highly correlated with the latter
78 (e.g., CEC correlated with clay content). In addition to these two rules, a minimum level of
79 variability of the targeted soil property across study regions is required to be accurately
80 estimated (Gomez et al., 2012a).

81 VNIR/SWIR airborne spectroscopy was successfully used to map a large range of soil
82 properties such as iron, soil organic carbon (SOC), and clay contents over bare soil surfaces
83 with high accuracy (e.g., Ben-Dor et al., 2002; Stevens et al., 2010; Gomez et al., 2008,
84 2012b; Chabrillat et al., 2019). More recently, the VNIR/SWIR multispectral satellites
85 Sentinel-2 (S2) enabled mapping topsoil properties over bare soil surfaces, such as SOC (e.g.,
86 Gholizadeh et al., 2018; Vaudour et al., 2019a, b, 2021; Žížala et al., 2019; Castaldi et al.,
87 2021; Dvorakova et al., 2021; Urbina-Salazar et al., 2021) and texture (e.g., Gholizadeh et al.,
88 2018; Gomez et al., 2019; Bousbih et al., 2019). Despite lower accuracy of S2 estimates
89 compared to airborne imaging spectroscopy, the global coverage and high revisit frequency of
90 S2 (five days at the equator) opened perspectives for regional mapping. However, the
91 selection of an S2 acquisition over a multi-temporal series in order to have an optimal
92 estimation of soil properties, such as clay content, raises questions due to multiple factors
93 influencing surface reflectance, and possibly statistical models adjusted to estimate soil
94 properties (Vaudour et al., 2019b; Gomez et al., 2019; Castaldi et al., 2019). Indeed soil
95 surface conditions such as soil moisture and roughness (e.g., due to tillage operations)
96 influence soil reflectance (Chabrillat et al., 2019). The Normalized Difference Vegetation
97 Index (NDVI, based on Red and NIR spectral bands) and the Normalized Burned Ratio 2
98 (NBR2) index (based on spectral bands around 1600 nm and 2200 nm) have been widely used
99 as indicators for photosynthetic vegetation and dry crop residues, respectively (e.g., for

100 Landsat 8 by [Demattê et al., 2018](#) and S2 by [Castaldi et al. 2019](#)). Nevertheless, while NBR2
101 follows a linear relationship with crop residue cover over dry soils, no correlation with residue
102 cover could be found over moist soils ([Dvorakova et al., 2020](#)). So it is assumed that NBR2
103 reacts both to crop residues and soil moisture, where high values of NBR2 indicate soils that
104 are moist and/or are covered by crop residues. In the way of creating a composite multi-date
105 bare soil image based on S2 images for soil organic carbon (SOC) prediction, [Dvorakova et](#)
106 [al., \(2021\)](#) and [Vaudour et al. \(2021\)](#) developed strategies that allow selecting S2 pixels with
107 minimal influence of crop residues, surface roughness or soil moisture, using spectral indices,
108 especially NDVI and NBR2 indices for detecting these disturbing factors.

109 Atmospheric conditions also strongly influence the Top of Atmosphere (TOA)
110 reflectance (Level-1C products), and the choice of atmospheric correction (AC) method used
111 to convert S2 Level-1C into Level-2A products may result in differences in terms of Bottom
112 of Atmosphere (BOA) reflectance. These AC methods use different atmospheric models and
113 hypotheses, which may affect the resulting BOA reflectance depending on the season, clouds,
114 sun azimuth, and elevation, and therefore the soil property estimations. Several AC methods
115 have been developed for multispectral images (e.g., Multi-sensor Atmospheric Correction and
116 Cloud Screening -MACCS, [Hagolle et al. \(2015a\)](#), updated and renamed « MACCS-ATCOR
117 joint algorithm » or MAJA, [Lonjou et al. \(2016\)](#); Sen2Cor, [Gascon et al. \(2017\)](#), [Louis et al.](#)
118 [\(2016\)](#); Land Surface Reflectance Code -LaSRC, [Vermote et al. \(2016\)](#)). MAJA and Sen2Cor
119 methods were both developed specifically for Sentinel-2 and have become widely used
120 standard products available from Copernicus and other platforms. LaSRC, originally
121 developed by NASA for Landsat, has been adapted recently for Sentinel-2, so it may also
122 become a standard product in the future (e.g., considered in the Brazil Data Cube, [Ferreira et](#)
123 [al., 2020](#)). These methods use quite different processing paths, from the atmospheric
124 parameter estimation to the radiative transfer model and temporal information, justifying

125 interest in their comparison. A first Atmospheric Correction Inter-comparison Exercise
126 (ACIX) was carried out under an international collaborative initiative to compare a set of AC
127 methods for optical sensors, including S2 (Doxani et al., 2018). However, because the
128 exercise continued, Doxani et al. (2018) did not draw common conclusions from all the
129 algorithms. In addition, Sola et al. (2018a, 2018b) evaluated four AC methods for S2 images,
130 highlighting minor differences between these AC methods. Finally, to the best of our
131 knowledge, no work has been conducted on the impact of AC methods on soil property
132 estimation.

133 Following, on the one hand, the studies from Doxani et al. (2018) and Sola et al.
134 (2018a, 2018b) evaluating AC methods for S2 images, and those from Gomez et al. (2019)
135 and Vaudour et al. (2019b) studying the impact of acquisition dates on prediction
136 performances of texture and SOC contents, the objective of this study is to evaluate the
137 impact of three major AC methods (Sen2Cor, MAJA, and LaSRC) along an S2 time series on
138 clay content estimation in the topsoil. The influence of these AC methods was also analysed
139 depending on bare soil pixels selections based on two spectral indices and several thresholds:
140 the NDVI (below 0.25, 0.3 and 0.35) and the combination of NDVI (below 0.3) and NBR2
141 (below 0.09, 0.12 and 0.15). Eleven S2 images covering a cropping cycle in 2017 in a
142 cultivated region in India (Karnataka State) were selected for this study. Soil samples were
143 collected over the study area and analyzed in a laboratory. Clay content was then estimated
144 from each BOA S2 image, corrected by the three AC methods, using multiple linear
145 regression models.

146

147 **2. Materials and Methods**

148 **2.1. Study area**

149 The Berambadi catchment is a subcatchment of the South Gundal located in the Deccan
150 Plateau of Southern India (Figure 1a), extending over 84 km² (Figure 1b). Our study area is
151 located in the eastern part of the Berambadi catchment, of which 60% of the surface is
152 dedicated to agriculture, with a high diversity of crops (e.g., sunflower, marigold, sorghum,
153 turmeric, maize, etc.) and high seasonal variability. The western part of the Berambadi
154 catchment is covered by forest (not shown in Figure 1b). The Berambadi catchment belongs
155 to the Kabini Critical Zone Observatory (AMBHAS, BVET, Tomer et al., 2015; Sekhar et al.,
156 2016), which is part of the OZCAR network (Gaillardet et al., 2018). The climate is tropical
157 subhumid with an average rainfall of 800 mm/year and a PET of 1100 mm (aridity index
158 P/PET of 0.7). The monsoon dynamics drive three main seasons: dry season (winter in
159 January and February, summer from March to May), Kharif (southwest monsoon season,
160 from June to September), and Rabi (north-east monsoon season, from October to December).
161 Red soils (Ferralsols and Chromic Luvisols) cover the uplands and hillslopes, while black
162 soils (Vertisols and Vertic intergrades) are mostly found in the valley bottom (Barbiero et al.,
163 2010). Uplands and hillslopes are mainly characterized by coarse soil texture (sandy loam)
164 due to erosion processes, whereas valleys bottoms are mainly characterized by finer soil
165 texture (clay) mainly caused by deposition processes (Gunnell and Bourgeon, 1997; Barbiero
166 et al., 2010).

167 Three cropping seasons regulate the farm system (Robert et al., 2017). During the
168 Kharif season (from June to September) corresponding to the rainy season, most of the
169 cropping area are cultivated with sorghum, maize, sunflower, marigold, as well as crops
170 grown in irrigated conditions, such as turmeric, onion, garlic, and banana. During the Rabi
171 season (from October to January) corresponding to the winter season, most irrigated plots are
172 cultivated with maize, horse gram, and vegetables. Finally, during the Summer season (from

173 February to May) corresponding to the hot and dry season, only few plots are cultivated, and
174 almost 90% of the cropland is bare land.

175 [Figure 1]

176

177 **2.2. Soil dataset**

178 A total of 164 topsoil samples were collected over Berambadi in November 2019 (Figure 1b).

179 All samples were composed of five sub-samples collected to a depth of 5 cm within a 10 m
180 \times 10 m square (one at the center and four at each corner) centered on the geographical position
181 of the sampling plot, as recorded by a Garmin GPS instrument. After sample homogenization,
182 approximately 20 g was devoted to soil property analysis. The samples were then air-dried
183 and sieved with a 2 mm sieve prior to being analyzed in the laboratory. The clay fraction was
184 determined using the pipette method as described by Piper (1966). The clay content ranged
185 between 58 and 622 g kg⁻¹ (mean = 252 g kg⁻¹, standard deviation = 124 g kg⁻¹ and skewness
186 = 5 g kg⁻¹).

187

188 **2.3. Sentinel-2 images**

189 Launched in 2015 and then 2017 the combination of both Sentinel-2 satellites (S2A and S2B)
190 delivers a revisit period of five days at the equator. The multispectral sensors acquired
191 information over thirteen spectral bands in the VNIR/SWIR spectral domain, with spatial
192 resolution ranging from 10 m to 60 m.

193 Twenty-six images from the S2 tile 43PFP were acquired over Berambadi between
194 February 1st and June 30th 2017. This corresponds to the summer season and beginning of
195 Kharif season in South India, when both a maximum of bare soil pixels and a minimum of
196 clouds can be encountered. As clay content is a perennial property in topsoil horizons, we can
197 assume that the clay content analyzed from soil samples collected in 2019 remained

198 unchanged compared to 2017, as assumed in previous studies (e.g., [Gomez et al., 2008](#);
199 [Loiseau et al., 2019](#)).

200

201 **2.3.1. Atmospheric correction method**

202 Atmospheric corrections were performed on the 26 images using three AC methods described
203 in the next subsections (MAJA, [Lonjou et al., 2016](#); Sen2Cor, [Louis et al., 2016](#); LaSRC ;
204 [Vermote et al., 2016](#)) providing three time series of atmospherically corrected images ([Figure](#)
205 [2a](#)). After atmospheric correction, only ten bands were retained for spectral modeling,
206 corresponding to the four 10 m resolution bands (B2, B3, B4 and B8, see Table 1 in
207 [Dvorakova et al., 2021](#)) and the six 20 m resolution bands (B5, B6, B7, B8A, B11 and B12,
208 see Table 1 in [Dvorakova et al., 2021](#)). The BOA products were provided with reflectance
209 *10000.

210 [Figure 2]

211

212 **a. Sen2Cor**

213 Since 2015, the Copernicus Open Access Hub has provided Level-2A products of S2 imagery
214 data over Europe, using the Sen2Cor processor developed by European Space Agency (ESA)
215 ([ESA, 2015](#); [Gascon et al., 2017](#); [Louis et al., 2016](#)). The Sen2Cor processor performs
216 atmospheric correction, terrain and cirrus correction, and scene classification applied to TOA
217 data.

218 The Sen2Cor 2.8 was applied to produce Level-2A images from Level-1C images using
219 default settings (available in the Ground Image Processing Parameters files delivered with
220 Sen2Cor official release, [https://step.esa.int/main/snap-supported-](https://step.esa.int/main/snap-supported-plugins/sen2cor/sen2cor_v2-8)
221 [plugins/sen2cor/sen2cor_v2-8](https://step.esa.int/main/snap-supported-plugins/sen2cor/sen2cor_v2-8)). The Planet Digital Elevation Model was used, and cirrus and
222 Bidirectional Reflectance Distribution Function (BRDF) corrections were deactivated. All

223 other parameters were set to their default value. In addition to Level-2A reflectance data,
224 Sen2Cor also produces an Aerosol Optical Thickness (AOT) map, a Water Vapour (WV)
225 map, and a Scene Classification Map (SCM) together with Quality Indicators (QI) for cloud
226 and snow probabilities. After conversion to Level-2A reflectance, Sen2Cor provides the four
227 VNIR spectral bands with their native spatial resolution of 10 m and the six 20 m resolution
228 bands, which were resampled to 10 m using nearest-neighbor interpolation.

229

230 **b. MAJA**

231 The MAJA processor (Lonjou et al., 2016; Hagolle et al., 2019) was initially developed to
232 perform cloud detection and atmospheric correction over time series of optical images
233 acquired at high resolution and under quasi constant viewing angles. MAJA combines Multi-
234 Mission Atmospheric Correction and Cloud Screening (MACCS) developed by the French
235 Centre National d'Études Spatiales (CNES) and ATCOR developed by the German
236 Aerospace Center (DLR). This spectro-temporal AC method was developed to process images
237 from Formosat-2, Landsat, VENμS, and S2 satellites. MAJA is based on a spectral
238 assumption of the link between red and blue spectral bands and a temporal assumption
239 assuming that a given neighborhood separated by a few days should yield similar surface
240 reflectance (Hagolle et al., 2015a; 2015b). In the current study, the MAJA correction was
241 processed with on-demand PEPS (Plateforme d'Exploitation des Produits Sentinel) processing
242 service, which uses eight S2 acquisitions prior to each acquisition of interest to meet the
243 temporal assumption ([https://labo.obs-mip.fr/multitemp/on-demand-sentinel2-l2a-processing-
244 with-maja-on-peps](https://labo.obs-mip.fr/multitemp/on-demand-sentinel2-l2a-processing-with-maja-on-peps), last access on 2021-07-01). As Sen2Cor, MAJA provides spectral bands
245 with their native spatial resolution. After Level-2A conversion, we used nearest-neighbor
246 interpolation to convert all 20 m bands to 10 m resolution.

247

248 c. **LaSRC**

249 Landsat Surface Reflectance Code (LaSRC) is an AC method initially developed to convert
250 TOA radiance to BOA reflectance for Landsat 8 collection, which was recently adapted to
251 Sentinel-2 (Vermote et al., 2018). The algorithm performs atmospheric correction, assuming a
252 Lambertian-plane-parallel atmosphere, and using the Second Simulation of the Satellite
253 Signal in the Solar Spectrum (6S) “Urban Clean” model (Vermote et al., 2016; USGS, 2016).
254 The atmospheric parameters required for the inversion include surface pressure (from the
255 National Center for Environmental Prediction Global Data Assimilation System—NCEP
256 GDAS weather model), water vapor (derived from the MODIS near-infrared channels), ozone
257 (from NCEP GDAS), and aerosol properties (AOT and Angstrom exponent) (Claverie et al.,
258 2018). The aerosol properties are estimated using the comparison between assumed surface
259 reflectance ratios computed from MODIS time series and Sentinel-2 TOA reflectance ratios.
260 The correction algorithm is applied at a spatial resolution of 10 m, which is the output
261 resolution for all bands. Therefore, the L1C bands with a resolution of 20 and 60 m are
262 resampled with a nearest-neighbor method in a pre-processing step.

263

264 **2.3.2. Cloud mask**

265 Pixels identified as clouds or cloud shadows were masked over the S2 time series, using the
266 cloud mask product provided by MAJA (Figure 2a). The MAJA cloud mask is computed with
267 a combination of mono-temporal and multi-temporal approaches (Hagolle et al., 2010) and
268 accounts for different types of clouds (low, high, thin) and corresponding projected shadows
269 (Baetens et al., 2019). Recent cloud mask algorithms inter-comparisons highlighted that the
270 MAJA cloud mask algorithm provided similar to better performances over other cloud and
271 shadow mask algorithms on a large variety of environments (Baetens et al., 2019 ; Tarrío et
272 al., 2020).

273

274 **2.3.3. Bare soil selection**

275 S2 pixels identified as non-cultivated land were masked using a land-use map available for
276 the study area (AMBHAS Team, 2015) (Figure 2a). These pixels correspond to urban areas,
277 bodies of water, and natural vegetation (forest). Among the 26 images available over the
278 study area, eleven were finally kept for this study, for which the surface of cultivated land
279 outside clouds and cloud shadows covered more than 95% of the Berambadi catchment.

280 After masking pixels corresponding to non-cultivated land, bare soil were
281 discriminated from photosynthetic vegetation, based on a thresholding applied on the
282 normalized difference vegetation index (NDVI) calculated using the red band B04 (665 nm)
283 and the NIR band B08 (842 nm) (Figure 2a). Three NDVI thresholds common to all images
284 were defined based on visual interpretation from an expert with field knowledge and literature
285 (e.g., Stevens et al., 2008; Vaudour et al., 2016; Lu et al., 2013), : (i) below 0.25, (ii) below
286 0.3 and (iii) below 0.35. Pixels corresponding to bare soil were also differentiated from crop
287 residue and moist soil, applying thresholding on the Normalized Burned Ratio 2 index
288 (NBR2) calculated using the SWIR1 band B11 (1610 nm) and the SWIR2 band B12 (2202
289 nm) (Figure 2a). As setting a threshold for the NBR2 index might be difficult without relevant
290 field observation (Dvorakova et al., 2020 and 2021), four thresholds of NBR2 were tested: (i)
291 no threshold, (ii) below 0.15, (iii) below 0.12 and (iv) below 0.09. . This study focused on the
292 common pixels identified as bare soil for all images to compare clay predictions obtained for
293 the different acquisitions and AC methods.

294 Topsoil water content is directly related to the time since last precipitations. As soil
295 moisture affects spectra (Diek et al., 2016; Vaudour et al., 2019b) and in order to compare the
296 NBR2 response with rainfall events, we calculated the number of days from the last rain for
297 each S2 image to hint at the topsoil humidity conditions, based on daily rainfall measured in

298 the Maddur village, on the West part of the Berambadi catchment. Among these eleven
 299 selected images, two images were acquired following heavy rainfall (cumulative rainfall of
 300 25.5 mm and 18.4 mm over the last two days before the S2 data acquisition on 08-03-2017
 301 and 04-04-2017, respectively), three images were acquired following moderate rainfall
 302 (cumulative rainfall of 6.5 mm, 5.5 mm and 9 mm over the last five days before the S2 data
 303 acquisition on 25-03-2017, 24-04-2017, and 07-05-2017), and six images were acquired after
 304 more than five days without rain (03-02-2017, 16-02-2017, 23-02-2017, 26-02-2017, 28-03-
 305 2017, and 27-04-2017).

306

307 **2.4 Spectral Measures**

308 The BOA reflectances produced with the three different AC methods were compared pairwise
 309 to estimate the spectral similarity between AC methods over the eleven dates (Figure 2b).
 310 These spectral similarity analysis were performed using R software (R Development Core
 311 Team, 2015). The coefficient of correlation $r_{i,t}$ was calculated between the BOA reflectance
 312 values as follows:

$$313 \quad r_{i,t} = cor(Refl_{i,t}^{AC1}, Refl_{i,t}^{AC2}) \quad (1)$$

314 where $Refl_{i,j}^{AC1}$ and $Refl_{i,j}^{AC2}$ are the BOA reflectance values obtained from AC methods 1
 315 and 2, at band i for the S2 image acquired at date t . The coefficient of correlation r was
 316 calculated over the N bare soil pixels, for which soil sample was collected and clay content
 317 was analyzed.

318 As the coefficient of correlation $r_{i,t}$ does not reflect the bias in the data (due to change
 319 in albedo in our case), the bias $Rbias_{i,t}$ between two AC methods at band i for the S2 image
 320 acquired at date t was calculated as follows:

$$321 \quad Rbias_{i,t} = \sum_{k=1}^N \frac{Refl_{i,t,k}^{AC1}}{N} - \sum_{k=1}^N \frac{Refl_{i,t,k}^{AC2}}{N} \quad (2)$$

322 where $Refl_{i,t,k}^{AC1}$ and $Refl_{i,t,k}^{AC2}$ are the BOA reflectance values for pixel k , at band i for the S2
 323 image acquired at date t , obtained from one AC and another, respectively, and N is the
 324 number of considered pixels. The bias $Rbias_{i,t}$ was calculated over the N bare soil pixels for
 325 which soil sample was collected and clay content was analyzed.

326 The spectral angle was used to analyze the spectral similarity between BOA
 327 reflectance spectra. The spectral angle $SA_{k,t}$ considers the whole spectra and not a single band
 328 as $r_{i,t}$ and $Rbias_{i,t}$ (Kruse et al., 1993). It was calculated between the BOA spectrum obtained
 329 from one AC and another, for pixel k and the S2 image acquired at date t , as follows:

$$330 \quad SA_{k,t} = \frac{\sum_{i=1}^{nb} Refl_{i,t,k}^{AC1} * Refl_{i,t,k}^{AC2}}{\sqrt{\sum_{i=1}^{nb} Refl_{i,t,k}^{AC1^2}} \sqrt{\sum_{i=1}^{nb} Refl_{i,t,k}^{AC2^2}}} \quad (3)$$

331 where $Refl_{i,t,k}^{AC1}$ and $Refl_{i,t,k}^{AC2}$ are the BOA reflectance values at band i obtained from one AC
 332 method and another, respectively, for pixel k , and nb is the total number of spectral bands (10
 333 in our case). $SA_{k,t}$ ranges between 0 and 1, with low values corresponding to low spectral
 334 similarity and high values corresponding to high spectral similarity. The spectral angle
 335 mapper SA_t was finally calculated for the S2 image acquired at date t , over the N bare soil
 336 pixels for which soil sample was collected and clay content was analyzed, as follows:

$$337 \quad SA_t = \frac{1}{N} \sum_{k=1}^N SA_{k,t} \quad (4)$$

338 where $SA_{k,t}$ was calculated from Equation (3).

339

340 2.5. Regression model

341 Regression models and analysis were performed using R software (R Development Core
 342 Team, 2015), and both the ade4 (Dray and Dufour, 2007) and pls packages (Mevik and
 343 Wehrens, 2007) were used.

344 A Multiple Linear Regression (MLR) method was used to produce clay maps
 345 estimated from S2 images. MLR is a multivariate approach adjusting a linear relationship

346 between a dependent (response) variable (Y-variable, i.e., clay content in the present case),
347 and a set of predictor variables (X-variables, i.e., S2 spectra in the present case) (Tenehaus,
348 1998) (Figure 2d). A restrictive selection of pixels corresponding to bare soil may result in
349 small sample size available to train and validate regression models using independent
350 datasets. In this work, a k-fold cross-validation (CV) was used. The original dataset was
351 randomly divided into k sub-datasets. Then, k-1 sub-datasets were used as training data, and
352 the remaining one was used as validation data. The CV process was repeated k times, and the
353 model performance was evaluated by averaging prediction error obtained for the k sub-
354 datasets. The k-fold CV method can take full advantage of data, as each part of the original
355 dataset is randomly divided and used for both training and testing. Here, 10-fold cross-
356 validation (CV) was used to build robust methods for estimating the accuracy of MLR models
357 and repeated 5 times.

358 Finally, three statistical criteria were used to assess model performances: mean
359 absolute error (MAE), root mean square error ($RMSE_{cv}$), and Pearson correlation coefficient
360 (R^2_{cv}) of cross validation (Figure 2e). These statistical criteria were calculated taking into
361 account the 10-fold CV and 5 repetitions.

362

363 **3. Results**

364 **3.1 Bare soil coverage analysis**

365 The distribution of NDVI values calculated over all unmasked pixels showed positive skew
366 for all acquisition dates and AC methods with values lower than 0.25 (Figure 3). Larger
367 differences in NDVI distributions between the three AC methods were obtained for the S2
368 image acquired on 08-03-2017 (Figure 3). The median value of NDVI distributions obtained
369 using the LaSRC method (around 0.21) was slightly higher than the median obtained using
370 Sen2Cor and MAJA methods (around 0.19) regardless of the date (Figure 3).

371 [Figure 3]

372

373 The distribution of NBR2 values calculated over all unmasked pixels showed positive skew
374 for all acquisition dates and AC methods with values lower than 0.2 (Figure 4). The median
375 value of NBR2 distributions obtained using the LaSRC (around 0.12 along the dates) method
376 was slightly higher than the median obtained using Sen2Cor and MAJA methods (around 0.9
377 and 0.10 along the dates, respectively), regardless of the date (Figure 4). The AC methods
378 affected the NBR2 values (Figure 4) more than NDVI values (Figure 3) as NBR2
379 distributions differed from an AC method to another, especially for the S2 images acquired on
380 08-03-2017, 04-04-2017, 24-04-2017, 27-04-2017 and 07-05-2017.

381 [Figure 4]

382

383 Bare soil coverage selected using both NDVI and NBR2 thresholding varied depending on
384 both acquisition dates and AC methods (Table 1). The acquisition date providing the
385 maximum bare soil coverage varied among AC methods. As an example, using a NDVI
386 below 0.3, the S2 image providing the maximum bare soil coverage was the one acquired on
387 26-02-2017 from MAJA-corrected S2 images (with 85.9%), on 08-03-2017 from Sen2Cor-
388 corrected S2 images (with 85.9%), and on 23-02-2017 from LaSRC-corrected S2 images
389 (with 82.7%) (Table 1). The acquisition date providing the minimum bare soil coverage also
390 varied from an AC method to another. Still using a NDVI below 0.3, the S2 image providing
391 the minimum bare soil coverage was the one acquired on 07-05-2017 along MAJA-corrected
392 S2 images (with 76.3%), on 03-02-2017 along Sen2Cor-corrected S2 images (with 73.4%)
393 and on 08-03-2017 along LaSRC-corrected S2 images (with 75.1%) (Table 1). Regardless of
394 the NDVI and NBR2 thresholds, the image acquired on 08-03-2017 presented maximum bare
395 soil coverage for all AC methods (range from 5.7 to 20.75%, Table 1). Regardless of the

396 NDVI and NBR2 thresholds, the Sen2Cor method provided the highest range of bare soil
397 coverage along the acquisition dates (range from 9.18 to 18.49%, Table 1). Finally, a majority
398 of LaSRC-corrected S2 images provided slightly less bare soil coverage than those corrected
399 by Sen2Cor and MAJA, regardless of the NDVI threshold or the combination of NDVI and
400 NBR2 thresholds (Table 1).

401 [Table 1]

402
403 The bare soil pixels that were common for all S2 images corrected by MAJA, Sen2Cor and
404 LaSRC covered from 67.3% of the surface based on a NDVI below 0.35 to 1.2% based on
405 NDVI and NBR2 below 0.3 and 0.09, respectively (Table 2). This resulted in varying sample
406 sizes of bare soil locations with clay content information, which ranged from 122 samples
407 based on NDVI below 0.35 to 2 samples based on NDVI and NBR2 below 0.3 and 0.09,
408 respectively (Table 2). Due to the poor bare soil coverage on 08-03-2017 using NBR2 below
409 to 0.12 or lower, especially from the LaSRC method (Table 1, Figure 4), the bare soil
410 coverage common to all dates did not exceed 7%, which allowed 12 or less samples (Table 2).

411 As the combinations of NDVI below 0.3 and NBR2 below 0.09 or 0.12, were too
412 restrictive, this study used the four datasets reaching more than 45 collected samples
413 identified as bare soil pixels (i.e. extracted using NDVI below 0.25, 0.3, 0.35 and using the
414 combination of NDVI and NBR2 below 0.3 and 0.15, respectively) to train MLR models
415 (Figures 2c and d) and apply them to the corresponding bare soil coverage. These four dataset
416 displayed close distributions with a range between 58 and 592 g kg⁻¹, a mean around 220 g kg⁻¹,
417 and skewness from 7 to 8.7 g kg⁻¹ (Table 2).

418 [Table 2]

419

420 3.2 BOA reflectance analysis based on bare soil pixels identified with NDVI below 0.3

421 The effect of AC methods on BOA reflectance values was investigated based on the 111 bare
422 soil pixels for which a measured clay content value was available for training the regression
423 models, identified with a NDVI below 0.3 (Table 2). The BOA reflectance values obtained for
424 each image corrected by the three AC methods were compared based on these 111 bare soil
425 pixels. Reflectance obtained with LaSRC showed lower values for almost all bands, except
426 for B04 on 08-03-2017 (Figure 5C) and for B03 (Figure 5B). Reflectance obtained with
427 Sen2Cor from bands B04 to B12 (665 nm to 2190 nm, red boxplots, Figures 5 from C to J)
428 showed systematically higher values than reflectance obtained with MAJA and LaSRC
429 (yellow and blue boxplots, respectively, Figures 5 from C to J), except on 08-03-2017 for
430 Band 04. The largest difference in BOA reflectance obtained among the three AC methods at
431 the same date was observed on 08-03-2017 and mainly over the visible spectral bands
432 (Figures 5A, B, and C). The SA_t values (Equation (4)), computed for pairwise comparison
433 among AC methods over the 111 topsoil spectra, are high, up to 0.994, showing high
434 similarities between spectra corrected by different AC methods (data not shown).

435 [Figure 5]

436

437 The Pearson's correlation coefficient $r_{i,t}$ (Equation (1)) computed between the 111 topsoil
438 spectra corrected by MAJA and Sen2Cor were very high for all bands and dates, ranging from
439 0.93 to 0.99 (Figure 6A1). The lowest correlations were obtained for B02 (490 nm) and B03
440 (560 nm), with a mean of 0.97, while the NIR and SWIR bands showed a very high
441 correlation (> 0.99) (Figure 6A1). Along the eleven dates, the lowest correlations were
442 obtained for the S2 image acquired on 08-03-2017 (mean of 0.99 over the spectral bands),
443 while the highest correlations were obtained for the S2 image acquired on 16-02-2017 (mean
444 up to 0.999 over the spectral bands) (Figure 6A1).

445 The correlations $r_{i,t}$ between the 111 topsoil spectra corrected by MAJA and LaSRC
446 followed similar patterns to correlations $r_{i,t}$ calculated between the 111 spectra corrected by
447 Sen2Cor and LaSRC (Figures 6A2 and A3). The correlations $r_{i,t}$ were very high varying from
448 0.90 to 0.99, except for B02 (from 0.52 and 0.90) and B03 (from 0.79 and 0.97). Along the
449 eleven dates, the lowest correlations were obtained for the S2 image acquired on 08-03-2017
450 (mean of 0.88 over the spectral bands, Figures 6A2 and A3).

451
452 The $Rbias_{i,t}$ (in absolute value) between the 111 spectra corrected by MAJA and Sen2Cor
453 ranged from 2.5 (B02 at 490 nm, on 23-02-2017) to 400.0 (B12 at 2190 nm, on 24-04-2017)
454 for all the bands and dates (Figure 6B1). Along the 10 spectral bands, the highest $Rbias_{i,t}$
455 were obtained for the B12 (2190 nm) (mean of 302.2 in absolute value) while the lowest
456 $Rbias_{i,t}$ were obtained for the B02 (490 nm) (mean of 68 in absolute value) (Figure 6B1).
457 Along the eleven dates, the highest $Rbias_{i,t}$ were obtained for the S2 image acquired on 04-
458 04-2017 (mean of 220.5 in absolute value), while the lowest $Rbias_{i,t}$ were obtained for the S2
459 image acquired on 16-02-2017 (mean of 76.3 in absolute value) (Figure 6B1).

460 The $Rbias_{i,t}$ (in absolute value) between the 111 spectra corrected by MAJA and
461 LaSRC varied from 0.00 (B8A at 865 nm, on 28-03-2017) to 249.6 (B02 at 490 nm, on 08-
462 03-2017) for all bands and dates (Figure 6B2). Along the 10 spectral bands, the highest
463 $Rbias_{i,t}$ were obtained for the B02 (490 nm) (mean of 112.3 in absolute value) while the
464 lowest $Rbias_{i,t}$ were obtained for the B08A (865 nm) (mean of 10.9 in absolute value) (Figure
465 6B2). Along the eleven dates, the highest $Rbias_{i,t}$ were obtained for the S2 image acquired on
466 08-03-2017 (mean of 135.9 in absolute value), while the lowest $Rbias_{i,t}$ were obtained for the
467 S2 image acquired on 25-03-2017 (mean of 33.2 in absolute value) (Figure 6B2).

468 The $Rbias_{i,t}$ (in absolute value) between the 111 spectra corrected by Sen2Cor and
469 LaSRC varied from 20 (B02 at 490 nm, on 03-02-2017) and 460.8 (B05 at 705 nm, on 08-03-

470 2017) for all the bands and dates (Figure 6B3). Along the 10 spectral bands, the highest
471 $R_{bias_{i,t}}$ were obtained for the B05 (705 nm) (mean of 310.9 in absolute value), while the
472 lowest biases were obtained for the B08A (865 nm) (mean of 128.8 in absolute value) (Figure
473 6B3). Along the eleven dates, the highest $R_{bias_{i,t}}$ were obtained for the S2 image acquired on
474 08-03-2017 (mean of 308.4 in absolute value), while the lowest $R_{bias_{i,t}}$ were obtained for the
475 S2 image acquired on 16-02-2017 (mean of 100.8 in absolute value) (Figure 6B3).

476 [Figure 6]

477

478 3.3 Clay predictions based on bare soil pixels identified with NDVI below 0.3

479 The effect of AC methods on clay content estimations was firstly investigated based on bare
480 soil pixels obtained with NDVI below 0.3. MLR models were built from each S2 image and
481 each AC method using the 111 topsoil samples identified with NDVI below 0.3, providing 33
482 MLR models. Performances for the prediction of clay content strongly varied depending on
483 the acquisition date, with R^2_{cv} ranging from 0.49 to 0.72 for MAJA, from 0.45 to 0.71 for
484 Sen2Cor, and from 0.43 to 0.71 for LaSRC (Table 3).

485 The difference in R^2_{cv} obtained when comparing dates of acquisition for each AC
486 method varied from 0.23 (obtained from MAJA-corrected S2 images) to 0.29 (obtained from
487 LaSRC-corrected S2 images) (Table 3), suggesting a strong impact of the acquisition date on
488 clay content estimation. Based on MAJA- and Sen2Cor-corrected S2 images, the best MLR
489 model was obtained from the S2 image acquired on 24-04-2017 ($R^2_{cv} > 0.71$ and $RMSE_{CV} >$
490 6.50% , Table 3), while based on LaSRC-corrected S2 images, the best MLR model was
491 obtained from the S2 image acquired on 25-03-2017 (R^2_{cv} of 0.71 and $RMSE_{CV}$ of 6.38%,
492 Table 3). The MLR models built from the S2 image acquired on 04-04-2017 provided poor
493 performances for clay content prediction ($R^2_{cv} < 0.55$ and $RMSE_{CV} > 6.50\%$, Table 3) for all
494 AC methods.

495 The difference in R^2_{cv} obtained when comparing the three corrected S2 images for a
496 given date of acquisition varied from 0.00 (obtained on 27-04-2017) to 0.07 (obtained on 08-
497 03-2017) (Table 3), suggesting a poor impact of the AC method selection on clay content
498 estimation. The three AC methods provided very close regression performances for six S2
499 images (03-02-2017, 26-02-2017, 25-03-2017, 28-03-2017, 24-04-2017 and 27-04-2017,
500 Table 3). Considering the five remaining S2 images, MAJA provided corrected S2 images
501 associated to the best clay content estimations for four acquisition dates, while LaSRC
502 provided corrected S2 images associated to the best clay content estimations for only one
503 acquisition dates (04-04-2017, Table 3).

504 Finally, the largest difference in R^2_{cv} obtained when comparing AC methods for a
505 given date of acquisition was 0.07, while the largest difference in R^2_{cv} obtained when
506 comparing dates of acquisition for each AC method was 0.29 (Table 3). So it suggested a
507 stronger impact of the acquisition date on clay prediction performance compared to the AC
508 method.

509 [Table 3]

510

511 **3.4 Influence of bare soil pixels identification on clay predictions**

512 The analysis of AC influence on clay content estimations was extended to varying bare soil
513 pixels identifications: with NDVI below 0.25 and 0.35, and with the combination of NDVI
514 and NBR2 below 0.3 and 0.15, respectively. Based on these bare soil pixels identifications,
515 MLR models were built from MAJA-, Sen2Cor- and LaSRC-corrected images.

516 Regardless of the AC methods and thresholdings, the poorest performances were
517 obtained from the S2 image acquired on 08-03-2017 (Table 3). Using MAJA-corrected S2
518 images, the best performances were obtained from the S2 image acquired on 24-04-2017,
519 independently from thresholding. Using Sen2Cor- and LaSRC-corrected S2 images, the best

520 performances were obtained from S2 images acquired on 25-03-2017 and 24-04-2017,
521 depending on thresholding (Table 3).

522 Regression models performances obtained based on bare soil coverage selected using
523 NDVI below 0.3 and 0.35, were similar. Using such NDVI thresholds, the largest difference
524 in R^2_{cv} obtained when comparing AC methods for a given date of acquisition was 0.07, while
525 the largest difference in R^2_{cv} obtained when comparing dates of acquisition for each AC
526 method was 0.29 (Table 3). Using the 84 topsoil samples identified with NDVI below 0.25,
527 the largest difference in R^2_{cv} obtained when comparing AC methods for a given date of
528 acquisition was 0.04, while the largest difference in R^2_{cv} obtained when comparing dates of
529 acquisition for each AC method was 0.31 (Table 3). Using the 47 topsoil samples identified
530 with NDVI below 0.3 and NBR2 below 0.15, the largest difference in R^2_{cv} obtained when
531 comparing AC methods for a given date of acquisition was 0.08, while the largest difference
532 in R^2_{cv} obtained when comparing dates of acquisition for each AC method was 0.26 (Table 3).
533 These results suggest that the acquisition date has stronger influence than the AC method on
534 clay prediction performance. Finally, applying restrictive NDVI thresholds or combining
535 NDVI and NBR2 thresholds did not systematically improve the models performance for a
536 given acquisition date (Table 3).

537

538 **4. Discussion**

539 *Variation in bare soils coverage*

540 The extent of bare soil coverage varied depending on acquisition date, AC method and index
541 thresholds (Table 1), with more variability due to dates and index thresholds than AC
542 methods. While the non-cultivated lands mask (urban areas, bodies of water, and natural
543 vegetation forest) was unique regardless of AC methods and date of acquisition, the cloud,
544 and land masks were consistent only among the three AC methods but not among the eleven

545 dates, and the photosynthetic vegetation masks varied between AC methods and the date of
546 acquisition. Indeed, the clouds were masked using one cloud mask specific to each S2 date
547 provided by MAJA (Table 1, Section 2.3.2), and the photosynthetic vegetation was masked by
548 applying a threshold on the NDVI calculated using the red band B04 (665 nm) and the NIR
549 band B08 (842 nm) of each S2 image.

550 As clouds only affected three dates (08-03-2017, 04-04-2017, and 27-04-2017, Table
551 1) and bands B04 and B08 differed slightly depending on the S2 acquisition date (Figures 5C
552 and G, respectively), the variability of bare soil coverage along the dates was mainly due to
553 the cut-off induced by NDVI and NBR2 values, which may vary according to some changes
554 in soil surface conditions (due to humidity, roughness), new crop growing within the dates
555 (Vaudour et al., 2019b) or crop residues (Dvorakova et al., 2021). The variability of NDVI
556 values for the same date depending on AC methods was also previously observed by Sola et
557 al. (2018a, 2018b).

558 The choice of an AC method showed very minor influence on bare soil coverage,
559 except for the image acquired on 08-03-2017 (e.g., from 58.22% to 78.97% using NDVI <
560 0.25 and LaSRC- and Sen2Cor-corrected S2 images, Table 1). For this specific date, such
561 difference may be due to some unmasked clouds (including cirrus) over the Berambadi (Table
562 1) affecting the reflectance along the spectra, hence the NDVI values. Tarrío et al. (2020)
563 showed that although MAJA detects clouds and cloud shadows fairly well, producing few
564 errors of omission, the majority of omitted clouds for MAJA was composed of high cirrus
565 clouds.

566

567 *Variation in BOA spectra (based on bare soil pixels identified with NDVI below 0.3)*

568 Atmospheric corrections performed with Sen2Cor resulted in higher reflectance values for
569 bands B04 to B12 (665 nm to 2190 nm, red boxplots, Figures 5 from C to J), except for the

570 date 08-03-2017 in Band 04 (Figure 5C), compared to others AC methods. These results are
571 in agreement with those obtained by Sola et al. (2018a, 2018b), who reported higher BOA
572 reflectance obtained by Sen2Cor compared to MAJA and the 6S model. Li et al. (2018) also
573 observed an overestimation of surface reflectance by Sen2Cor, especially for bright pixels,
574 and suggested that may be due to an AOT overestimation. Reflectance obtained with LaSRC
575 showed systematically lower values for almost all bands, with the exception for B04 on 08-
576 03-2017 (Figure 5C) and generally for B03 (Figure 5B), than reflectance obtained with MAJA
577 and Sen2Cor. These results are still in agreement with Sola et al. (2018a), who reported lower
578 BOA reflectance obtained by the 6S model compared to MAJA and Sen2Cor. Moreover,
579 reflectance values obtained on 08-03-2017 from bands B05 to B12 (705 nm to 2190 nm,
580 Figure 4) were the lowest of the time series, which may be explained by a higher soil moisture
581 content on the 08-03-2017 image due to previous rainfall, likely to cause a general decrease in
582 reflectance (Minasny et al., 2011) or some unmasked clouds (including cirrus).

583 Variations in performances between AC methods might be explained by several
584 factors: i) MAJA is based on a multitemporal approach, while other methods are based on
585 individual acquisitions and ii) the computation of AOT and water vapor as well as the
586 radiative transfer model differs among methods, iii) each AC method uses its own cloud mask
587 method to screen the main cloudy pixels for which AC parameter estimations (such as AOT)
588 would probably not converge. Regarding this last factor, a gap-filling operation was
589 performed to fill AOT values for cloudy pixels (either by using a constant value for MAJA or
590 interpolation for Sen2Cor and LaSRC) before applying atmospheric correction with radiative
591 transfer models. Such differences in cloudy pixels may result in local differences in correction
592 accuracy, especially in the surroundings of shadowed or clouded pixels, and therefore
593 contribute to a small part of the difference between the accuracy of MAJA and the other
594 methods. However, considering the low rate of omission of MAJA cloud detection exposed in

595 [Baetens et al. \(2019\)](#), as well as the spatial homogeneity of the atmospheric correction
596 parameters, this cannot fully explain the differences in accuracy observed between the
597 methods.

598 Despite these differences observed for BOA reflectance produced by the three AC
599 methods ([Figures 5 and 6](#)), the spectra were highly correlated regardless of the date of
600 acquisition, except for 08-03-2017, and regardless of the bands, except the B02 (490 nm).
601 Such high correlations are in accordance with the results obtained by [Padró et al. \(2017\)](#), who
602 compared BOA reflectances obtained by several AC methods, including Sen2Cor, with in situ
603 reflectances, and obtained high correlations regardless of the AC methods ($r^2 > 0.9$).

604

605 *AC method as minor driver of prediction performances*

606 Our MLR regression models for clay content estimations provided performances ([Table 3](#)) in
607 accordance with the performances obtained by [Shabou et al. \(2015\)](#) and [Gasmi et al. \(2021\)](#)
608 using multispectral Landsat images over Mediterranean contexts, and by [Bellinaso et al.](#)
609 [\(2021\)](#) and [Vaudour et al. \(2019a\)](#) using S2 images over tropical and temperate/Mediterranean
610 contexts, respectively.

611 Regardless of the bare soil coverage selection, the AC method had a limited influence
612 on MLR performances for clay content prediction ([Table 3](#)). Indeed the MLR models built
613 from MAJA-, Sen2Cor- and LaSRC-corrected S2 images provided close performances ([Table](#)
614 [3](#)). This is in accordance with [Marujo et al. \(2021\)](#), who expected that LaSRC and Sen2Cor
615 would converge to produce consistent and comparable data from both sensors. However, this
616 is not in accordance with [Sola et al. \(2018a\)](#), who showed MAJA provided better
617 performances in the analysis per land cover compared to Sen2Cor.

618

619 *S2 acquisition date as the main driver of prediction performances*

620 Regardless of the bare soil coverage selection, the date of S2 image acquisition showed a
621 stronger influence than the AC method on the performances for clay content prediction (Table
622 3). As previously shown by Demattê et al. (2018) and then Castaldi et al. (2019), Vaudour et
623 al. (2019b) and Dvorakova et al. (2021), soil surface conditions impact the accuracy of SOC
624 prediction models. This seems the case through our observed time-series, for which the dates
625 acquired shortly after a rainfall event resulted in the lowest performance (on 08-03-2017 and
626 04-04-2017, Table 3), presumably due to higher soil moisture.

627 As the date of S2 image acquisition appears to be very important for soil property
628 estimation, and this even more compared to the AC method, future researches could focus on
629 identifying rules for selecting the best date of image acquisition. This was started by selecting
630 dates with average driest condition or lower crop residues over bare soils based on NDVI and
631 NBR2 spectral indices and enabled to obtain the best prediction performances using single-
632 dates images (Castaldi et al., 2019; Vaudour et al., 2019b; Urbina-Salazar et al., 2021).

633 Similarly, rules for date selection may be used to create multi-temporal image
634 composites composed of multirate images stacked over the same tile. From now, multi-
635 temporal image composites are built from either pixelwise thresholding based on the
636 minimum pixel value (Loiseau et al., 2019; Vaudour et al., 2021), average reflectance value
637 (e.g., Demattê et al., 2018; Gasmi et al., 2019; Dvorakova et al., 2021), median reflectance
638 value (Castaldi et al., 2021; Luo et al., 2022), or considering a trade-off between average per
639 date-indices and maximum bare soil coverage (Vaudour et al., 2021) along a multirate
640 satellite series. Multiple studies suggested using multi-temporal image composites to
641 maximize bare soil coverage for soil property prediction (e.g., Demattê et al., 2018; Vaudour
642 et al., 2021; Dvorakova et al., 2021). However, current researches do not converge towards a
643 common conclusion about their benefits compared to single-date images, especially in terms
644 of soil property prediction accuracies. While Gasmi et al. (2019) showed that multi-temporal

645 image composites based on mean spectral reflectance from bare soil pixels along a Landsat-
646 TM time series allowed to increase both the prediction accuracy of soil clay content and
647 mapping coverage, [Vaudour et al. \(2021\)](#) showed that none of the multi-temporal image
648 composites based on pixelwise or per-date bare soil reflectance along a S2 time series
649 improved model performance for SOC prediction compared to the best single-date image.
650 Therefore, benefit and methodologies for multi-temporal mosaicking are critical questions
651 remaining to be explored.

652

653 **5. Conclusions**

654 The influence of three atmospheric corrections, namely MAJA, Sen2Cor and LaSRC, on
655 eleven S2 images was evaluated based on clay content estimation over bare soil pixels. Our
656 study highlighted the influence of the S2 acquisition date and AC method on model
657 performances adjusted for clay content estimations, with more variability induced by the
658 acquisition date than the AC method. Thus, the influence of the choice of an AC on the
659 estimation of soil properties can be considered as moderate compared to soil surface
660 conditions, such as moisture, crop residue or roughness, which may be strongly variable in
661 space and time. As regression models performances were close from one AC method to
662 another, this work did not allow to consider one AC method to be the best method prior
663 estimating clay content. As Sen2Cor provided performances for clay content estimations close
664 to MAJA and LaSRC methods, and since ESA provides corrected imagery with Sen2Cor, this
665 AC method might be a satisfactory choice. Finally, as soil properties such as organic carbon
666 and iron are key properties influencing soil radiometric properties in a different manner than
667 clay, the impact of AC methods on the estimation of such topsoil properties could be further
668 investigated to test the robustness of our conclusions.

669

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679

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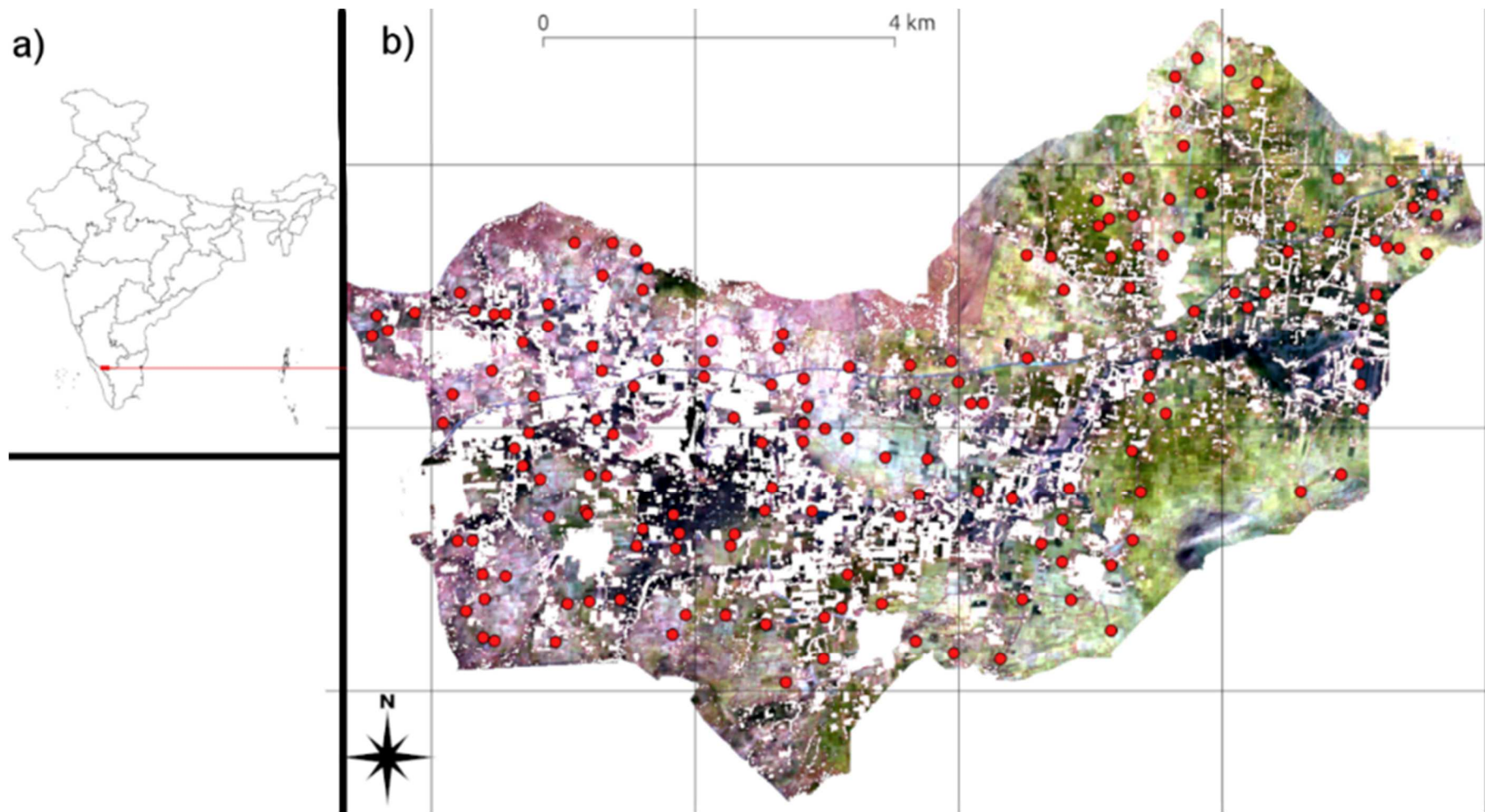


Figure 1: a) Location of the Berambadi catchment in India (red dot) and b) the 164 collected soil surface samples (red dots) plotted over the S2 image acquired on 03-02-2017 (colored composite of bands 08 (R), 04 (G), and 03 (B)). White pixels correspond to masked pixels (non-bare soils).

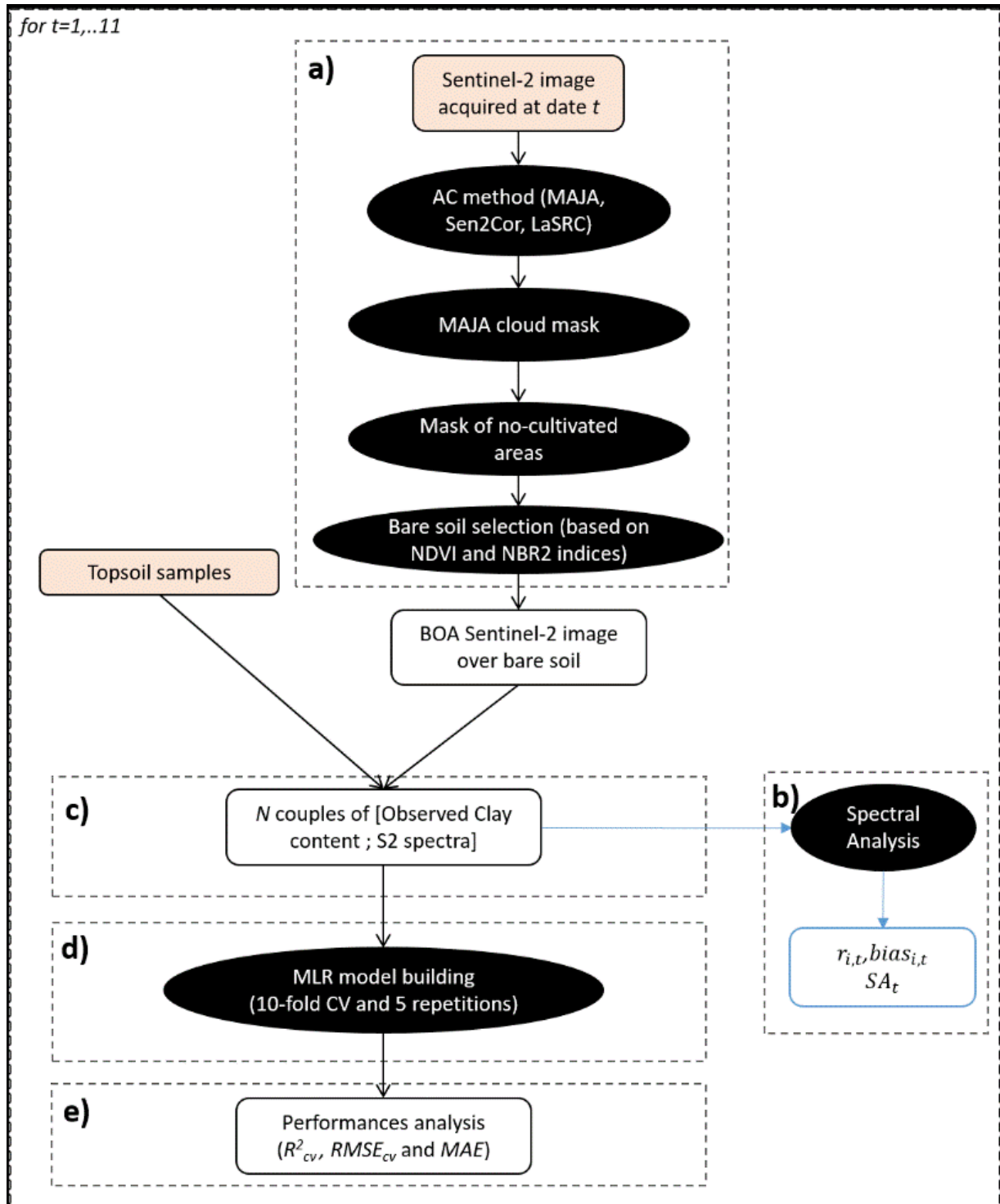


Figure 2: Workflow. a) Preparation of the BOA S2 images restricted to bare soil pixels, b) spectral analysis of BOA images across the time-series and obtained from the 3 AC methods, c) preparation of both calibration and validation datasets for d) model building and e) validation of clay predictions.

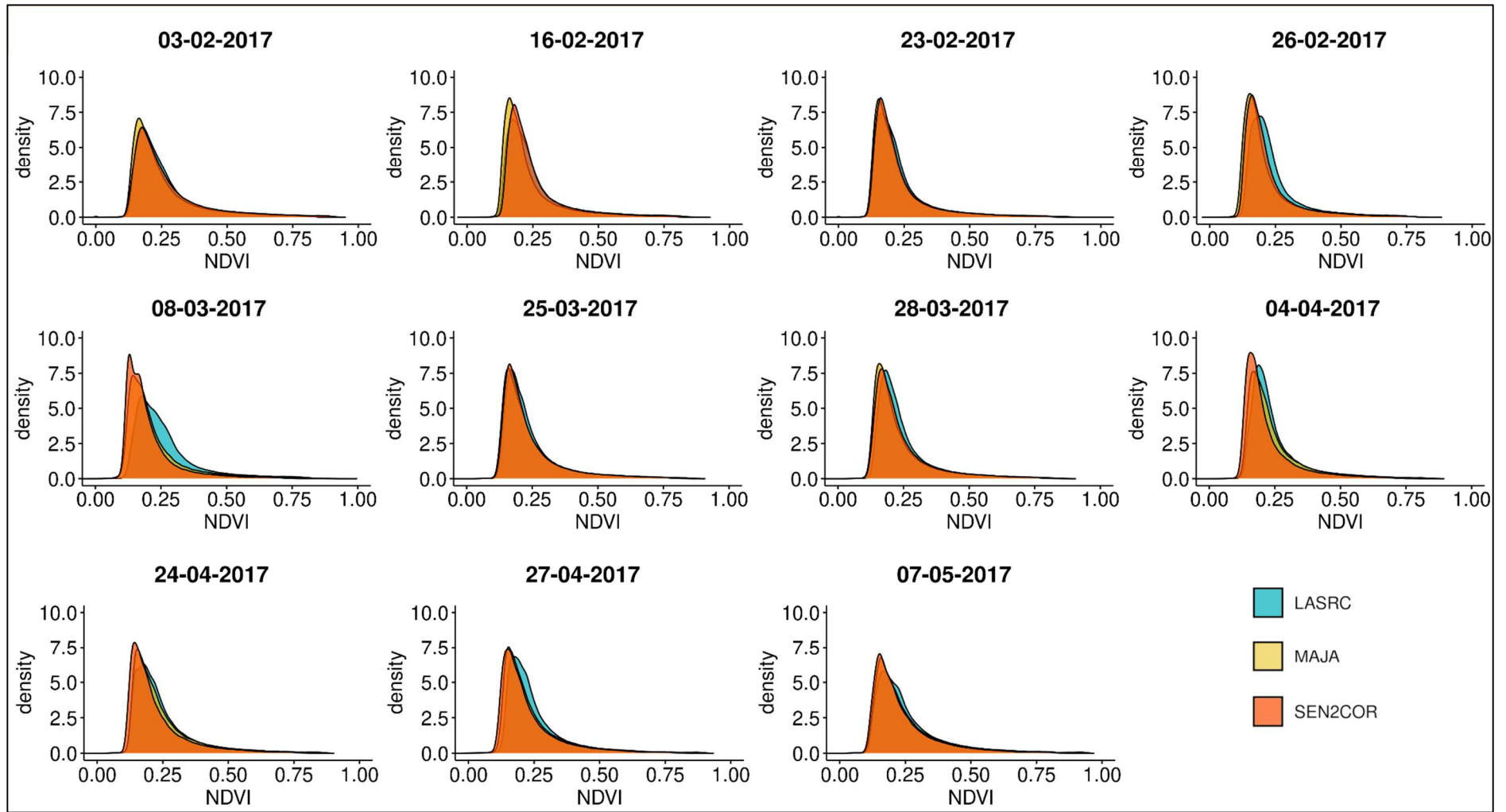


Figure 3: Histograms of NDVI values calculated for each S2 image corrected by the MAJA (yellow), Sen2Cor (orange), and LaSRC (cyan) AC methods.

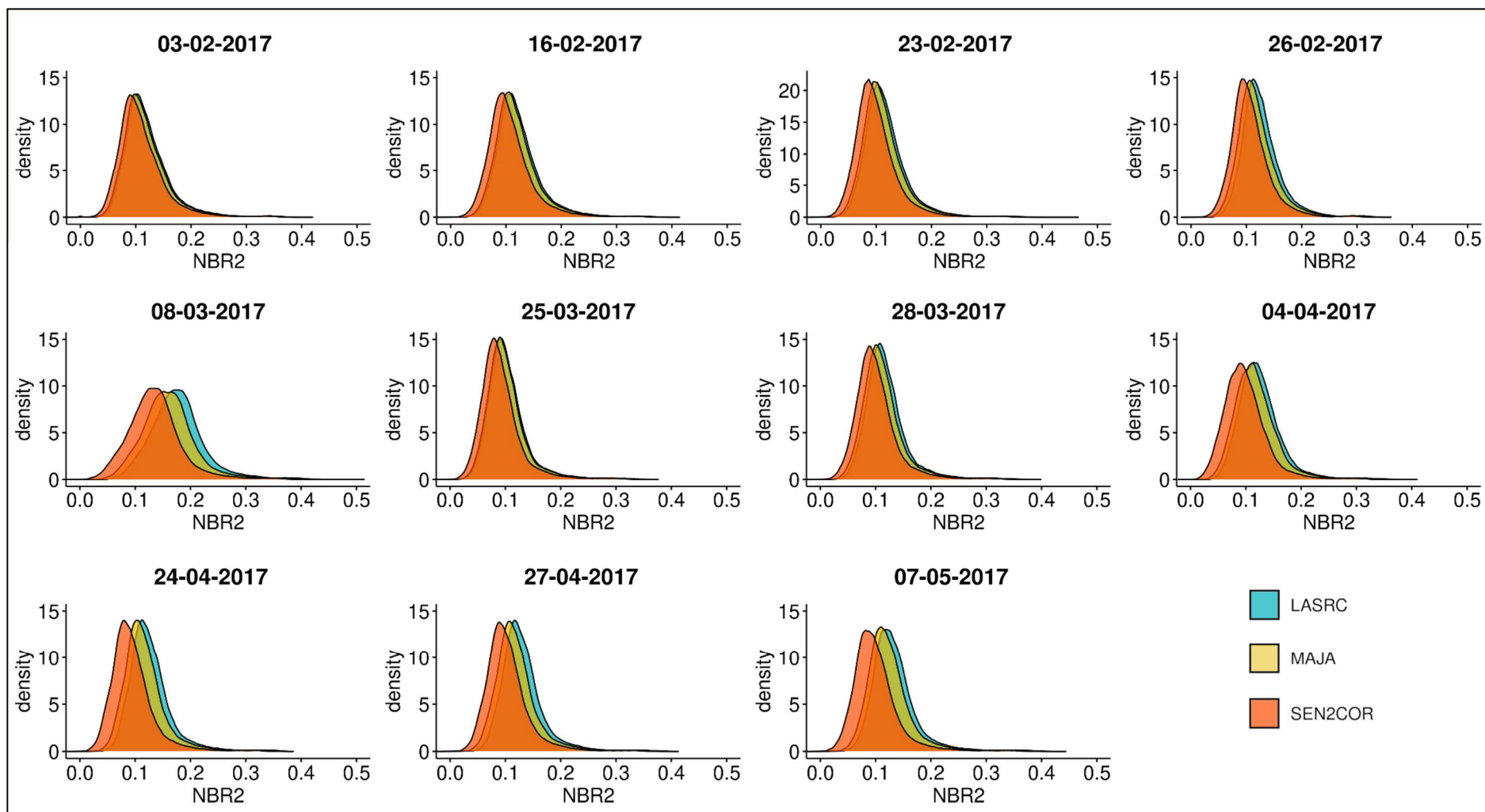


Figure 4: Histograms of NBR2 values calculated for each S2 image corrected by the MAJA (yellow), Sen2Cor (orange), and LaSRC (cyan) AC methods.

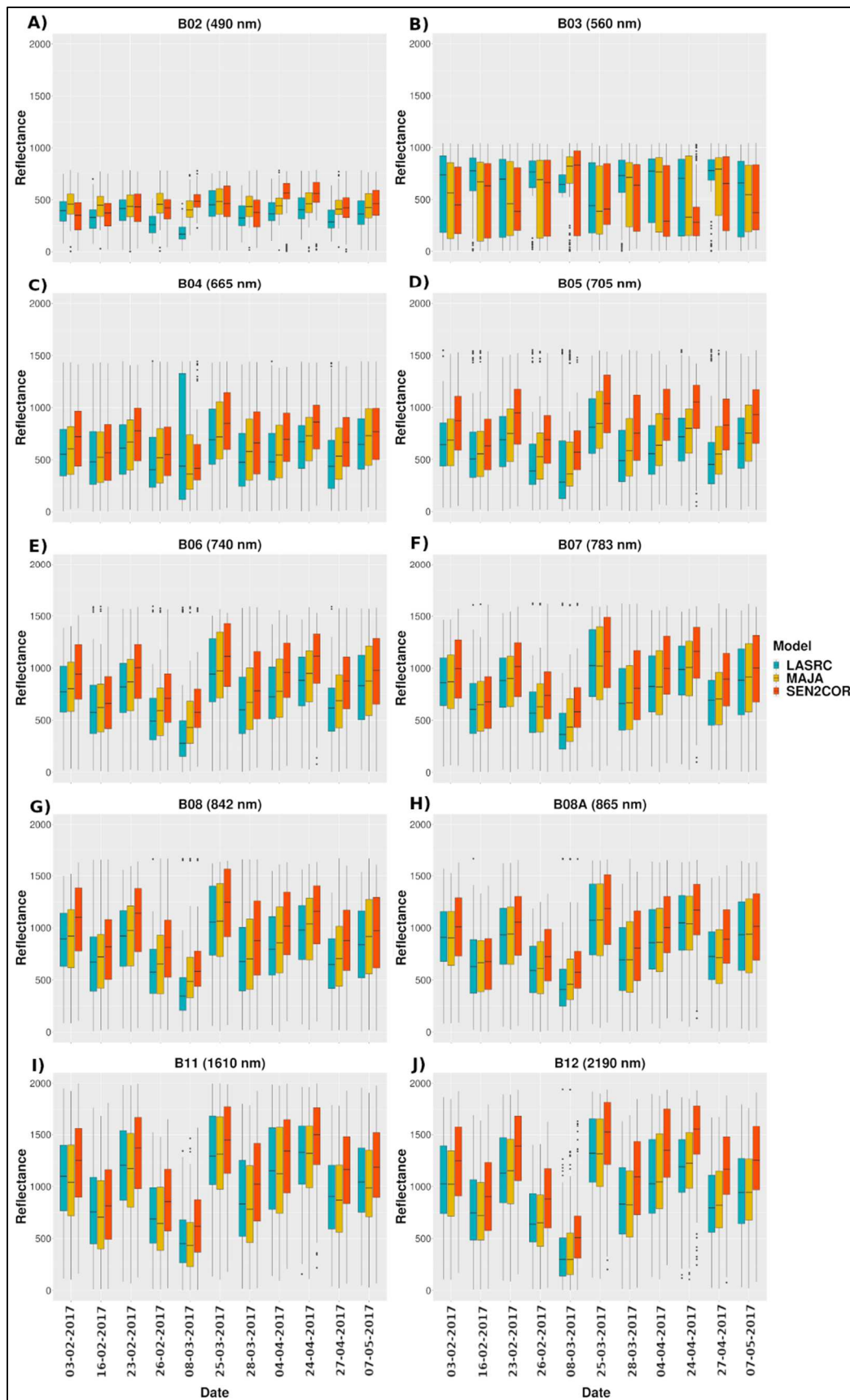


Figure 5: Distribution of BOA reflectance obtained with MAJA (yellow), Sen2Cor (orange) and LaSRC (cyan) at each spectral band and for each date.

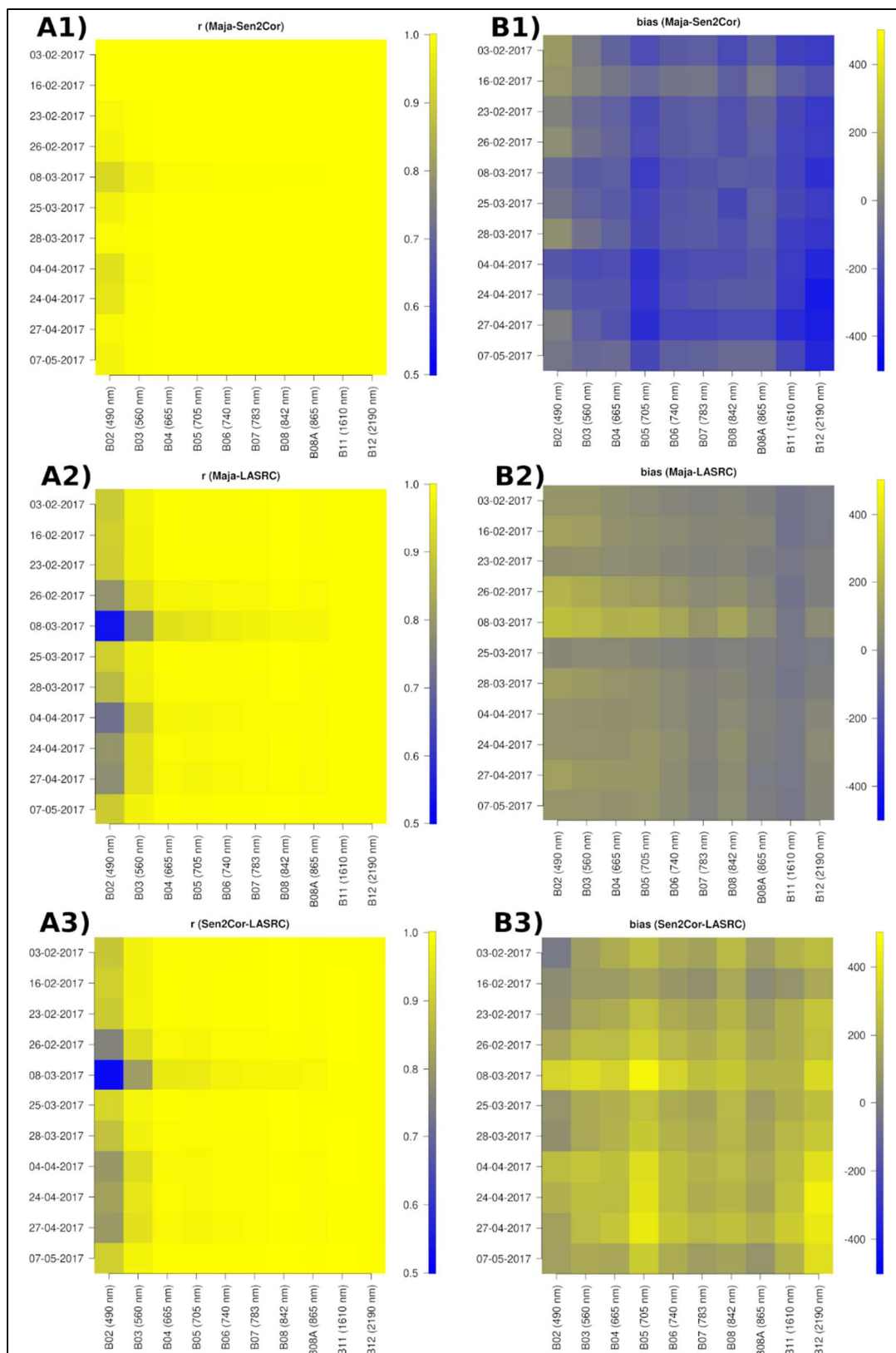


Figure 6: Coefficient of correlation $r_{i,t}$ (A1, A2 and A3) and $Rbias_{i,t}$ (BA, B2 and B3) calculated between the BOA reflectance obtained by MAJA and Sen2Cor (A1 and B1), by MAJA and LaSRC (A2 and B2), and by Sen2Cor and LaSRC (A3 and B3) for the 111 spectra.

Table 1: Proportion of masked cloud surface (%) over each S2 image. Proportion of bare soil (%) over each S2 image obtained by each AC method and the tested combinations of NDVI and NBR2 thresholds. The range (difference between the largest and smallest values of bare soil coverage) was also calculated along both AC methods and S2 dates. Highest ranges of bare soil pixels along both AC methods and S2 dates are highlighted in bold.

		S2 acquisition date (DD-MM-YYYY)											Range along the dates (%)
		03-02-2017	16-02-2017	23-02-2017	26-02-2017	08-03-2017	25-03-2017	28-03-2017	04-04-2017	24-04-2017	27-04-2017	07-05-2017	
% of masked cloud surface		0	0	0	0	3	0	0	0.3	0	1.4	0	
NDVI < 0.25	% of bare soil with MAJA	66.45	74.5	75.18	78.7	71.78	70.43	72.16	68.42	67.69	68.38	65.03	13.67
	% of bare soil with Sen2Cor	60.48	67.85	74.6	77.42	78.97	70.59	70.28	76.4	73.96	71.97	68.89	18.49
	% of bare soil with LaSRC	61.01	66.78	72.74	70.09	58.22	69.26	67.64	67.79	66.3	65.92	62.7	14.52
	Range of bare soil along the 3 AC methods (%)	5.97	7.72	2.44	8.61	20.75	1.33	4.52	8.61	7.66	6.05	6.19	
NDVI < 0.3	% of bare soil with MAJA	77.9	83.4	83.8	85.9	80.5	80.8	81.8	79.8	79.1	78.6	76.3	9.6
	% of bare soil with Sen2Cor	73.4	79.9	83.5	85.6	85.9	81.1	80.9	85.1	83.3	81.3	79.5	12.5
	% of bare soil with LaSRC	75.2	79.1	82.7	82	75.1	80.9	80.2	80.5	78.7	79.2	75.3	7.6
	Range of bare soil along the 3 AC methods (%)	4.5	4.3	1.1	3.9	10.8	0.3	1.6	5.3	4.6	2.7	4.2	
NDVI < 0.35	% of bare soil with MAJA	84.36	88.44	88.56	90.03	85.37	86.69	87.27	86.28	85.58	84.77	83.22	6.81
	% of bare soil with Sen2Cor	80.97	86.12	88.4	90.15	89.58	87.02	86.77	90.13	88.74	86.9	85.92	9.18
	% of bare soil with LaSRC	82.71	85.27	87.99	87.85	83.81	87	86.67	87.29	85.82	86.46	82.93	5.28
	Range of bare soil along the 3 AC methods (%)	3.39	3.17	0.57	2.3	5.77	0.33	0.6	3.85	3.16	2.13	2.99	
% of bare soil with MAJA		23.69	20.73	27.24	17.57	5.11	40.7	26.07	22.37	21.42	18.42	15.34	35.59

NDVI < 0.3 and NBR2 < 0.09	% of bare soil with Sen2Cor	33.89	35.86	44.66	32.25	14.81	55.19	41.83	44.34	51.25	38.96	44.38	40.38
	% of bare soil with LaSRC	20.22	17.26	21.82	10.42	2.08	37.83	19.17	15.16	11.83	9.65	8.53	35.75
	Range of bare soil along the 3 AC methods (%)	13.67	18.6	22.84	21.83	12.73	17.36	22.66	29.18	39.42	29.31	35.85	
NDVI < 0.3 and NBR2 < 0.12	% of bare soil with MAJA	58.51	57.64	64.29	57.21	17.21	73.28	64.02	55.51	59.32	54.8	50.48	56.07
	% of bare soil with Sen2Cor	62.82	67.21	73.72	69.59	36.37	77.75	72.56	73.38	77.52	70.72	72.09	41.38
	% of bare soil with LaSRC	55.24	52.96	59.46	47.06	10.15	71.79	57.71	48.09	47.64	42.63	39.64	61.64
	Range of bare soil along the 3 AC methods (%)	7.58	14.25	14.26	22.53	26.22	5.96	14.85	25.29	29.88	28.09	32.45	
NDVI < 0.3 and NBR2 < 0.15	% of bare soil with MAJA	74.78	78.25	80.58	79.76	40.01	80.35	79.75	74.97	77.06	75.21	72.43	40.57
	% of bare soil with Sen2Cor	72.51	78.5	82.55	83.55	63.42	80.87	80.17	83.84	82.75	80.05	78.9	20.42
	% of bare soil with LaSRC	72.13	74.04	78.71	74.72	27.63	80.29	77.53	73.13	73.89	72.1	68.22	52.66
	Range of bare soil along the 3 AC methods (%)	2.65	4.46	3.84	8.83	35.79	0.58	2.64	10.71	8.86	7.95	10.68	

Table 2: Proportion of common bare soil surface (%) for all S2 images corrected by MAJA, Sen2Cor and LaSRC. Number of collected topsoil samples located over these common bare soil pixels and statistics calculated for each dataset of topsoil samples. These values were calculated for each tested combinations of NDVI and NBR2 thresholds.

	Statistics on collected topsoil samples located over common bare soil pixels						
	% of common bare soil	Number of samples	min (g kg⁻¹)	max (g kg⁻¹)	mean (g kg⁻¹)	Standard deviation (g kg⁻¹)	skewness (g kg⁻¹)
NDVI < 0.25	40.0%	84	58	592	212	117	8.7
NDVI < 0.3	57.2%	111	58	592	224	120	7.8
NDVI < 0.35	67.3%	122	58	592	228	190	7.0
NDVI < 0.3 and NBR2 < 0.09	1.2%	2	58	64	61	3.9	0
NDVI < 0.3 and NBR2 < 0.12	7%	12	58	449	147	109	18.0
NDVI < 0.3 and NBR2 < 0.15	21%	47	58	592	228	119	7.0

Table 3: Validation performances for the estimations of soil clay content obtained from the MLR models built from MAJA- Sen2Cor- and LaSRC-corrected S2 images. Best performances in terms of both R^2_{cv} and $RMSE_{cv}$ for each date are highlighted in bold.

			S2 acquisition date (DD-MM-YYYY)										
			03-02-2017	16-02-2017	23-02-2017	26-02-2017	08-03-2017	25-03-2017	28-03-2017	04-04-2017	24-04-2017	27-04-2017	07-05-2017
NDVI < 0.25	MAJA	R^2_{cv}	0.62	0.63	0.62	0.61	0.52	0.74	0.71	0.62	0.80	0.73	0.72
		$RMSE_{cv}$	7.77	7.56	7.82	7.66	8.74	6.44	6.94	7.86	5.68	6.56	6.47
		MAE	5.55	5.84	5.97	5.72	6.88	5.02	5.45	6.29	4.68	5.23	5.22
	Sen2Cor	R^2_{cv}	0.62	0.63	0.62	0.60	0.50	0.76	0.72	0.62	0.80	0.74	0.72
		$RMSE_{cv}$	7.86	7.68	7.77	7.81	8.93	6.34	7.00	7.74	5.70	6.52	6.59
		MAE	5.54	5.89	5.88	5.71	7.09	4.94	5.48	6.21	4.70	5.20	5.30
	LaSRC	R^2_{cv}	0.63	0.62	0.62	0.59	0.50	0.75	0.72	0.63	0.77	0.72	0.68
		$RMSE_{cv}$	7.61	7.58	7.61	7.91	8.94	6.21	6.84	7.64	6.16	6.50	6.99
		MAE	5.39	5.66	5.85	5.99	7.09	4.79	5.47	6.02	5.07	5.35	5.73
NDVI < 0.3	MAJA	R^2_{cv}	0.62	0.65	0.64	0.65	0.49	0.71	0.66	0.51	0.72	0.68	0.68
		$RMSE_{cv}$	7.57	7.29	7.44	7.34	8.77	6.60	7.19	8.56	6.50	6.98	6.89
		MAE	5.67	5.57	5.62	5.41	6.98	5.18	5.68	6.79	5.17	5.53	5.73
	Sen2Cor	R^2_{cv}	0.61	0.64	0.61	0.64	0.45	0.71	0.67	0.48	0.71	0.68	0.67
		$RMSE_{cv}$	7.68	7.35	7.54	7.38	9.15	6.63	7.15	8.73	6.56	6.97	6.98
		MAE	5.70	5.63	5.66	5.40	7.30	5.20	5.62	6.89	5.24	5.48	5.84
	LaSRC	R^2_{cv}	0.62	0.62	0.62	0.64	0.43	0.71	0.66	0.55	0.71	0.68	0.64
		$RMSE_{cv}$	7.43	7.31	7.38	7.44	9.26	6.38	7.17	8.20	6.62	6.70	7.10
		MAE	5.50	5.54	5.61	5.51	7.40	4.94	5.68	6.44	5.34	5.39	5.88
NDVI < 0.35	MAJA	R^2_{cv}	0.63	0.64	0.61	0.65	0.51	0.74	0.71	0.55	0.75	0.70	0.70
		$RMSE_{cv}$	7.47	7.42	7.59	7.42	8.51	6.32	6.72	8.27	6.14	6.68	6.79
		MAE	5.59	5.60	5.69	5.59	6.76	4.91	5.28	6.55	4.79	5.22	5.61
	Sen2Cor	R^2_{cv}	0.62	0.63	0.62	0.62	0.47	0.73	0.70	0.51	0.75	0.70	0.69

